Physical Processes in Convection: Chemistry

Ken Pickering University of Maryland

With input from Ron Cohen, Laura Pan, Eric Apel, John Hair, Rich Ferrare

Outline

- Chemical tracer transport in convection
- Lightning NO_x production
- Effects of convection on ozone
 - In-cloud
 - Downwind
 - Wrapping of stratospheric air
- Convective transport of short-lived halogen species
- Wet scavenging of soluble gases
- New particle formation in convective outflow
- Pyroconvection

DEEP CONVECTIVE CLOUDS & CHEMISTRY (DC3)

- **Field campaign** to **simultaneously** use extensive ground & airborne instrumentation to investigate midlatitude deep convection
 - May-June 2012
 - 3 sampling regions
- Focus on storm behavior, lightning, surface emissions, LNO_x, anvil chemistry, and scavenging of chemical species
- DC3 objectives:
 - Investigate active convection and its outflow post-event (12-48 hrs)
 - Study impact on composition and chemistry of upper troposphere



The CONTRAST Experiment (Guam, Jan-Feb 2014) An example of how model, satellite and in situ observations work together to accomplish the science goals

Science Question:

• How important is halogen chemistry to ozone destruction and the associated chemistry-climate interaction?

NCAR Model:

- CAM-chem (CCM) predicts 25% of the chemical ozone loss in the tropical UT is represented by halogen chemistry.
- Halogen chemistry is important to stratospheric ozone depletion, but the injection of halogens into the stratosphere is poorly known; convective lifting of VSLS into TTL can only be provided by in situ observations
- The Field Campaign- CONTRAST "CONvective TRansport of Active Species in the Tropics"
- Unprecedented halogen data from ocean surface to the tropopause
- CAM-chem was run in forecast mode to provide chemical forecast for the flights
- Post-campaign comparison in progress





NASA Global Hawk



Chemical Tracers as Diagnostics of Convective Transport

- Trace gases can be used to illustrate the magnitude and direction of transport within a convective cloud.
- Use gases that have chemical lifetimes much longer than the duration of the convective event.
- Typical examples: CO, O₃, C₂H₆, C₃H₈, CH₃I
 Additional gases ? Need to sample at numerous altitudes.



~50 ppbv increase due to upward transport Cummings et al., 2017, in prep.



~40 ppbv decrease due to upward transport~2 ppbv loss due to chemistry

May 29-30, 2012 DC3

Tracers Used for Model Evaluation

		Low level inflow		Upper levels			
DC3 Storm				Affected by storm outflow		Unaffected by storm outflow	
		CO	03	СО	03	СО	03
May 21 st	Aircraft	150.5 (±9.6)	71.4 (±3.0)	100.2 (±4.5)	143.4 (±25.2)	75.1 (±3.4)	214.3 (±7.6)
	WRF-Chem	152.5 (±2.2)	61.8 (±2.3)	94.2 (±6.7)	147.3 (±25.2)	79.7 (±0.4)	213.0 (±14.3)
May 29 th	Aircraft	132.3 (±3.1)	32.6 (±0.4)	123.1 (±3.6)	80.0 (±4.8)	104.4 (±5.4)	82.2 (±7.0)
	WRF-Chem	136.3 (±0.3)	44.1 (±3.6)	123.2 (±14.2)	84.7 (±12.9)	96.3 (±3.4)	97.1 (±6.4)
June 11 th	Aircraft	117.5 (±4.3)	33.9 (±3.5)	107.9 (±5.0)	111.1 (±16.0)	72.6 (±3.1)	155.3 (±20.2)
	WRF-Chem	112.0 (±7.8)	45.9 (±4.2)	108.8 (±2.1)	101.4 (±14.4)	69.8 (±0.7)	161.8 (±6.1)



Positive values of Vertical Flux Divergence (VFD) indicate entrainment; negative values indicate detrainment

At mature stage, supercell storm has strongest anvil level detrainment, but MCS is much longer lived storm

Li et al., 2017 submitted

Some Literature Estimates of LNO_x Production Per Flash

Method	Moles NO/flash (Notes)	
Reference		
Theoretical	1100 (CG), 110 (IC)	
Price et al., 1997		
Laboratory	~103	
	Wang et al., 1998	
A/C data, cloud model DeCaria, et al., 20	345-460 (STERAO-A) 05	
A/C data, cloud model al., 2010	590-700 (CRYSTAL-FACE)	Ott et
	500 (Mean midlat. model)	
Ott et al., 2010		
A/C data, cloud model Cummings et al., 2	500 - 600 (Hector) 2013	
A/C data, cloud model Cummings, 2017	82 (DC3 OK supercells)	
Aircraft data Huntrieser et al., 2	70-210 (TROCCINOX) 2008	
Aircraft detest estimate: 2 (Schtrinser et ald	50 moles/f lash3856169811 y9,2 ^{wi} 8)TgN/yr Huntrieser, 2007) – factor of 4 uncertainty	
Aircraft data	70-179 (ΔΜΜΔ)	

LNOx Production in DC3 Storms



Pollack et al., (2016) estimated mean LNOx production per flash in 3 DC3 storms using volume- and flux-based approaches.

Estimates ranged from 107 to 332 moles per flash.

Cummings et al. (2017, in prep.) simulated the 29-30 May DC3 high flash rate supercell storm in Okla. using WRF-Chem and a upward ice flux flash rate prediction scheme.

Several LNOx production scenarios were tested: 82 moles per flash yielded best with anvil aircraft data; LMA: mean flash extent 7.6 km → ~10 moles/km

Relating LNO_x Production to Lightning Characteristics



Correlations based on data from 23:40-03:00 UTC (red +)

Cummings et al., 2017, in prep.

Need flash rate, extent, energy from LMA

9

Possible Underestimate of LNOx Production per Flash

- Nault et al. (2015) found that methyl peroxy nitrate, pernitric acid, nitrogen pentoxide, and alkyl nitrates are more significant sinks for NOx in convective outflow in DC3 than expected.
- Much more NOx in convective outflow was converted to higher oxides of N.
- Result is 2 3 hour lifetime for NOx in upper tropospheric convective outflow rather than the assumed few days.
- Nault et al. (2017, in prep.) find that reinterpretation of many previous measurements in light of the short NOx lifetime leads to greater estimates of LNOx production per flash (510 moles per flash in tropics and 550 moles/flash in midlatitudes).
- Nault et al. also show GEOS-Chem simulations with updated chemistry support 665 – 700 moles/flash over SE US when constrained by OMI NO2 column satellite data.

Further Investigation of Short NOx Lifetime and Larger LNOx Production/Flash Needed

- Storm penetrating aircraft (A-10) needed to sample NOx produced by lightning within or near storm core, where most lightning occurs.
- Combine the storm penetrating observations with anvil sampling (DC-8 or G-V) at a variety of altitudes and downwind distances to obtain a comprehensive view of NOx within a convective system.
- Will allow verification of the proposed 2-3 hour NOx lifetime and large LNOx production per flash.
- Alternatively, overfly storms with high-altitude aircraft to remotely sense NO2 in storm cores.
- Both methods require detailed flash information from LMAs, GLM, or aircraft.

Lightning NO_x Production During GOES-R Validation Flights March – May 2017 Scott Janz and Matt Kowalewski, NASA/GSFC Ken Pickering and Dale Allen, Univ. of MD

- Geo-CAPE Airborne Simulator (GCAS from GSFC) and Fly's Eye GLM Simulator (FEGS from MSFC) flew on NASA ER-2 during GOES-R Validation Mission.
- Provided a preview of the synergy available when GOES-R Geostationary Lightning Mapper (GLM) and geostationary Tropospheric Emissions Monitoring of Pollution (TEMPO; 2019 or later) are both in orbit.
- Data from validation flights and from these satellite instruments will allow retrieval of co-located flash counts and NO₂ column amounts, which will lead to improved estimates of lightning NO_x production per flash.



FEGS: Rich Blakeslee Mason Quick NASA/MSFC

Strong correlation of NO₂ columns and lightning optical pulses

Effects of LNO_x on Ozone



Liaskos et al. (2015) showed factor of 4 uncertainty in LNOx production leads to 40-60% uncertainty in tropical upper tropospheric O_3 in NASA GEOS-5 model

Effects to consider:

- Short-term loss of ozone inside storm
- Efficient photochemical ozone production in upper tropospheric storm outflow
- Wrapping of stratospheric ozone downward around storm anvil

Need storm penetrating aircraft to examine titration of ozone in storm Need outflow sampling at several downwind transport times/distances Need to determine how frequently stratospheric air wraps around anvil

Ozone Chemistry Within Thunderstorm Clouds



 O_3 loss occurs inside highly electrified clouds due to NO + O_3 reaction in presence of large NO mixing ratios from lightning.

Ott et al. (2007) simulation of EULINOX storm near Munich, Germany using 3-D GCE model with off-line chemistry

Found O_3 loss of up to 9 ppbv (avg. 4 ppbv) in regions with large NOx mixing ratios from lightning and pollution

No O_3 loss due to transport in this case.

DC3 2013 June 21 MCS Flight DC-8 and GV data MM Model comparison



During the DC3 experiment we had the exceptional opportunity to study the outflow from an MCS (mesoscale convective system).

The highly instrumented GV aircraft followed the highly instrumented DC8 aircraft in a daylong study of the outflow from the previous night's MCS.





Flight tracks

 MM agrees generally well with measurements

- Photochemistry produces ozone in outflow (at 10 km altitude) – spikes represent strat intrusion
- TOGA and WAS measured toluene and other VOCs that contribute to formaldehyde (Fried, DFGAS) and ozone (Ryerson, NOAA) formation



GV Measurements: VOCs – *TOGA* (Apel, Hills, Hornbrook, NCAR/ACD; Riemer, U. Miami), O₃ (Campos, NCAR/ACD), Formaldehyde (Fried, NCAR/EOL) *DC8 Measurements*: VOCs - Whole air sampler (Blake, UC-Irvine); O₃ (Ryerson, NOAA), Formaldehyde (Fried, NCAR/EOL) *NCAR Master Mechanism*: Detailed 0-D model – Chemistry – Apel, Lee-Taylor, Madronich (NCAR/ACD)

Airborne O₃ DIAL profiles directly reveal transport of stratospheric ozone near deep convective storms during DC3

Altitude (km)

Stratospheric Transport WRF Model

Profile of ozone and depolarization (355nm) approaching Mesoscale Convective System (MCS)



"Thunderstorms enhance tropospheric ozone by wrapping and shedding stratospheric air" Pan, L. L., et al. (2014), Geophys. Res. Lett., 41, 7785-7790, doi:10.1002/2014GL061921



Guam, Jan-Feb 2014

Tropical Storms – a mechanism for impacting the oxidizing capacity of the upper troposphere

Short-lived boundary layer species observed at the tropical upper troposphere on the GV near active convection

CONTRAST

Lifetime: Acetone ≈ weeks Formaldehyde ≈ hours Acetaldehyde ≈ hours



Tropical Convection – a link between oceanic biology and global ozone chemistry

- Collaborative observations from the GV, the Global Hawk, and the BAe146
- First complete measurements of this kind over western Pacific
- Provide quantitative constraints for global chemistry climate models



Wet Scavenging of Soluble Species

- Scavenging efficiencies (SE) between inflow and outflow can be calculated from aircraft observations and from cloudresolved storm/chemistry simulations.
- Amount of soluble gas incorporated into cloud water is dependent on the degree of solubility of the gas (Henry's Law coefficient)
- But, SE is also dependent on the fraction of the gas that is retained on ice upon freezing of cloud water.



Bela et al. (2016) used the WRF-Chem model to determine SE for several soluble gases and the fraction of the gases retained on cloud ice

SE for all soluble species except HNO3 are highly sensitive to the ice retention fraction that is assumed.

Best match aircraft obs: 0% retention: HCHO, H2O2 100% retention: CH3OOH, SO2

Storm-penetrating aircraft needed to make observations of gases, cloud water, cloud ice in storm cores (where freezing is taking place; Can ice particles be collected and later analyzed for gas content?

New Particle Formation in Convective Outflow

- During DISCOVER-AQ field experiment in Maryland (July 2011) Eck et al. (2014) found enhanced aerosol optical depth (AOD) in the vicinity of fair weather cumulus clouds from surface AERONET sun photometer data.
- Surface Micropulse Lidar (MPL) and airborne High Spectral Resolution Lidar (HSRL) also measured large increases in aerosol signal (backscatter, extinction, AOD) at altitudes of fair weather cumulus clouds.
- In situ aircraft observations during vertical profiling also yielded enhanced aerosol scattering and volume after cumulus formation compared with before.
- Humid environment likely enabled easy detection.
- Need further case studies of this process; need new particle speciation; investigations of ability of these aerosols to aid new cloud formation.

AERONET Detection of New Particle Formation



10 14

14 18 10

Tes UTC.

10

22



Eck et al., 2014



HSRL Detection of New Particle Formation



Pyroconvection

- Convective clouds often form in association with biomass fires due to the large quantities of heat and water vapor released.
- Extreme cases of pyroconvective clouds reaching into the stratosphere and/or containing lightning can result from large fires.
- Also plumes from fires can become entrained into otherwise relatively clean thunderstorm events (ARCTAS, DC3, SEAC4RS)
- Needs:

Can composition of pyroconvection be determined only by remote sensing? How often do fire-induced clouds occur?

Summary

 Storm penetrating aircraft (A-10) is needed for : chemical tracer measurements better characterization of lightning NOx determining magnitude of ozone titration due to LNOx better understanding of wet scavenging We won't know much about storm core chemistry without it!

- Role of remote sensing should be expanded for: estimates of LNOx in storm cores investigation of new particle formation composition of pyroconvective clouds
- Need close collaboration with lightning observation and cloud physics measurement communities.