

Stable Boundary Layers

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SBL over land and water & ice Stockholm University Ζ Ug LLJ Weakly stable - strongly stable longlived Shallow O(m - 100m) T & U Weak mixing – long timescales Non-local in space Non-stationary Intermittent Gravity waves, horizontal gradients, slopes, surface interaction, low-level jet (LLJ), radiation, ..._



SBL in models



$$\overline{w'c'} = -K \frac{\partial C}{\partial z}$$

$$K = \left| \frac{\partial U}{\partial z} \right| l^2 F_{m,h}(Ri)$$

$$l : length \ scale$$

$$Ri = \frac{g}{\theta} \frac{\partial \theta}{\partial z} \left/ \left| \frac{\partial U}{\partial z} \right|^2$$

NWP models need a long tail formulation to get the synoptic scale right (Louis et al. 1981)

Stability functions for momentum



SBL in NWP models





Fig. 4. Difference in 2-m temperature (°C) averaged over January 1996 between simulations with two different stability functions in the ECMWF model (Viterbo et al. 1999; Beljaars 2012). (left) Impact in the 1994 version of the ECMWF model and (right) impact of the same change in the 2011 model version. Color range in legend indicates temperature differences between -4° and 10° C.

Global forecast model Stability functions affect the large scale forecast scores





ECMWF IFS Courtesy A. Beljaars

The ageostrophic flow





Deeper PBL gives larger friction velocity and larger total drag Cross-isobaric angle provides a measure the of ageostrophic flow

PBL over ocean







Mean first guess departure (ECMWF model – observation, ASCAT) of surface wind direction in degrees

Improves a bit when short-tail (MO) functions are used

Sandu et al., 2013





Cuxart et al.2006; Beare et al. 2006; Svensson and Holtslag 2009

GABLS1 PBL scheme intercomparison





Holtslag et al. 2013

GABLS

The ageostrophic flow





Experiments in CAM5.3 10-year AMIP-type experiments





1 degree resolution & 30 vertical layers

PBL scheme: Bretherton and Park, (2009)

Land model use long-tail functions

Default CAM5 (**CONTROL**) No atmospheric turbulence when Ri > 0.19

Lindvall, Svensson and Caballero, 2016



versions have specially for Atlantic sector

Control is closer to observations

Control: enhanced surface roughness based on subgrid scale terrain variability Short tail: Critical Ri of 0.19, Longtail for surface fluxes PBL Long tail: Only in PBL (and surface) Long tail: SBL and above (and surface)

Lindvall, Svensson and Caballero, 2016



PBL in transition ...



The atmosphere is always turbulent











Monin-Obukhov theory assumes a constant flux layer What about directional divergence? Aligned stress and wind vectors?





T_{2m} Short tail vs long tail





Jan 2011 Forecast analysis

Jan 2011 Short tail – long tail

> Doubling the **skin conductivity** on top of short tail increases the temperature with up to 2 degrees again

> > Sandu et al., 2013

Concluding remarks



- Improved understanding and modeling of SBL is important at all latitudes and all seasons
- Theory builds on horizontal homogeneity and stationarity PBL is always in transition
- Model development is done with all processes interacting
- Observations are needed for model development, most useful if all important processes are observed e.g. SHEBA, CASES99
- Understand how momentum is extracted from the atmosphere (kinetic and thermal energy conservation)
 - 1. Structure of the entire SBL and the free flow
 - 2. Flux divergence near the surface
 - 3. Horizontal and vertical heterogeneity
 - 4. Coupling with the surface