# Simulating Hurricane Joaquin (2015) with WRF-ARW: Challenges and progress



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# Objectives

- Perform a successful simulation of Hurricane Joaquin (2015) using the WRF-ARW model;
- Study the environment-vortex interactions that caused the storm to follow a looping track; and
- Examine the physical mechanisms responsible for two episodes of rapid intensification, including the relationship between the upper outflow and boundary-layer inflow through trajectory analysis.

# Overview

- "Looping" track:
- Two distinct rapid intensification (RI) episodes:
  A) 29 Sep 0600 UTC 01 Oct 1200 UTC (41 m s<sup>-1</sup> 2.25 day<sup>-1</sup>)
  B) Re-intensification; 03 Oct 0600 UTC 03 Oct 1200 UTC (7.7 m s<sup>-1</sup> 6 h<sup>-1</sup>)
- Both track and intensity dependent on UL ridge and trough configuration
- Operational models (except for ECMWF) performed poorly in storm track prediction



# Major Challenges For Track Prediction

- 1) sensitivity of track to initial vortex structure
- 2) Hypothesis: a stronger, deeper vortex should be more influenced by mid-to-upper level steering flows.



2) accurately modeling upper-level ridge-trough configuration

- Strategies:
- Initialize WRF using European Center Ensemble (ECENS) Reanalysis data
- Generate bogus vortex (Kwon and Cheong *MWR* 2009) relocated SE of best-track position
- Use WRF Data Assimilation System (WRFDA, Huang et al. *MWR* 2009) to generate an improved analysis at initialization time

# 24-h WRF forecasts initialized every 6 h, from 28 Sep 00 UTC – 29 Sep 12 UTC

30

29

28





#### WRF-ARW 1-km Control Simulation Results

WRF-predicted (green squares) and best-track 6-hourly storm positions superimposed over the SST change from initial time to 1200 UTC 02 Oct



#### Control Structures at Peak Intensity Time

- Model-simulated fields (a,c) at 0600 UTC 03 Oct

- SSMI/S 91-GHz color composite (b) with satellite AMVs (d) at 1200 UTC 03 Oct



#### **Steering Flow Analysis**

- Winds averaged horizontally over a 6 deg × 6 deg box surrounding storm center
- (a,b) : winds as a function of height
- (c,d) : deep-layer 200 900 hPa mass-weighted averages
- (e,f) : storm motion vectors



## Using 4DVAR Data Assimilation to Improve Model Initialization

- Like 3DVAR data assimilation, minimizes a cost function to find the analysis solution where both model errors and observational errors are smallest
- Unlike 3DVAR, data can be assimilated over a much larger time window → more observations, model errors are "flow dependent"
- Uses a linear tangent/adjoint WRF model to advance the analysis through the time window to compare against observations → also provides a dynamical constraint in cost function minimization



# WRFDA strategy for Joaquin (2015)

- Generate background error matrix using a 1-month period of WRF forecasts, using the difference between the 24-h and 12-h forecasts verifying at the same time as a proxy for model error (NMC Method, Parrish and Derber *MWR* 1992)
- Experiment with both "cold start" and cycling WRFDA-initialized runs
- Assimilate as many observations as possible, including satellite AMV winds, station soundings, aircraft reports (AIREPS), recon plane dropsondes, satellite radiances
- For now, only perform DA on the outermost 27-km domain



# 5-day WRF simulation Initialized from WRFDA Analysis

- Using daily-averaged AVHRR SSTs as boundary conditions
- Assimilate AMVs and conventional observations over 6-hr window
- Re-intensification is now simulated





WRF-predicted (green squares) and best-track 6-hourly storm positions superimposed over the SST change from initial time to 1200 UTC 02 Oct

# Was the analysis at 28 Sep 1200 UTC Improved by Data Assimilation?

- WRFDA analysis NCEP reanalysis difference fields for ~500 hPa level
- Stronger Bahamas ridge
- Mid-upper level large scale flows become more northerly over Joaquin



#### Now Zooming in on Mid-upper Level Vortex

78W

80W

76W

74W

7ÓW

72W

68W

10

66W

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- WRF-generated vortex at 28 Sep 1200Z is weaker than the NCEP reanalysis vortex
- WRFDA has small impact on 500-hPa vortex structure



28 Sep 1200Z; WRF-generated background 500 hPa heights



## How Does WRFDA Impact Vortex at Later Initialization Time?

- Continue cycling process through 29 Sep 1800 UTC •
- Here we assimilate a wide variety of observations: .
  - AMVs, AIREPs, upper-air soundings
  - satellite radiances: AMSU-A, MHS, AIRS
  - 11 NOAA G-IV dropsondes (released at 2000 & 2100 UTC 29 Sep)
- WRFDA strengthens 500 hPa ridging NW of storm, but weakens vortex!





25

5880

5850

5840

5810

5800

5790

## Calculating Trajectories from WRF Model Output



- Trajectory analysis to address TCI objectives:
- Investigate relationship between high surface fluxes and outflow layer development through inner-core deep convection
- Study thermodynamics of convective bursts
- Compare the Wilma (2005) and Joaquin (2015) cases
- Calculating trajectory algorithms:
- Kinematic trajectories computed from gridded 5-min resolution model-output winds
- 2<sup>nd</sup> order Runge-Kutta forward and backward time integrations with a 10-s time step
- Advection correction (Shapiro et al. JAS 2015) used in time interpolations

# Summary

- CTL simulation (no DA) Successes:
  - Reproduces loop-shaped track
  - Captures timing and intensity of first RI period reasonably well
  - Inner-core convection and outflow layer structures compare well against satellite observations at time when the storm is most intense
- CTL simulation (no DA) room for improvement:
  - Does not show the observed slow counterclockwise turn to NE
  - Track errors increase substantially after 0000 UTC 03 Oct
  - No re-intensification period
- Preliminary results with assimilating AMVs:
  - Synoptic-scale analysis increments favorable for southward early track
  - Weakening of upper-level vortex structure a problem that still needs to be resolved
- Trajectory Analysis of Outflow Layer surface flux connection:
  - New methodology using advection correction currently being tested
  - Early results show that convective bursts originate from regions of PBL where surface heat fluxes are higher
  - Early results show that profiles of thermodynamic and microphysical variables through CBs traced to PBL are consistent with those shown in other modeling studies (Fierro et al. 2009, 2012)

# Plans For Future Work

• Continue trying to improve upon ECENS analysis with WRFDA:

- Perform single-observation tests to study impact of domain-tailored static background error, and fine-tune as necessary

- Assimilate radiance data

- Consider removing inner-core AMVs or accounting for possible height/speed biases

- Try multiple "outer loops" in 4DVAR minimization  $\rightarrow$  standard practice now for operational NWP centers using 4DVAR

• Dealing with the SST cooling impact on Joaquin's intensity forecast

- SST cooling parameterization being developed for WRF-ARW:

Liu, X., J. Wei, D.-L. Zhang, W. Miller, 2016 MWR accepted with major rev. Typhoon Matsa (2005) WRF simulation

- SST tendency computed at every model time step
- scheme generates ocean currents using the WRF-predicted 10-m horizontal winds
- vertical mixing and advection drive SST cooling, with mixed-layer depth accounted for
- SST recovery following TC passage also computed



• Perform statistical analysis on trajectory data on CBs versus background secondary circulation: thermodynamic characteristics, outflow layer composition

#### Advection Correction: accounting for advection of flow disturbances when performing time interpolations of gridded data

One-dimensional translating wave schematic

comparing linear time interpolation (LI) with

advection correction (AC) 40 U∆t w (m/s) 10 t<sub>o</sub> 30 20 368 0 ≻х 366 р 10 w (m/s) 364  $t_0 + \Delta t/2$ 10-AC 362 0 360 -10 358 0 →X 356 р w (m/s) -20 $t_0 + \Delta t$ 10--30 -40 -20 -10 10 30 -30 20 -40 Ò 40 0 ≻ X p't  $p'_{t_0} + \Delta t$ p km 50

z= 14 km vertical velocity (contours) with theta\_e (shaded) and storm-relative flow vectors

E K