



# Understanding Atypical Mid-Level Wind Speed Maxima in Hurricane Eyewalls

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Source: NASA/ISS

## Motivation

- Eyewall winds are typically maximized at the top of the boundary layer (~500-1000 m), and decrease monotonically upwards.
- In Patricia (2015), there is an atypical profile, with an additional maximum at mid-levels (4-5 km).
- Here, we show that this structure is likely a consequence of systematic unbalanced flow.

### **Typical Radius-Height Wind Structure**

Hurricane Ivan (2005)



#### Stern et al. (2014)

### Vertical Profiles of Tangential Wind



#### Stern et al. (2014)

### Wind Speed at z=2 km and Dropsonde Trajectories



### Wind Speed from TCI Dropsondes

#### Patricia (10/23/15), ~2000 UTC



### Wind Speed from P3 Doppler Analyses



#### SE to NW Leg (~2030 UTC)



### Azimuthal Mean Vt and Vr from P3 Doppler Analyses



#### This Atypical Wind Structure is Not Unique to Patricia



#### Rita 09/21/2005 height (km)

radius (km)

#### Felix 09/03/2007



#### **Idealized WRF Simulation**



### Modeling an Idealized Patricia



### **Example Axisymmetric CM1 Simulation**



#### CM1 Vt (t=37h), max = 81 m/s



### **Example Axisymmetric CM1 Simulation**

2

0

0 5



radius (km)



10 15 20 25 30 35 40 45 50 55 60

radius (km)

-15

-20

#### Gradient Wind (Vg)

### **Unbalanced Flow**

Kepert (2001)

#### Steady-state and Symmetric Radial Momentum Equation

$$u\frac{\partial u}{\partial r} + w\frac{\partial u}{\partial z} - \left(f + \frac{v}{r}\right)v = -\frac{\partial\phi}{\partial r} + DIFFUSION$$

#### If there is no advection or diffusion:

$$u\frac{\partial u}{\partial r} + w\frac{\partial u}{\partial z} - \left(f + \frac{v}{r}\right)v = -\frac{\partial \phi}{\partial r} + DIFFUSION$$

#### Then we get gradient wind balance

$$\left(f + \frac{v}{r}\right)v = \frac{\partial\phi}{\partial r}$$

### **Unbalanced Flow**

#### If advection and/or diffusion are non-zero, we get unbalanced flow

$$u\frac{\partial u}{\partial r} + w\frac{\partial u}{\partial z} - \left(f + \frac{v}{r}\right)v = -\frac{\partial\phi}{\partial r} + DIFFUSION$$

$$u\frac{\partial u}{\partial r} > 0 \quad \text{AND/OR} \quad w\frac{\partial u}{\partial z} > 0 \quad \text{Contribute to supergradient flow}$$
$$u\frac{\partial u}{\partial r} < 0 \quad \text{AND/OR} \quad w\frac{\partial u}{\partial z} < 0 \quad \text{Contribute to subgradient flow}$$

#### Kepert (2001)

## **Contributions of Radial and Vertical Advection**





### Relationship Between Vertical Structure and Size

#### Initial RMW=90 km

![](_page_15_Figure_2.jpeg)

#### Initial RMW=36 km

![](_page_15_Figure_4.jpeg)

![](_page_15_Figure_5.jpeg)

![](_page_15_Figure_6.jpeg)

![](_page_16_Figure_1.jpeg)

![](_page_17_Figure_1.jpeg)

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![](_page_19_Figure_1.jpeg)

![](_page_20_Figure_1.jpeg)

## Next Steps

Analysis of 3D idealized simulations in CM1

 Boundary layer model of Kepert forced by gradient wind fields from CM1 simulations.

 Kepert's model forced by observational analyses of gradient winds.

## Summary and Conclusions

- Patricia near peak intensity had an atypical wind profile, with both a boundary layer maximum and a mid-level (4-5 km) maximum.
- This structure has been seen in a few other intense/small TCs.
- The mid-level maximum is seen in simulations, and is a consequence of systematic unbalanced flow associated with a vertical oscillation in the radial winds.

## **Bonus Slides!**

![](_page_25_Figure_0.jpeg)

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