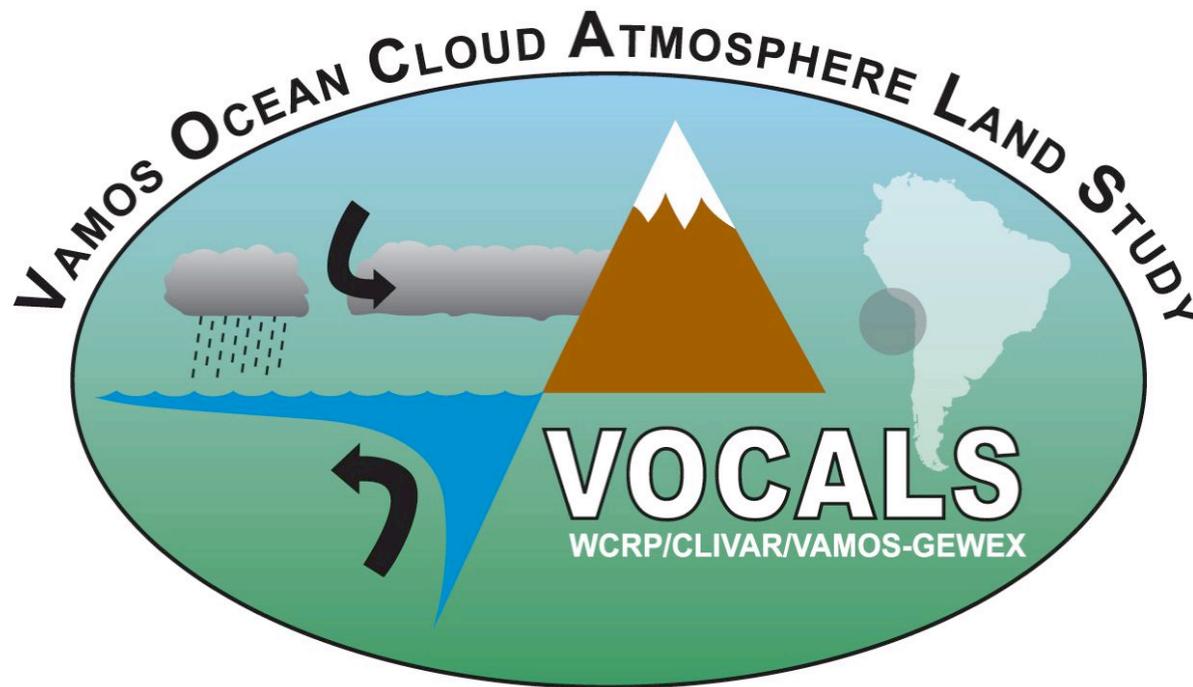
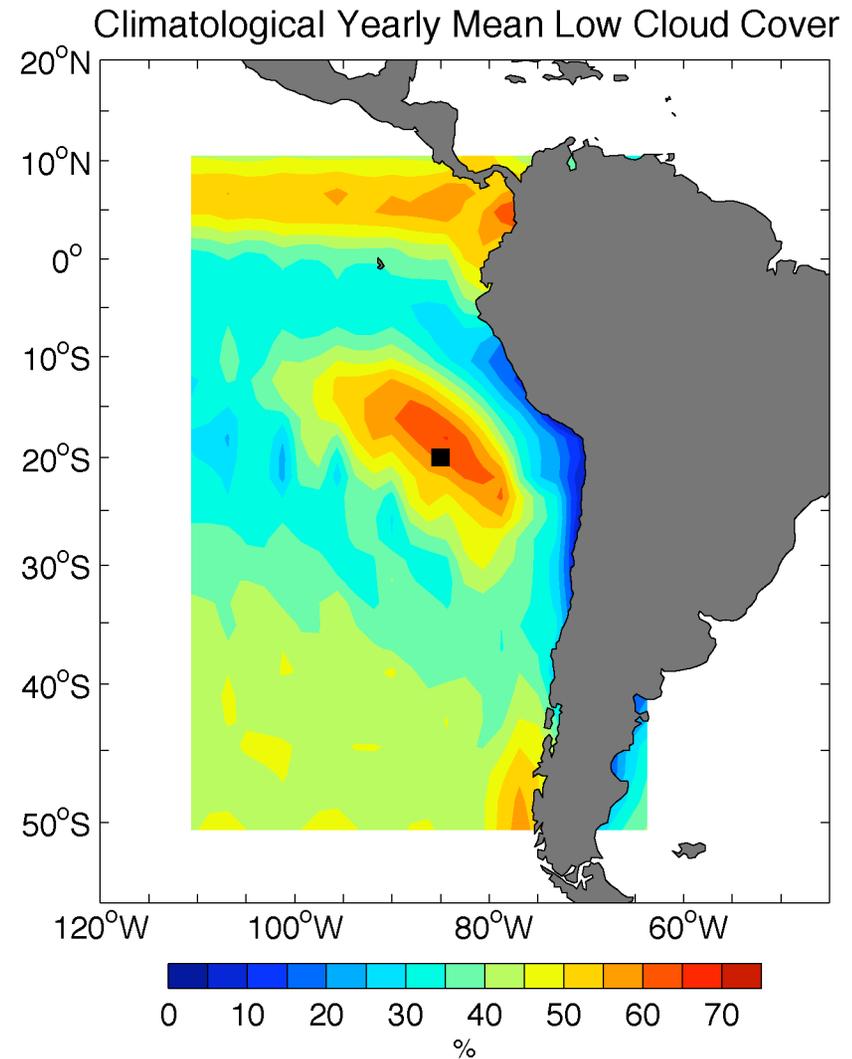


# VOCALS - Ocean Science and Implementation



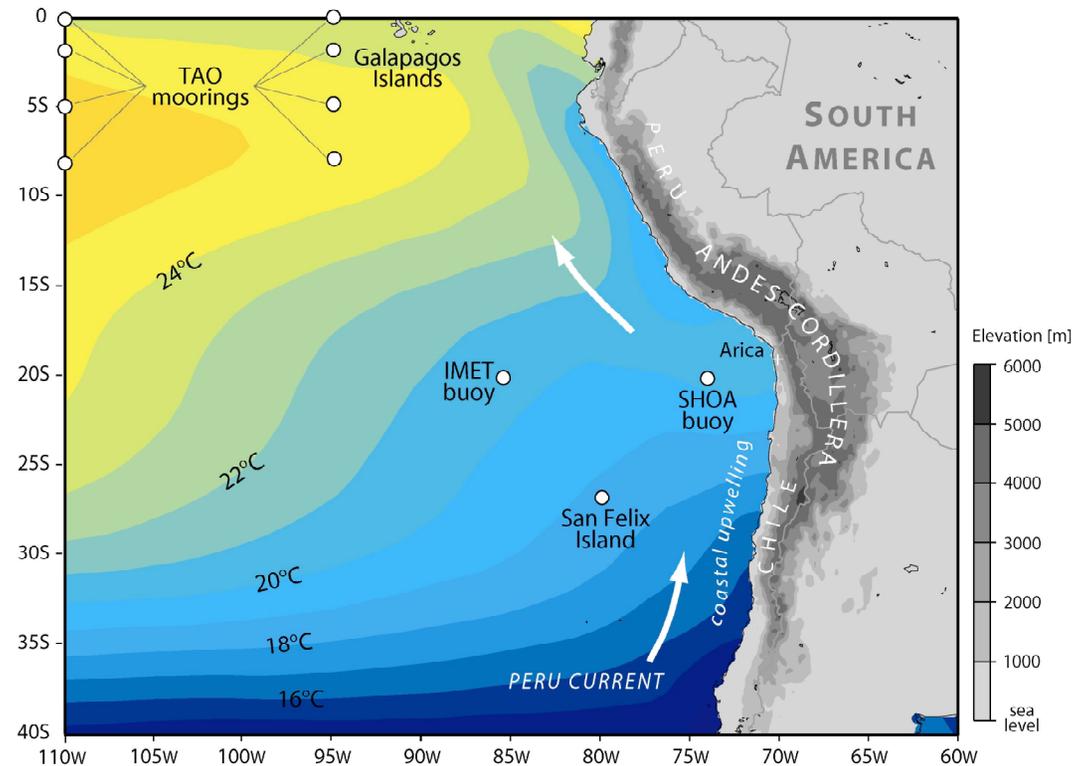
## An eastern boundary

- model biases
- cloud cover
- ocean-atmosphere



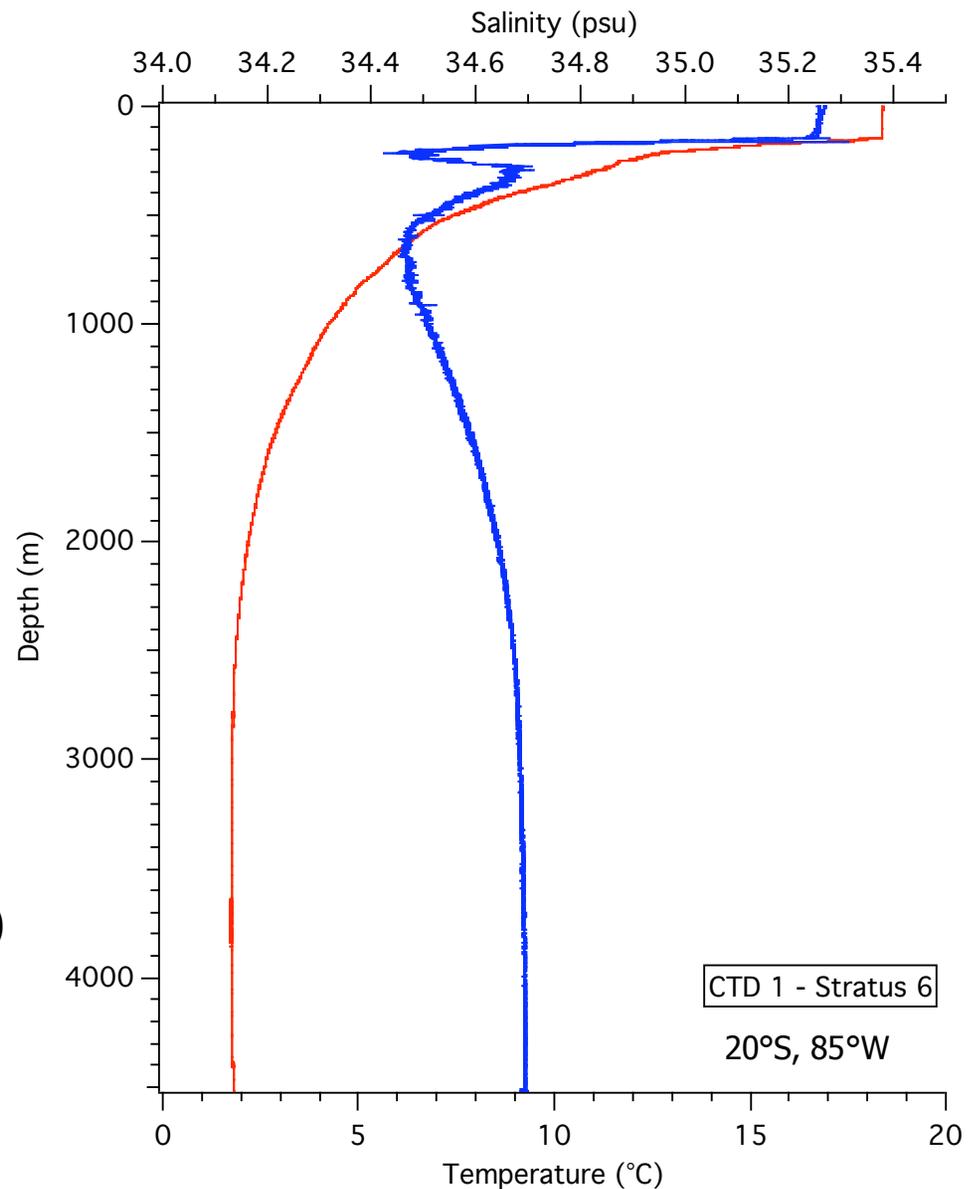
# The ocean setting - the Southeast Pacific (SEP)

- Persistent trade winds, coastal upwelling.
- Trade winds - directionally steady but vary in speed, with periods of low winds
- Low level of synoptic weather systems
- Peru/Chile Current flowing north and northwest.

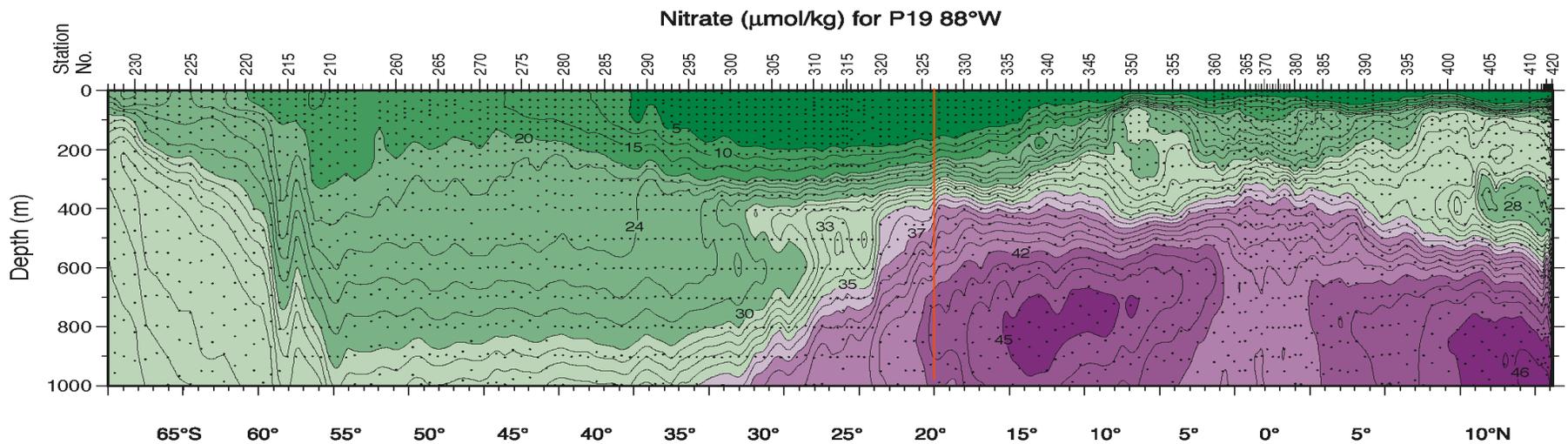
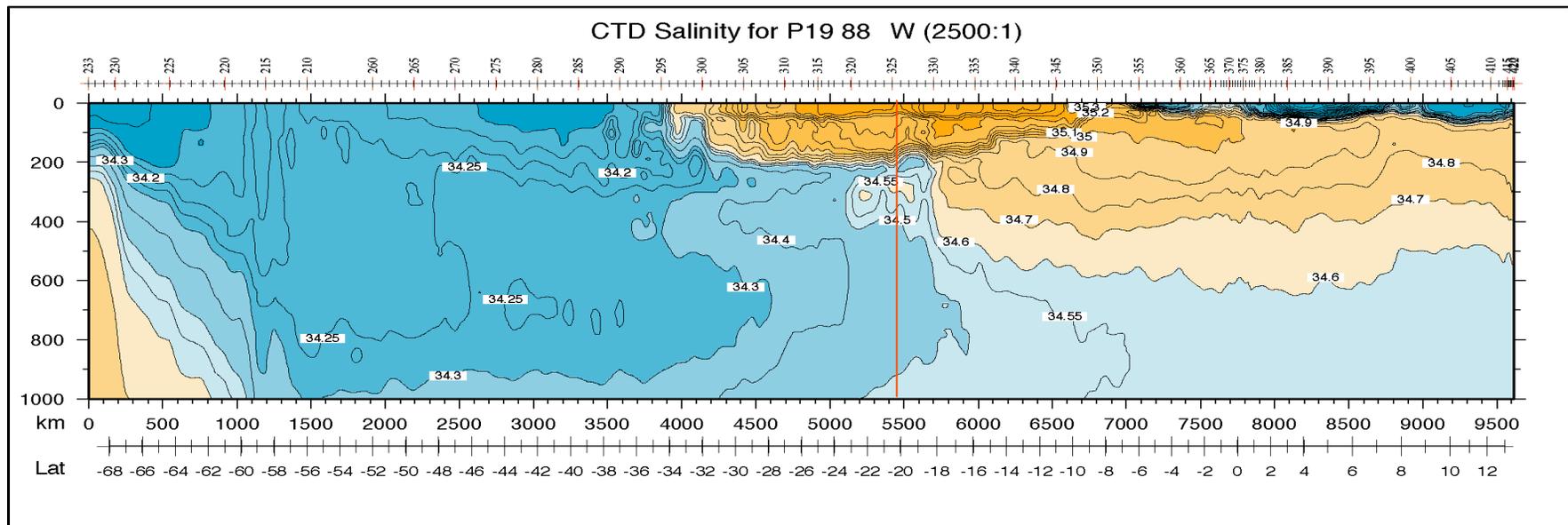


- A strongly evaporative, moderately warmed region producing temperate, salty surface water.
- Fresher water moving in below the surface layer.
- Below that a more saline layer and a second salinity minimum.
- VOCALS' goal of understanding controls on SST sets a focus on the surface layer
- VOCALS partners (Chile, Peru, France - PRIMO) interest is on the oxygen minimum layer below

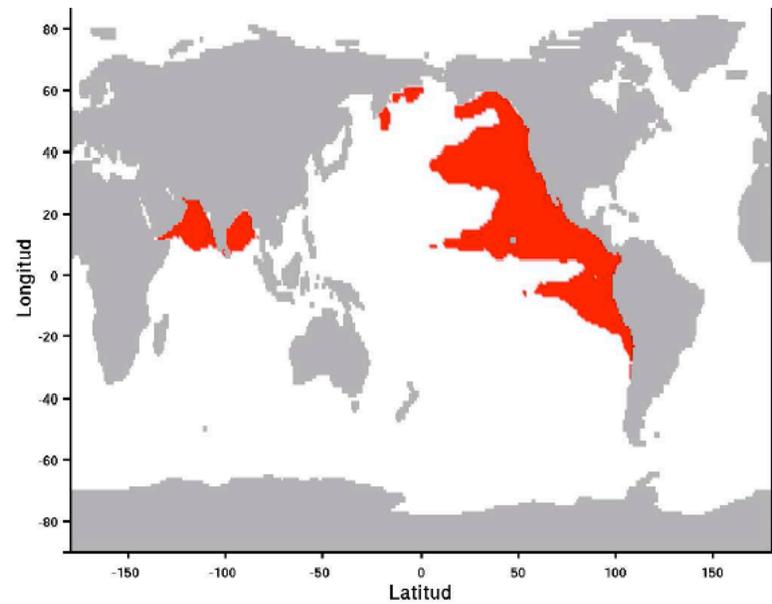
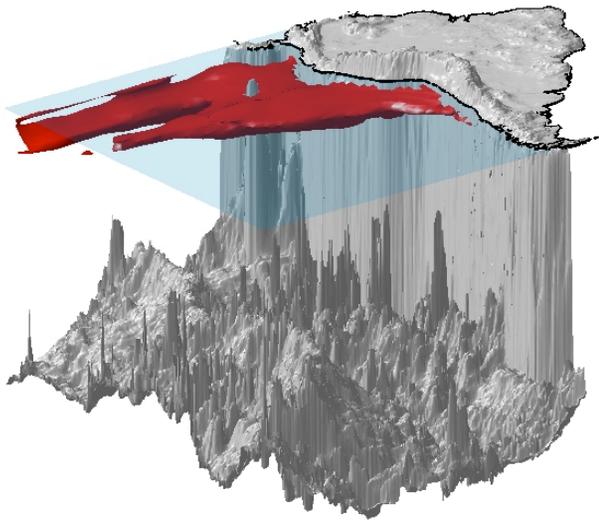
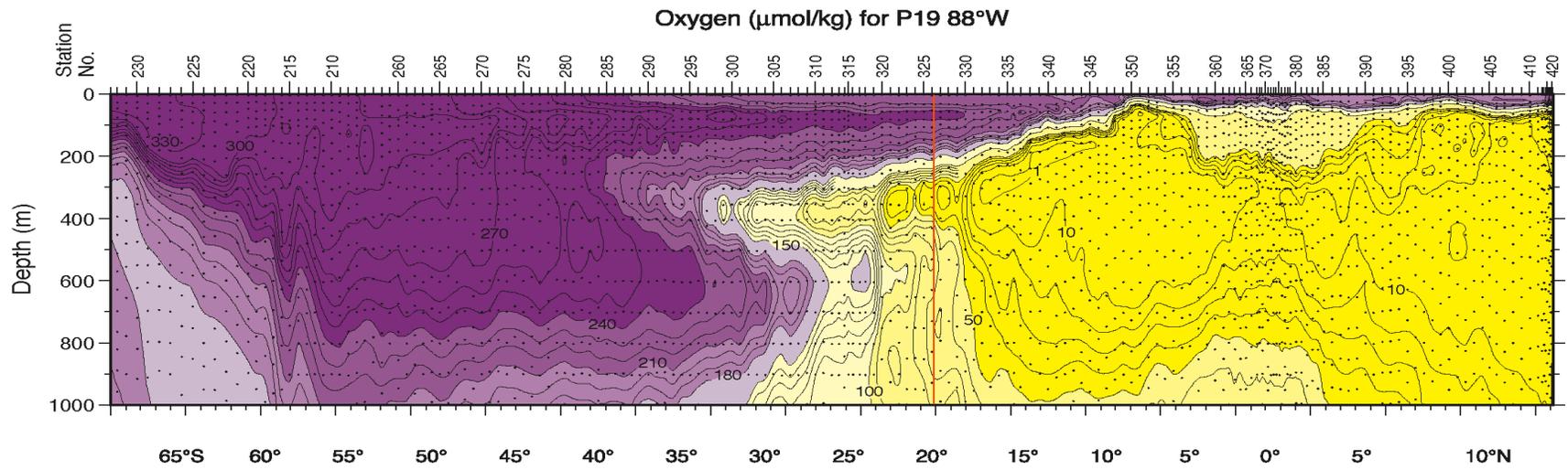
# The ocean setting



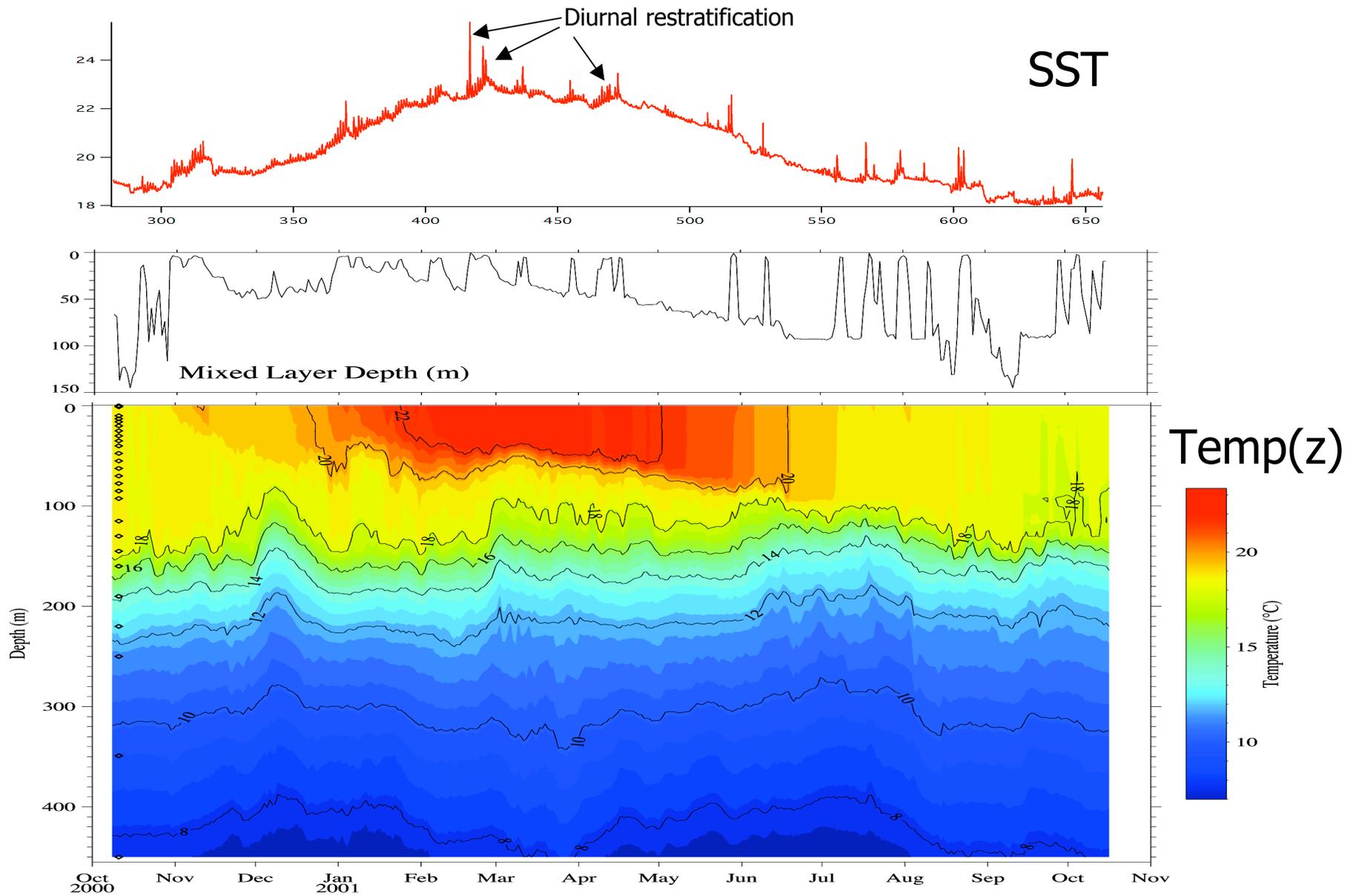
# The ocean setting



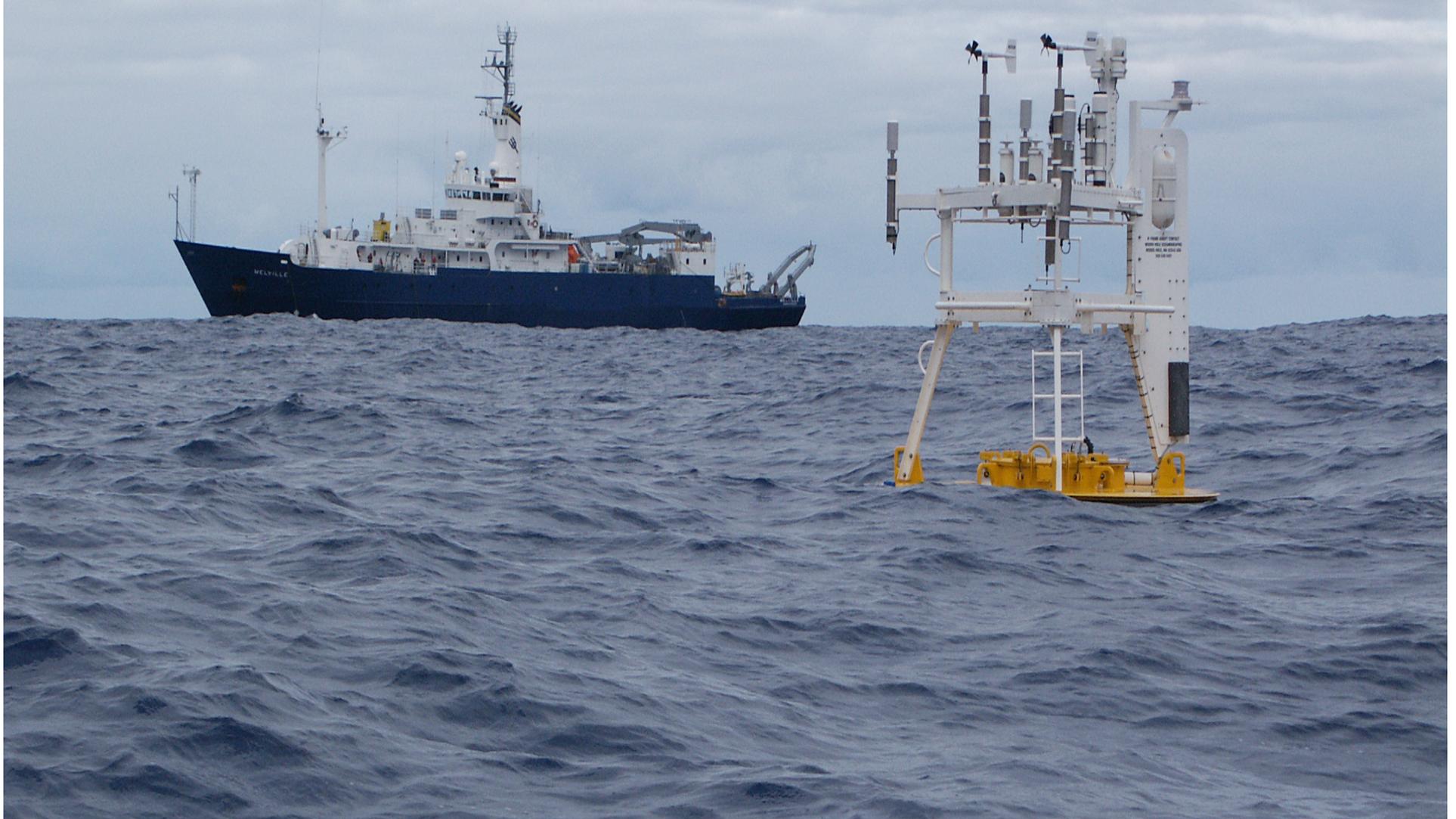
# The ocean setting



# Annual cycle in the upper ocean under the stratus deck



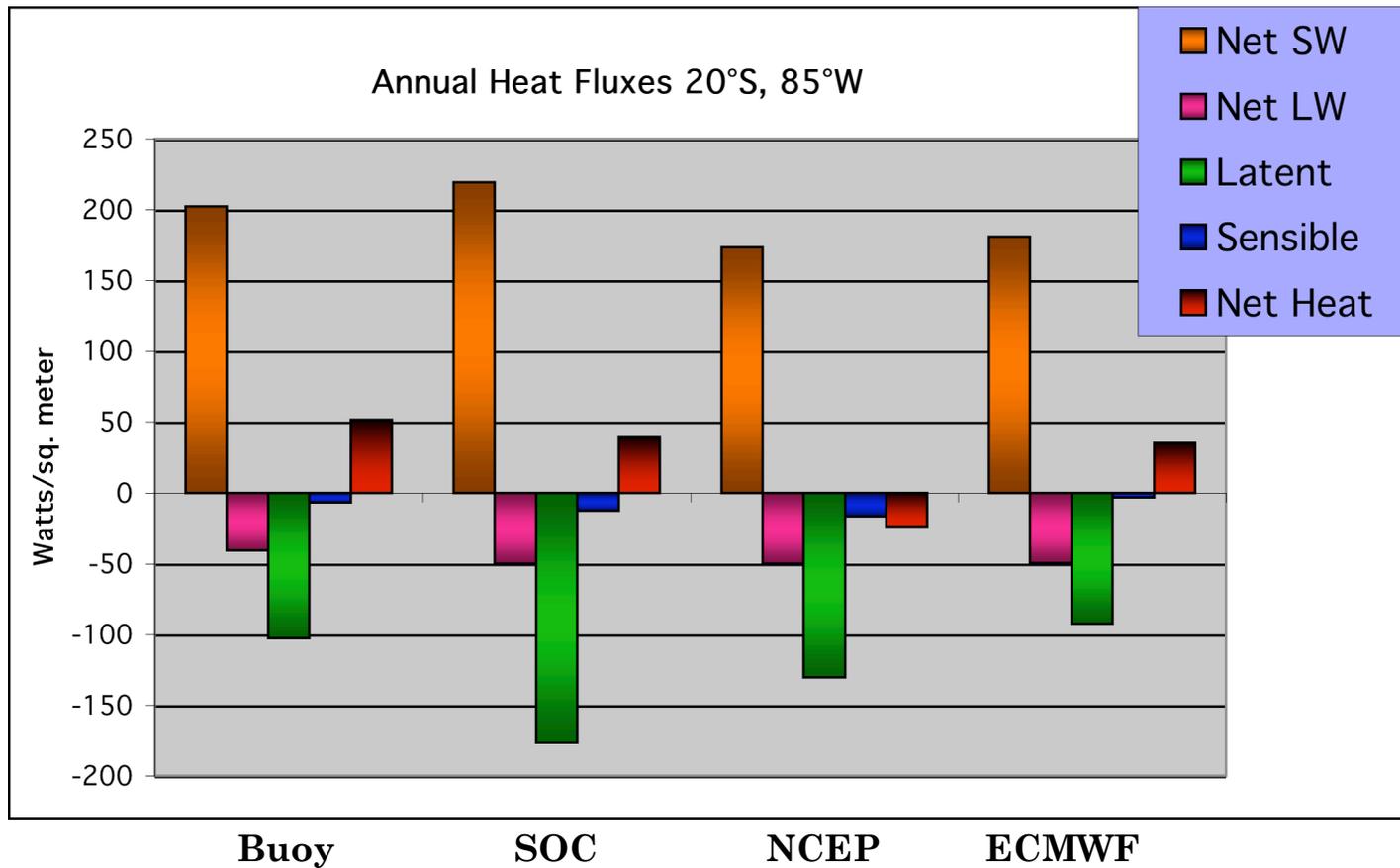
Pilot studies: To what extent are SST and upper ocean structure predictable from surface forcing alone?



Model-based surface fluxes are not accurate, for example, NCEP1 suggests a longer, cooler winter and little net heating of the ocean.

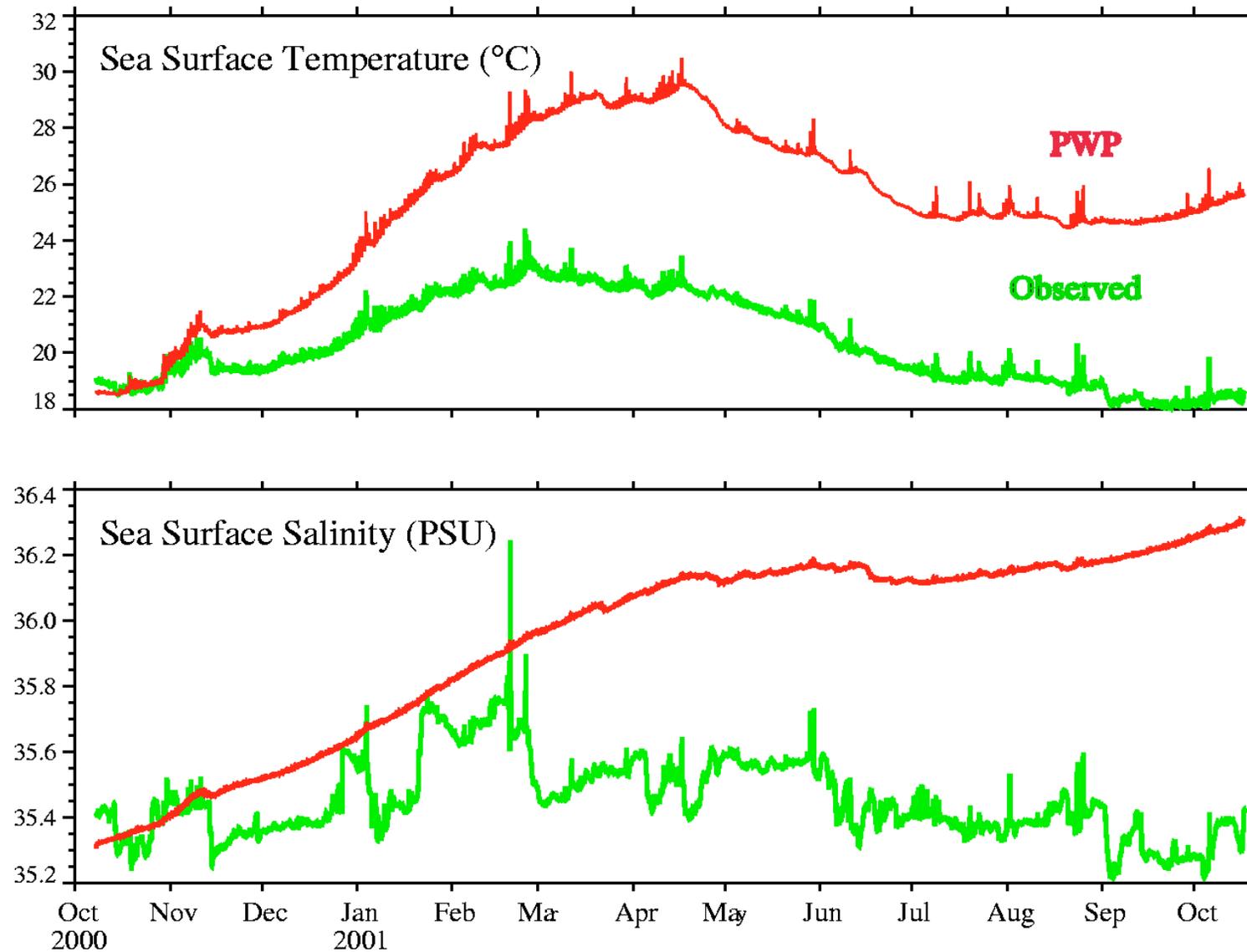
Annual net heat ranges from Buoy - more ocean gain than models, NCEP shows a loss  
NCEP stress 30% stronger

Models: rain (.07 to .3 m yr<sup>-1</sup>) Buoy: no rain to .03 m yr<sup>-1</sup>



Surface forcing from buoy driving a one-dimensional ocean model (PWP) produces a surface layer that is too warm and too salty.

Stratus1 Observed versus PWP (T & S Initialized from CTD cast)



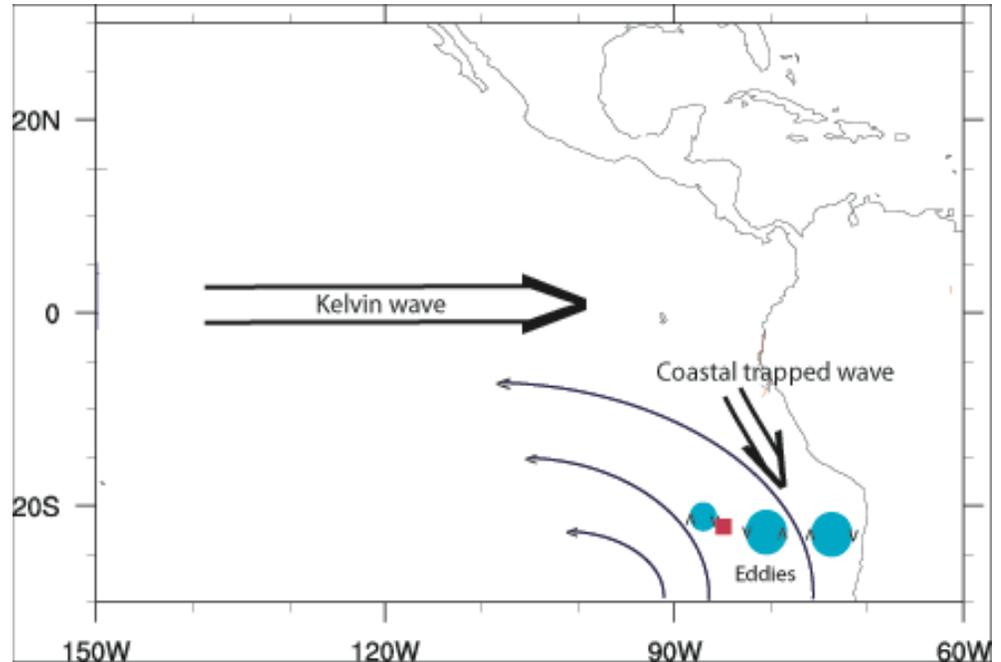
Weller

Additional cooling and freshening is needed. Possible mechanisms:

- Ekman (wind-driven surface layer) transport offshore of coastal water
- Open ocean downwelling/upwelling (Ekman pumping)
- Mixing with low saline water below
- Geostrophic currents (advection)
- Eddy processes, including horizontal transport enhanced vertical mixing

Remote as well as local forcing is possible, possible links to ENSO variability.

- Kelvin waves->coastal waves-> Rossby waves
- Displacement of S Pacific high pressure center



### Integrated Heat Content Equation

$$\int_0^{1\text{year}} \left( \frac{Q_{net}}{C_p \rho_0} - \int_{z_0}^0 \left( u \cdot \nabla T + w_E \frac{\partial T}{\partial z} + \frac{\partial \overline{u'T'}}{\partial x} + \frac{\partial \overline{v'T'}}{\partial y} \right) dz - \kappa_v \frac{\partial T}{\partial z} \Big|_{z=z_0} \right) dt \approx 0$$

Surface Flux

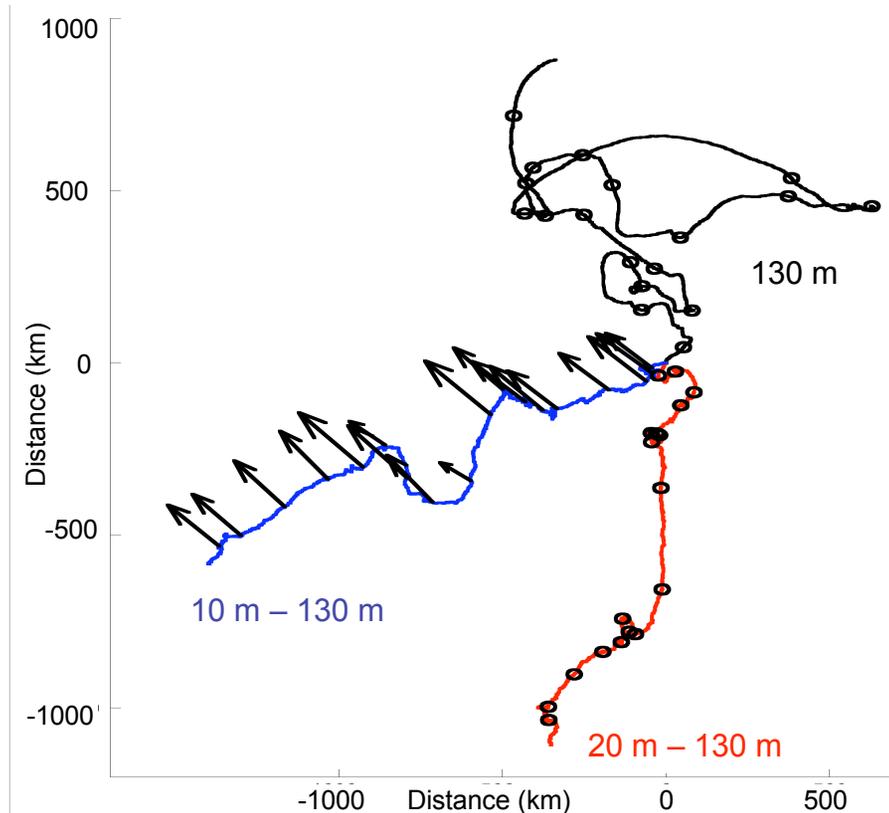
Advection

Ekman  
Pumping

Eddy Flux  
Divergence

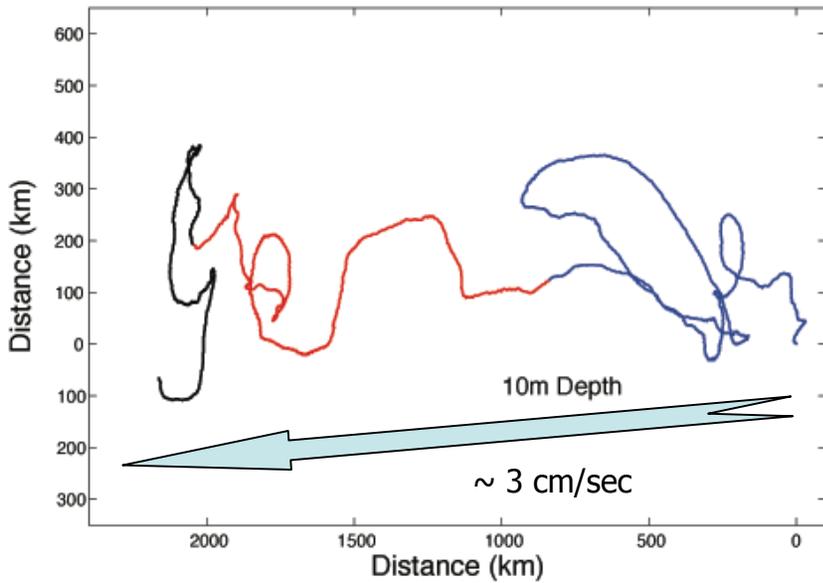
Vertical  
Diffusivity

# Steady Trade Winds to the NW, wind-driven surface flow to the Southwest

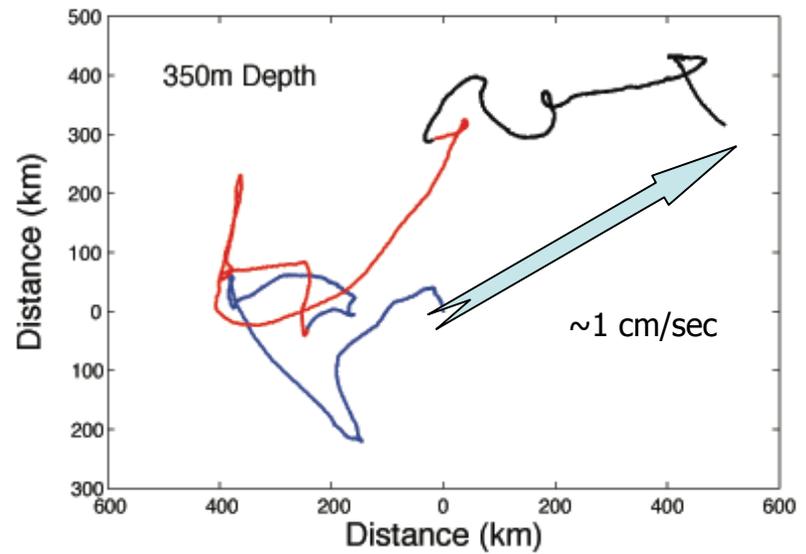


One-year displacements or progressive vector diagrams of velocities at 10 and 20 m relative to that at 130 m, as well as for the velocity at 130 m.

The surface water moves offshore under the influence of the wind.



3 year displacement at 10 m depth,  
a mean of  $\sim 3$  cm/sec

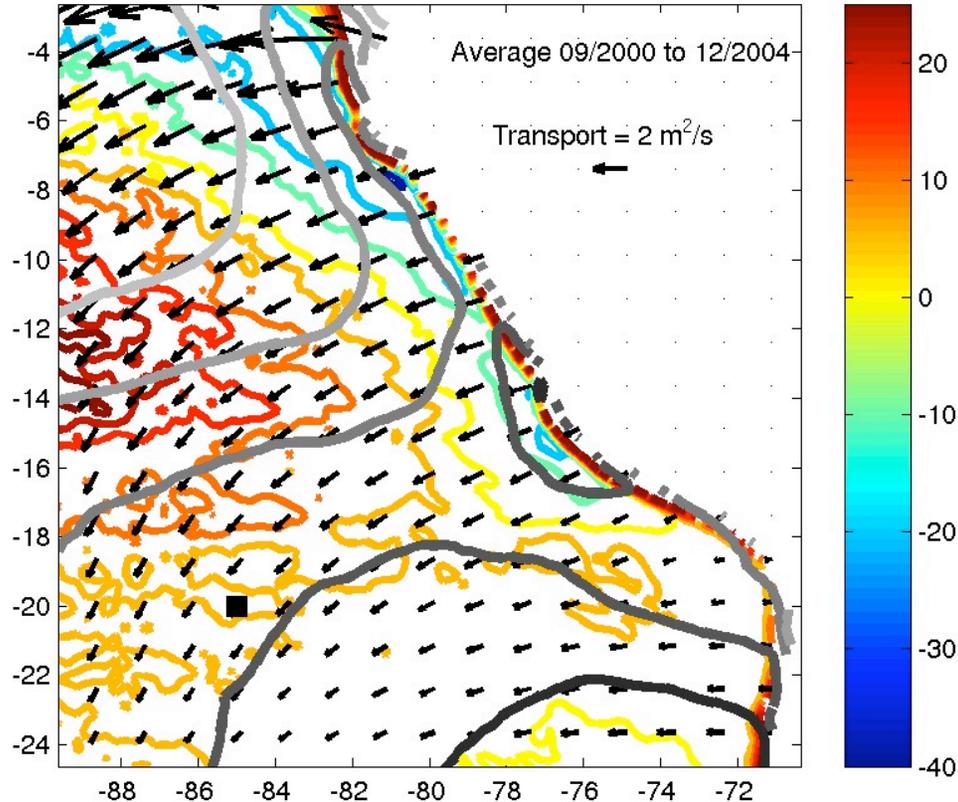


3 year displacement at 350 m depth,  
a mean of  $\sim 1$  cm/sec

In upper thermocline, 1-2 cm/sec annual mean  
Flows to NW.

QuikScat winds and TMI SST fields used to estimate the advective component of heat flux due to Ekman transport across SST gradients. Calculation done for weekly fields and then combined to get an annual average. The steadiness of the winds implies that the mean of the high-frequency product is close to the product of the means.

### Ekman Advection along SST gradients

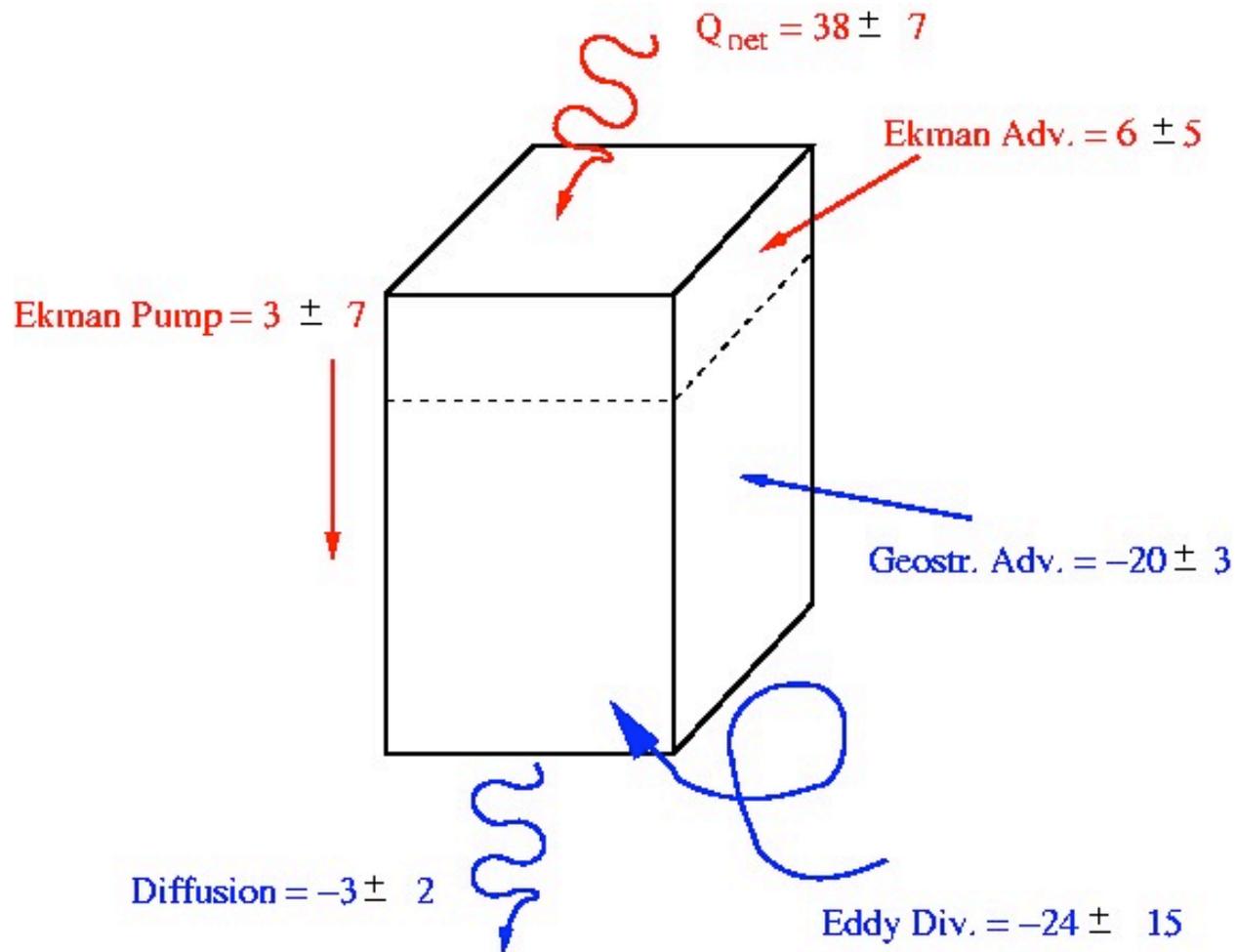


**Color Contours:** Annually averaged component of the heat flux due to advection by Ekman transport

**Gray Contours:** Annually averages SST

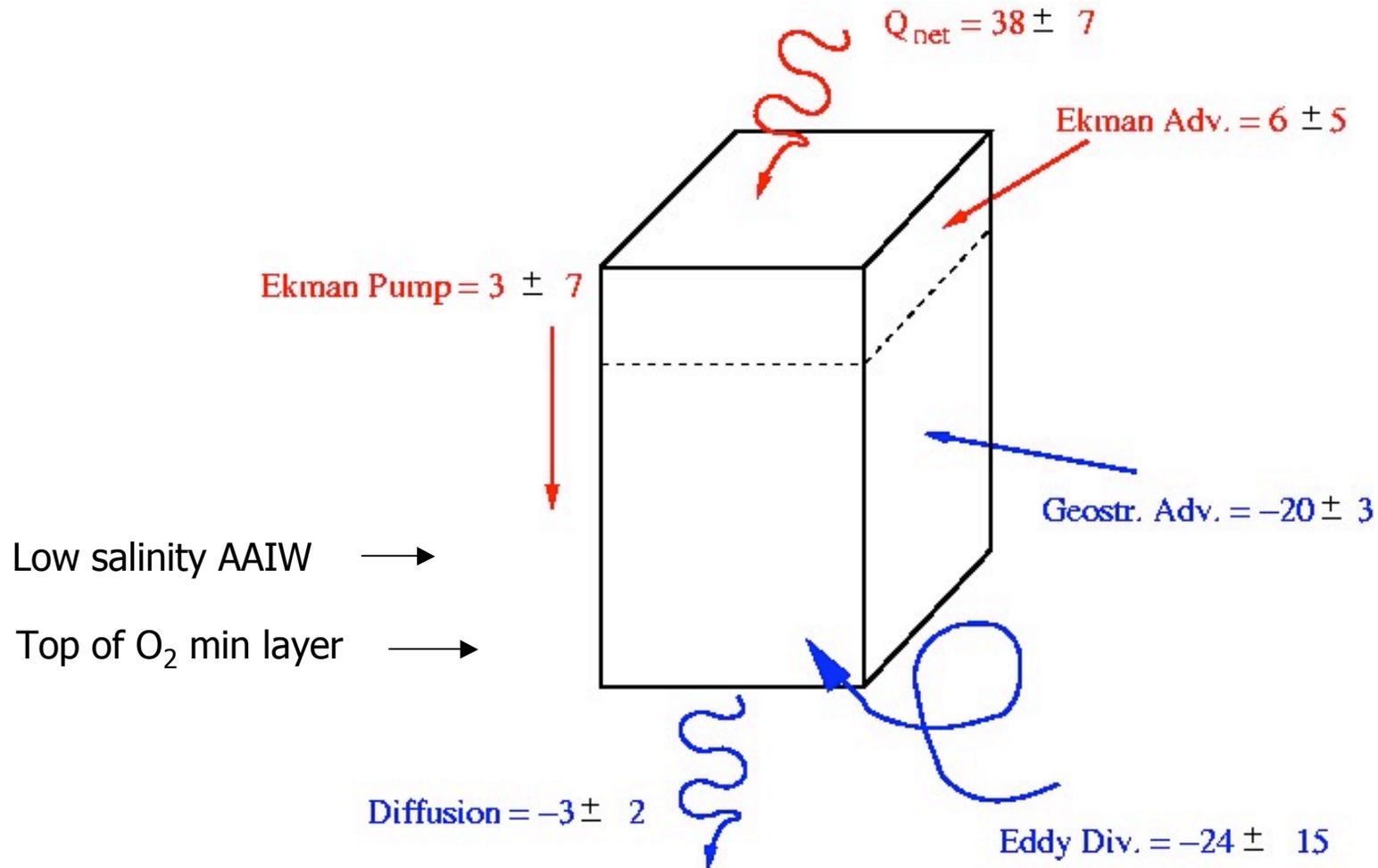
**Arrows:** Annually averaged Ekman transport

**Ekman Advection = 6 +/- 5 W/m<sup>2</sup>**

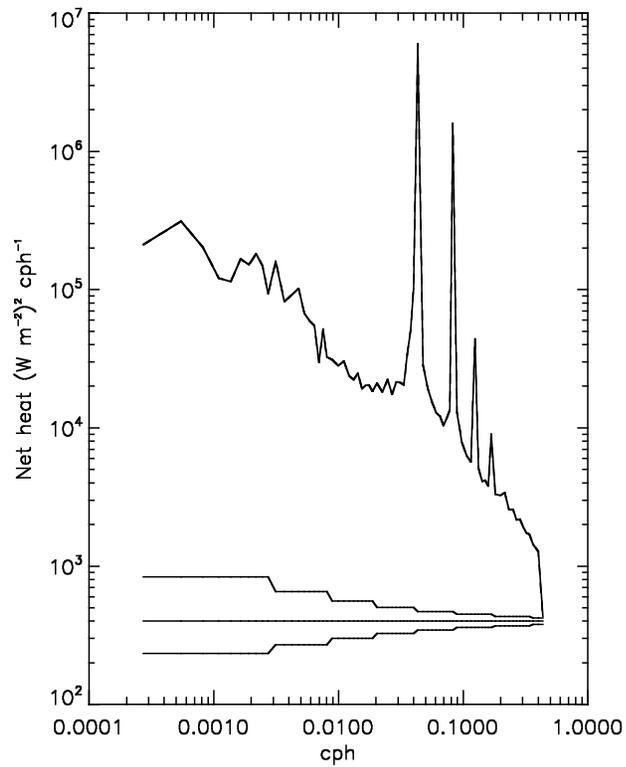


Best estimates of heat balance at Stratus mooring, 20°S, 85°W in a volume down to 250m.

Major terms: surface heating, geostrophic advection, eddy divergence.



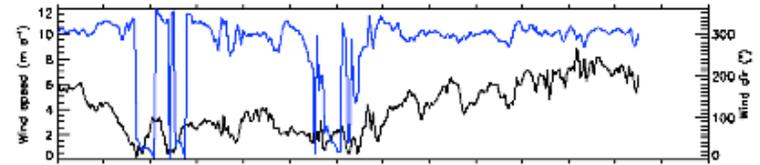
- Geostrophic advection from SW - heat, salt
- Eddy flux from the coast - heat, salt, nutrients
- Vertical mixing between upper thermocline and surface mixed layer
- Vertical exchange between surface layer and oxygen minimum layer



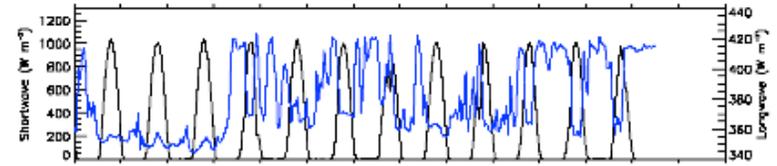
Power spectrum of net heat flux.

Of concern: Diurnal variability in forcing and in response.

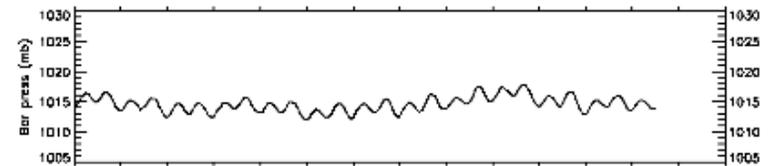
wind



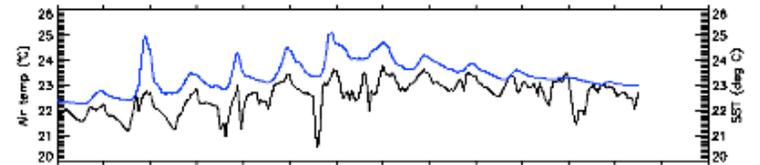
radiation



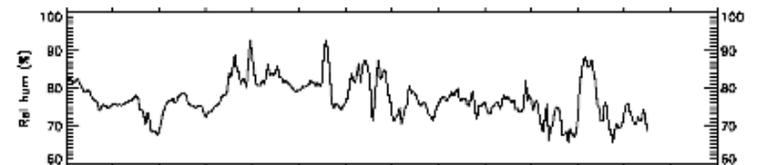
bar pres



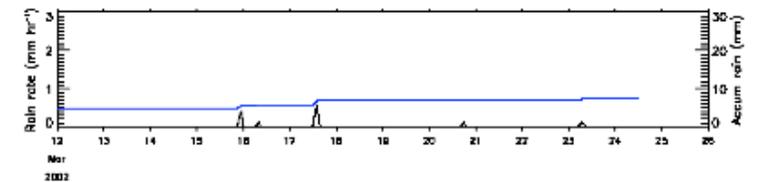
SST  
Air temp



Rel hum

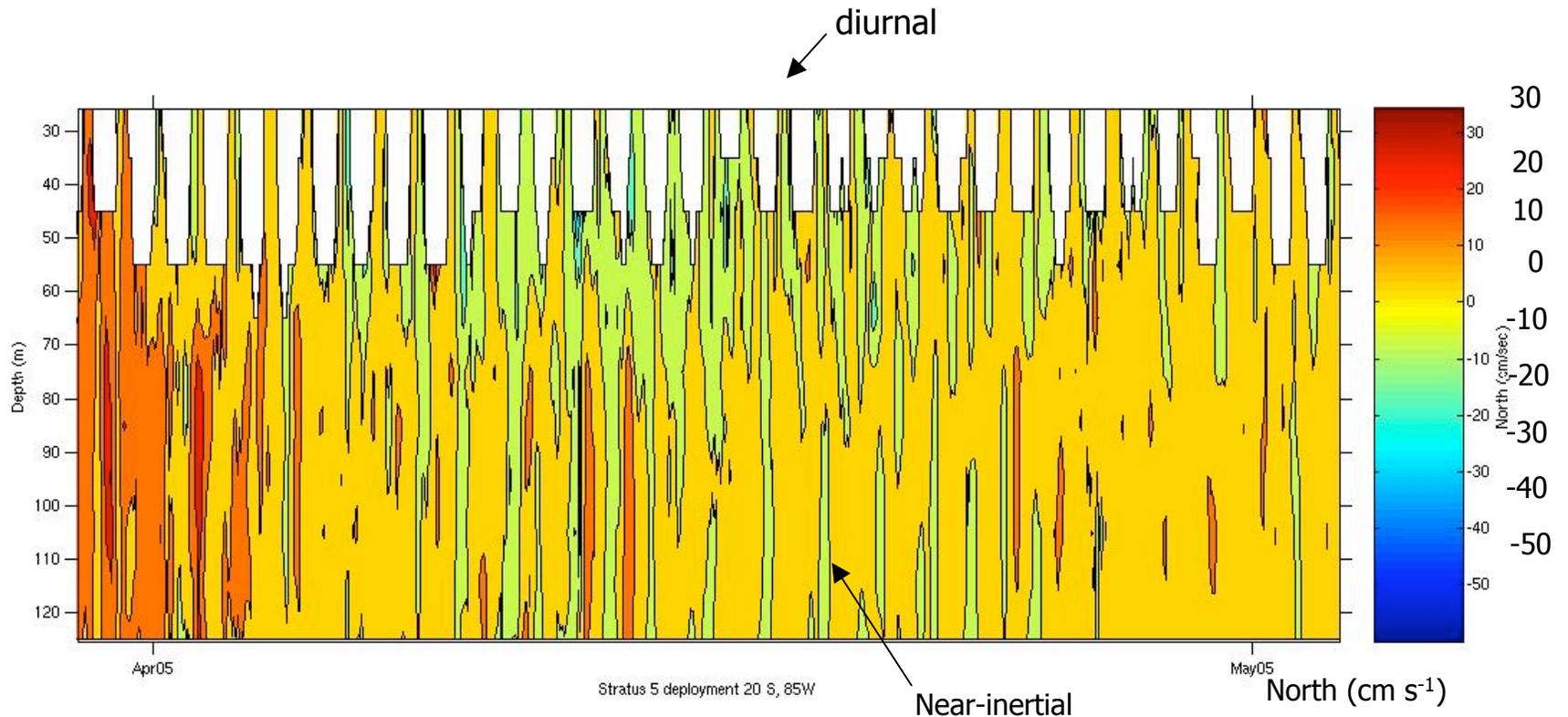


rain



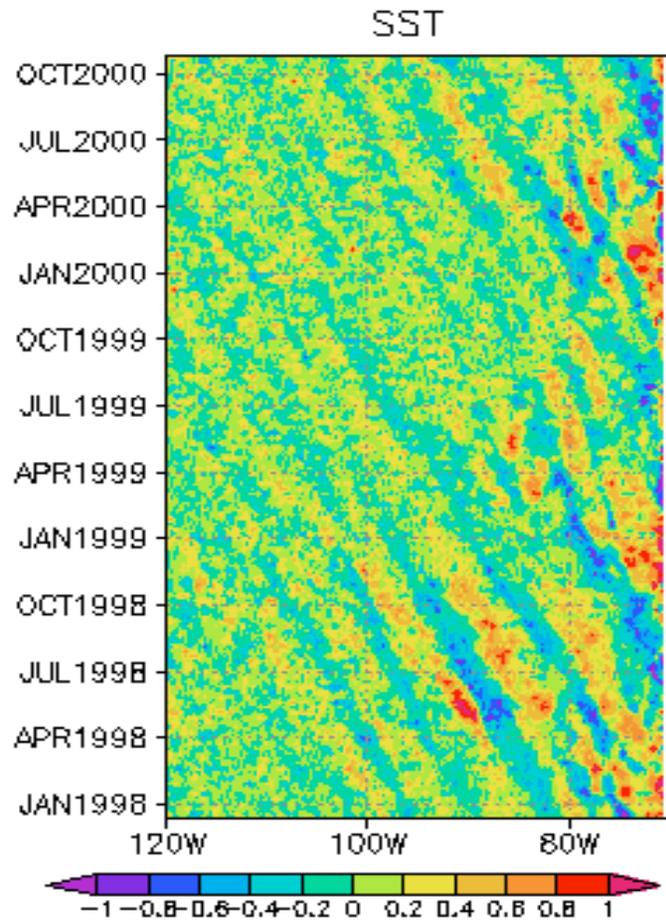
Low wind period - diurnal response

## Of concern: diurnal and near-inertial ocean



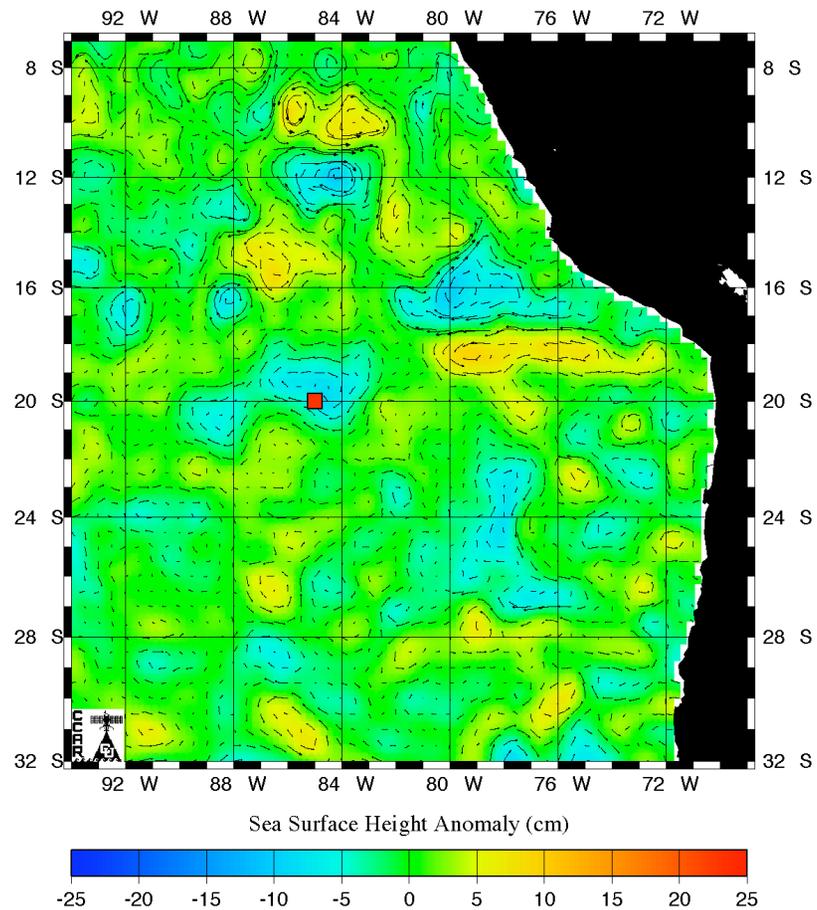
One month of upper ocean velocity (15 to 125 m) data at the Stratus mooring from an acoustic Doppler velocity profiler. Note at the surface, a diurnal (24 hour) cycle in the depth of the scattering layer modulates Doppler return. Below the surface layer, near-inertial (36 hour) oscillations are evident; these play a role in mixing.

Altimetric satellites show westward propagating eddies are typical of the region.

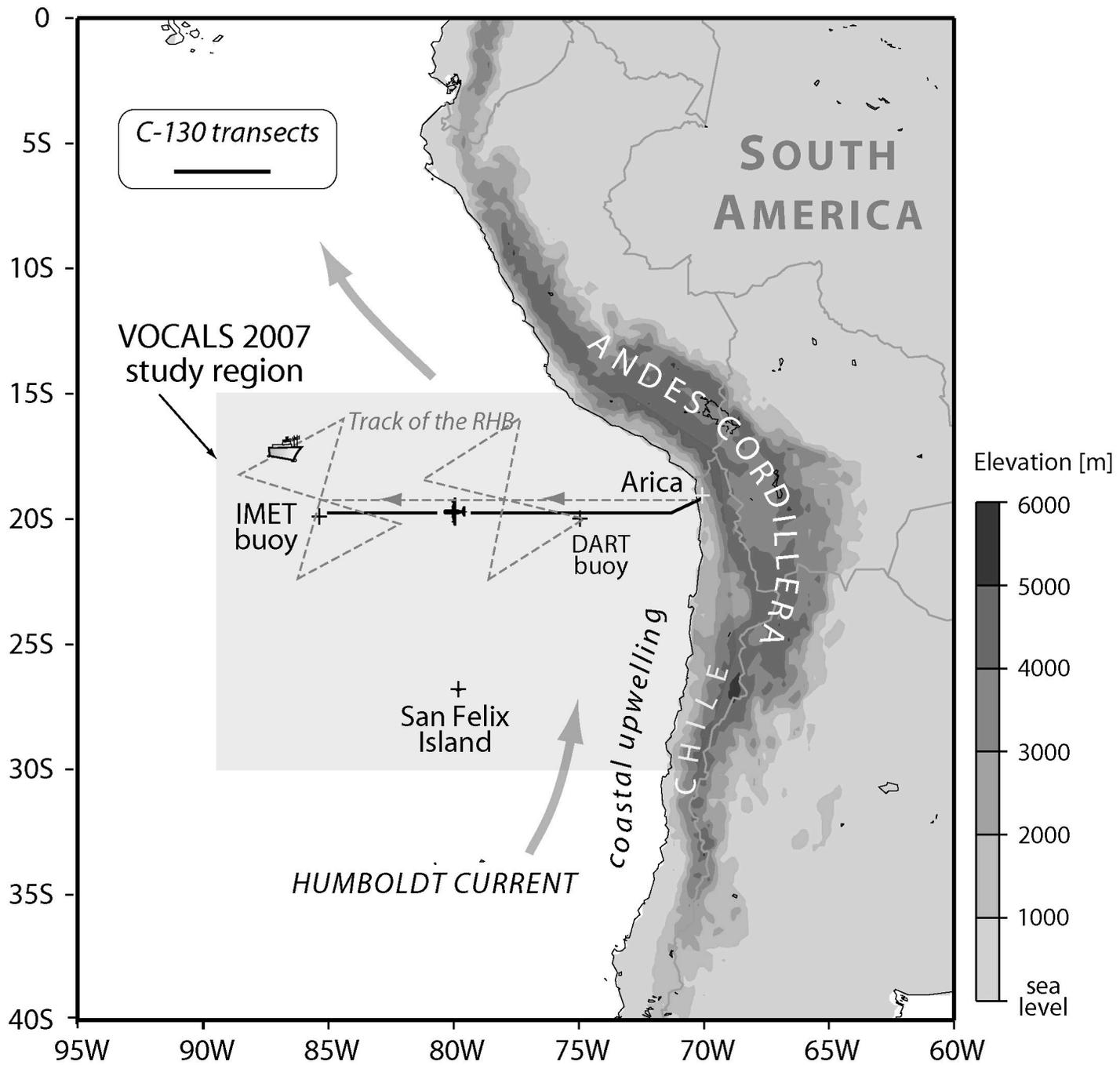


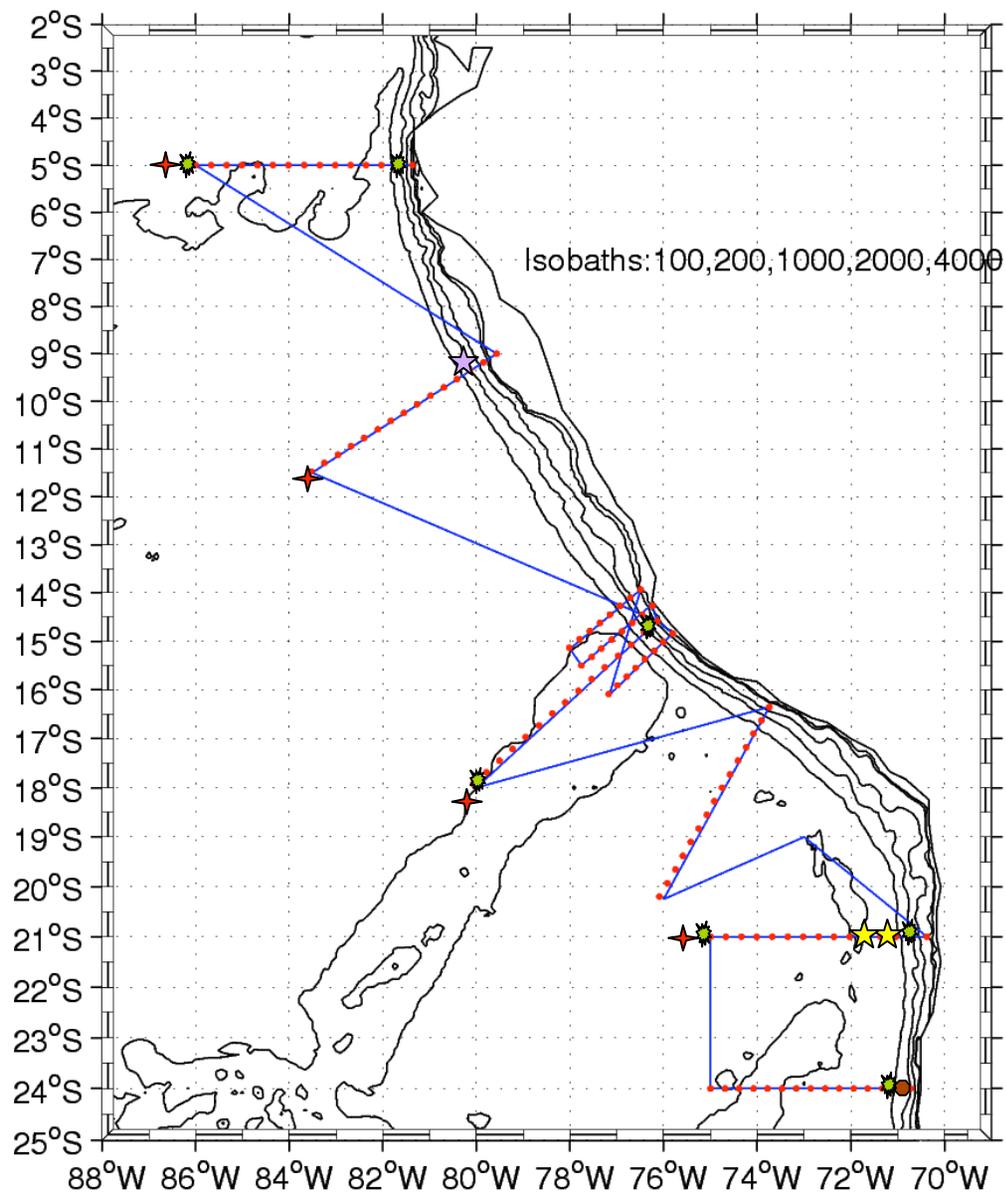
S-P Xie

Historical Mesoscale Altimetry - Mar 17, 1998



Weller





- 6 main sections, 250-300 nm
- 114 stations, 1000m or less
- 4 stations to the bottom 
- 1 « radiator », 100 nm
- Total time: 37 days
- Including:
- 7 x 2-day BGC fixed stations 
- 1 deep-core 
- 
- Existing moorings 
- Future ADCP mooring (pos. to be defined) 

## VOCALS ocean implementation:

Regional scale, annual cycle:

- Argo float, surface drifter
- Lagrangian tracking of eddy by sea gliders
- Satellite
- Model comparisons

High time, high vertical resolution, time series at two contrasting sites with accurate surface fluxes:

- Stratus (20°S, 85°W) NOAA Climate Obs - ongoing
- Chilean Tsunami (DART) buoy (20°S, 75°W) - Oct 2006-Oct 2008 (turn around planned for Oct 2006)

Vertical mixing processes at Stratus and DART sites over diurnal and near-inertial cycles:

- Microstructure profiling for 6-10 days at each
- Near-surface profilers to resolve diurnal layer

## VOCALS ocean implementation:

### Mesoscale sampling - the role of the eddies:

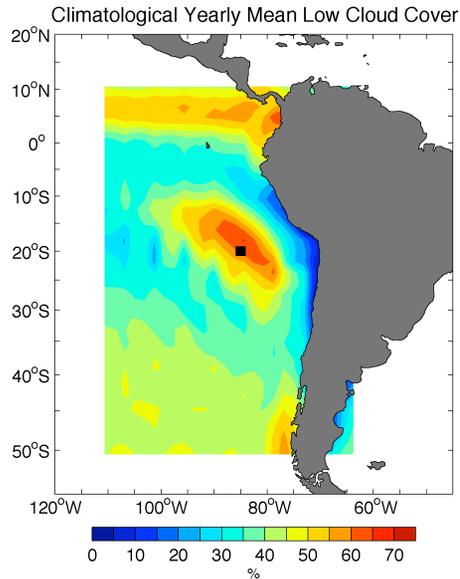
- SeaSoar surveys at Stratus, DART, coastal sites
- Sea glider deployment for one year
- XBT, ADCP sections

### Partnering

- Ongoing NOAA Climate Observation Program
- SOLAS
- PRIMO
- Chilean coastal program

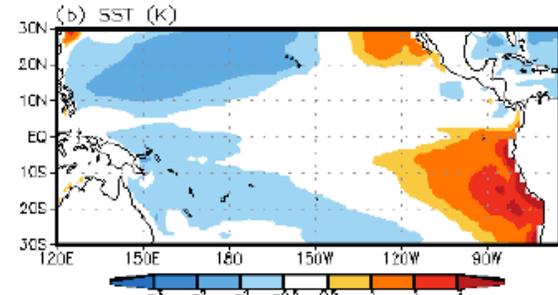
# summary

- Coupled atmosphere-ocean models
  - Tested across regimes, coastal upwelling to open ocean, offshore weak downwelling, eddy-influenced
  - Tested at sites (Stratus buoy, DART) at different phases of diurnal subsidence wave, different aerosol regimes
  - Clear separation of surface layer from upper thermocline
- Surface fluxes at 3 sites (also San Felix)
- Coupled physics-biology ocean models
  - Ability to get surface layer, oxygen minimum layer
  - Ability to get offshore nutrient transport
  - DMS production, impacts on optical absorption



Surface fluxes and ocean structure and evolution: Addressing model (atmospheric, coupled atmosphere-ocean, coupled physics-biology) performance in a region spanning several regimes

SST



CLD

