



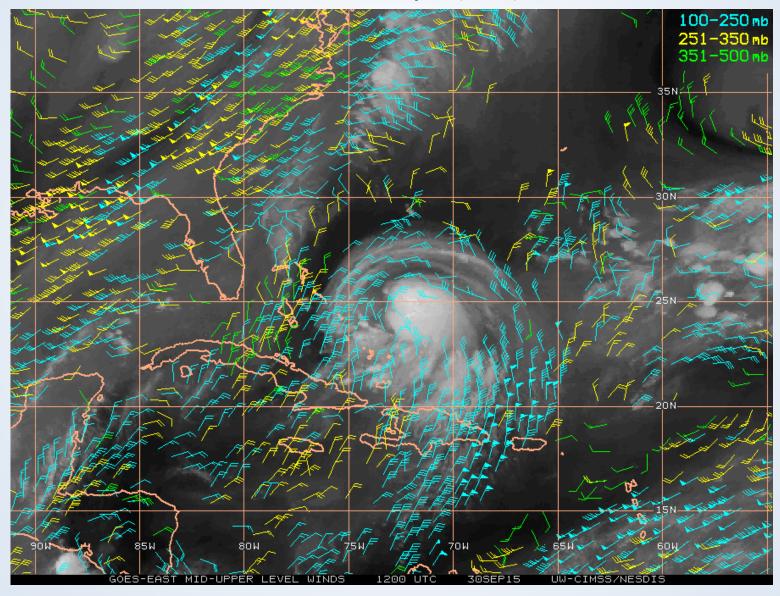
Simulating TC-trough interaction in an idealized modeling framework

Will Komaromi Jim Doyle Naval Research Lab - Monterey will.komaromi@nrlmry.navy.mil

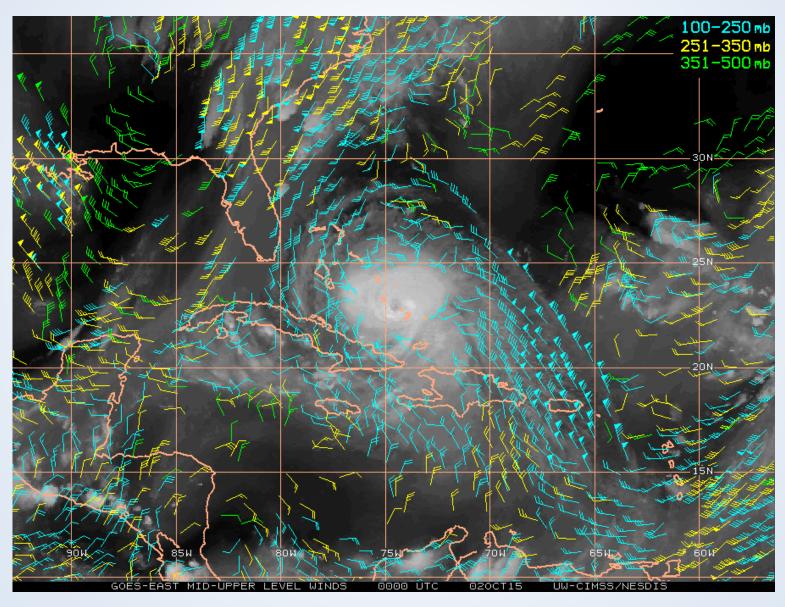
> *TCI Science Meeting 18 October 2016*

Stage 1: pre-interaction, outflow entirely to S

Hurricane Joaquin (2015)

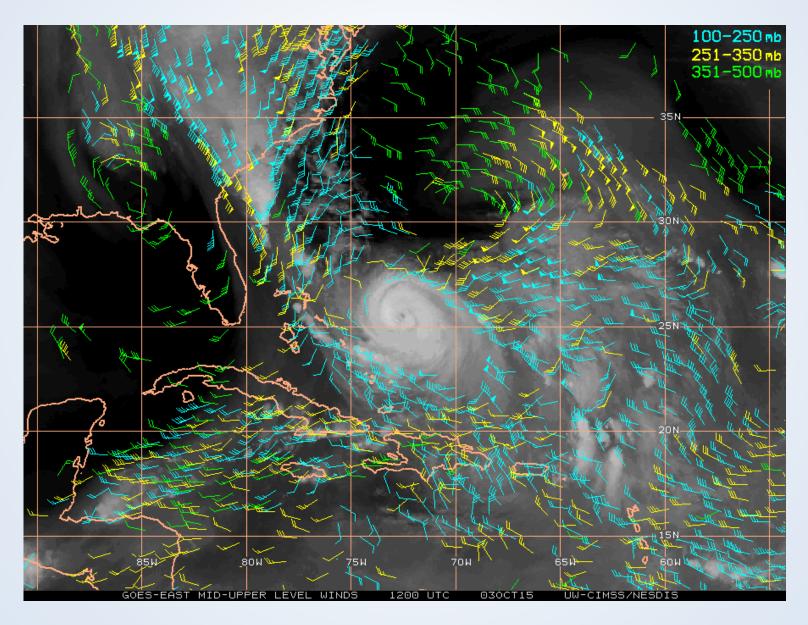


Stage 2: dual outflow channels develop, slight weakening as shear increases



Stage 3: Strongest outflow to NW & NE (but still have decent outflow to S)

peak intensity occurs at this time



How can trough interaction theoretically be favorable for TC intensification?

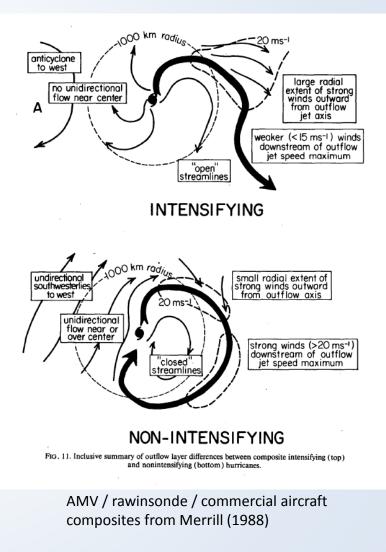
- Enhanced upper-level divergence
- Enhanced ascent associated with trough's secondary circulation and/or right-entrance region of jet streak
- TCs primarily interact with environmental flow at upper levels where inertial stability (I) is weaker, as I resists motion in radial direction. Reduction in I may allow stronger radial outflow to develop

$$I = \sqrt{\left(f + \frac{2v_t}{r}\right)(f + \zeta)} \text{ where } \zeta = \frac{1}{r} \frac{\partial(rv_t)}{\partial r}$$

 Flux convergence of angular momentum by azimuthal eddies (EFC) may trigger structural changes which ultimately lead to intensification of TC. For many real cases window of positive impacts associated with EFC may be short: EFC may be too far from TC, or shear may be too great once EFC is close enough to have effect on TC

$$\text{EFC} = -\frac{1}{r}\frac{\partial}{\partial r} \left(r^2 \overline{v_r' v_t'} \right)$$

Of course, trough interaction can also be unfavorable to a TC, primarily due to increased vertical wind shear



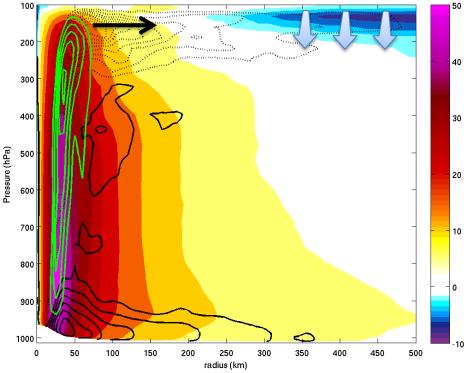
Model Configuration

- COAMPS-TC v4
- 5km res, 801x801 grid points (4005x4005 km)
- 40 vertical levels
- No cumulus parameterization
- Periodic in x, wall boundaries in y
- Modified Mellor-Yamada PBL scheme
- Lin et al. (1983) 5-class single moment microphysics
- Radiation off
- β-plane
- Fixed SST 28.0 °C
- θ (K) and q (g/kg) from Dunion (2011) MT sounding
- Initialized with rankine vortex
 - − rmw = 90 km, $V_t \rightarrow 0$ = 240 km

How can we test TC sensitivity to "environmental" forcing in outflow region?

- Option A: enhance V_r (similar to nudging, but with balanced pressure gradient) → tried without much success, won't be presenting today
- Option B: Reduce V_t at slightly larger r via zonal jet or trough in order to reduce I, which should result in environment more favorable for enhanced V_r. Trough interaction should also promote EFC, which may indirectly intensify the TC





Azimuthal avg V_t (ms⁻¹, shaded), V_r (2 ms⁻¹ increments solid<0, dashed>0), w (0.5 ms⁻¹ increments)

We define jet in terms of u_g. From u and v momentum equations, incorporating a large-scale flow:

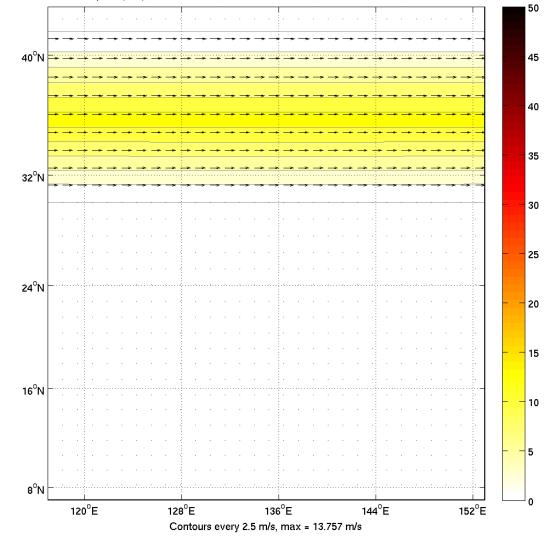
$$\frac{\partial u}{\partial t} \sim f(v - v_g)$$

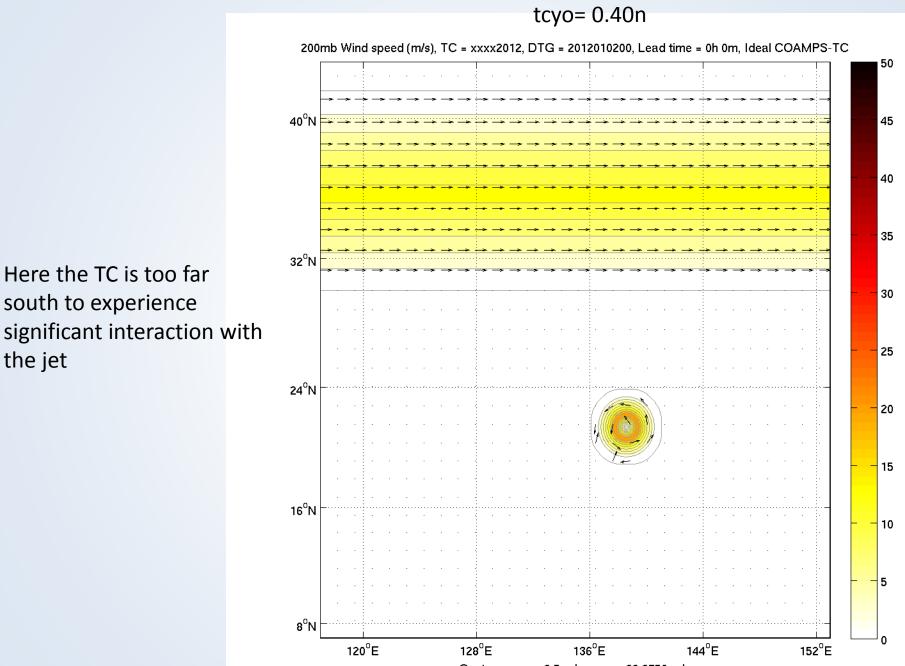
$$\frac{\partial v}{\partial t} \sim -f(u-u_g)$$

where
$$\begin{bmatrix} u_g \\ v_g \end{bmatrix}$$
 is balanced by a background $\begin{bmatrix} -\frac{\partial \prod}{\partial y} \\ \frac{\partial \prod}{\partial x} \end{bmatrix}$ and $\prod = \left(\frac{p}{p0}\right)^{\frac{R_d}{c_p}}$

Jet extends from ~500-100 hPa, strongest at 300 hPa decaying linearly above and below

200mb Wind speed (m/s), TC = xxxx2012, DTG = 2012010200, Lead time = 0h 0m, Ideal COAMPS-TC





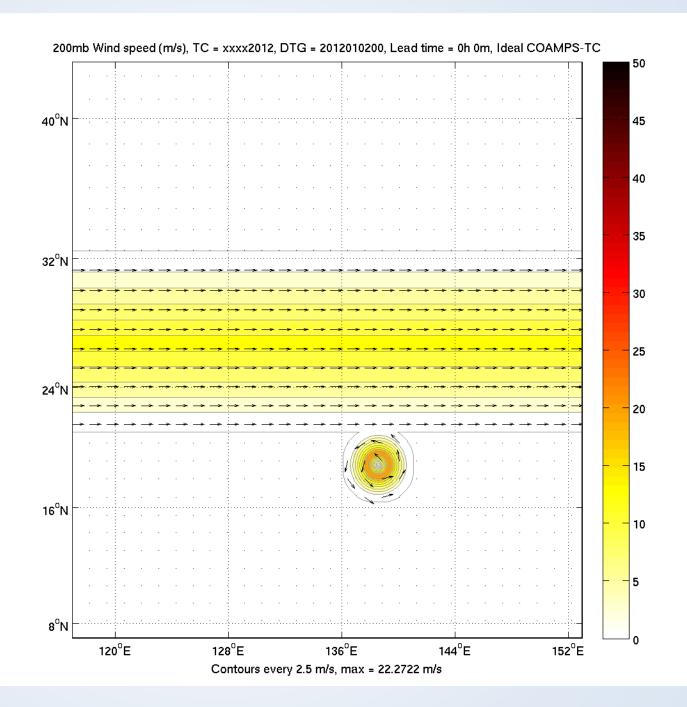
south to experience significant interaction with the jet

Contours every 2.5 m/s, max = 22.2758 m/s

TC + zonal jet

Here we move the jet S towards the TC (as opposed to moving the TC N towards the jet), otherwise changes in I will be dominated by changes in *f*

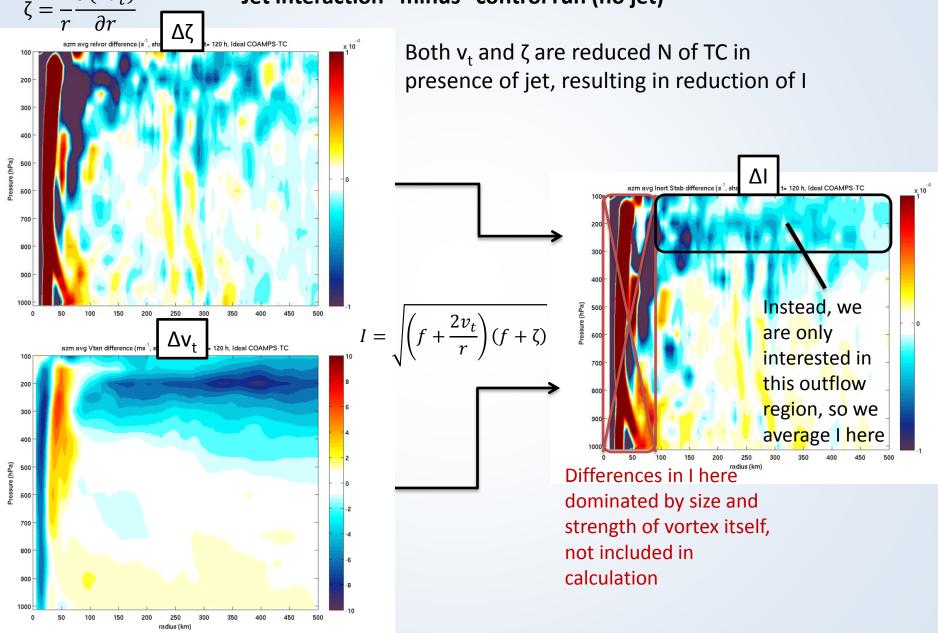
Here the TC interacts more strongly with the jet, although the TC appears to be sheared by the end of the simulation

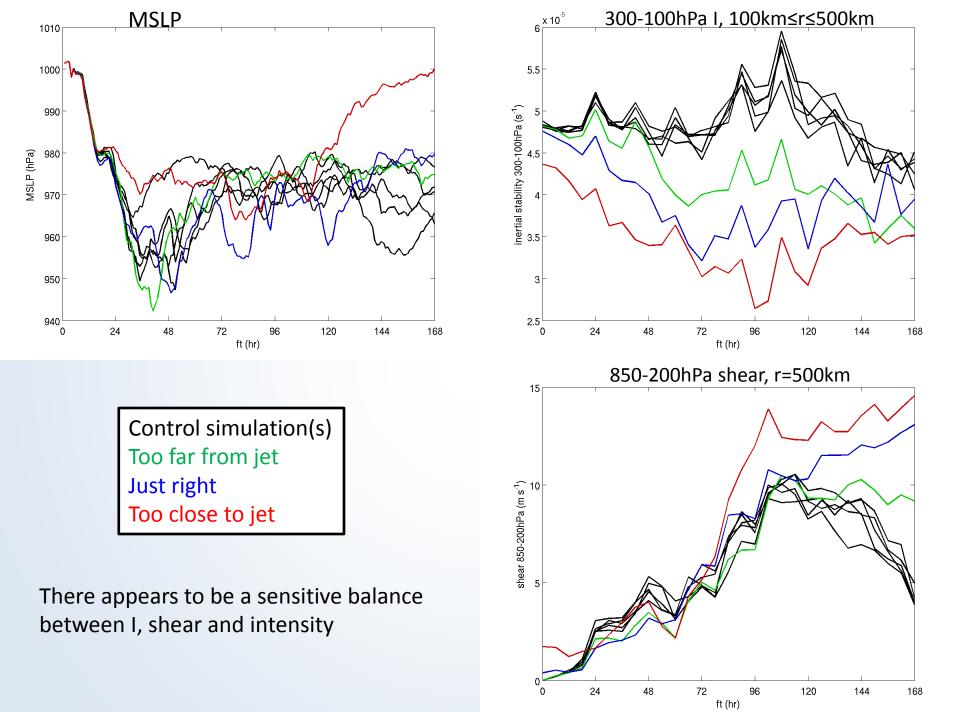


Quantifying change in inertial stability: azimuthal means in radius-pressure



 $1 \partial (rv_t)$



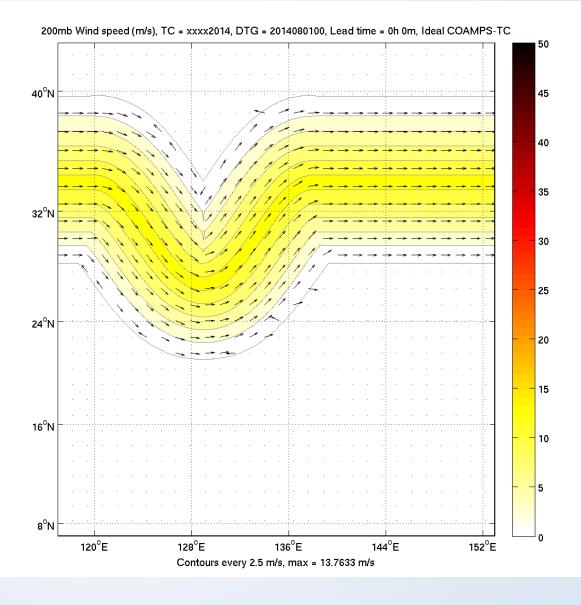


Jet extends from ~500-100 hPa, strongest at 300 hPa decaying linearly above and below

We define a trough using a cos wave:

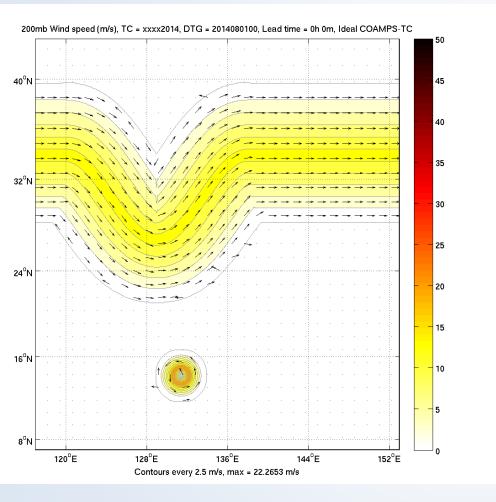
$$y = A\cos\left(\frac{4\pi x}{X} + \frac{5\pi}{3}\right) + y_0 - A$$

A = trough amplitude (here 400km) x = location in x X = size of domain in x y₀ = displacement in y

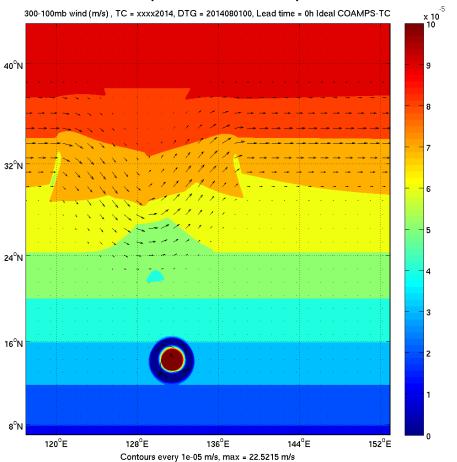


Minimal interaction: TC is too far south

200hPa wind

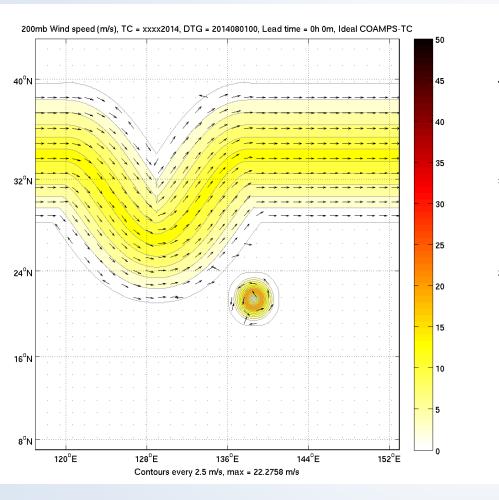


300-100hPa avg inertial stability (vortex relative)

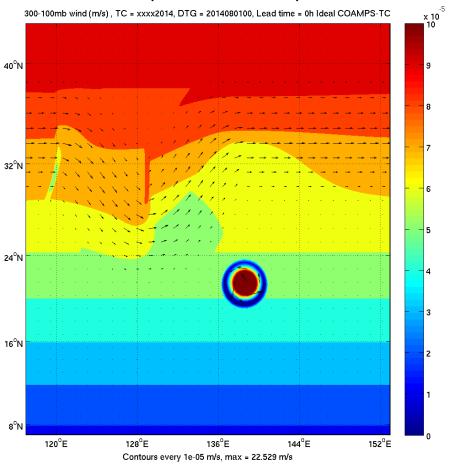


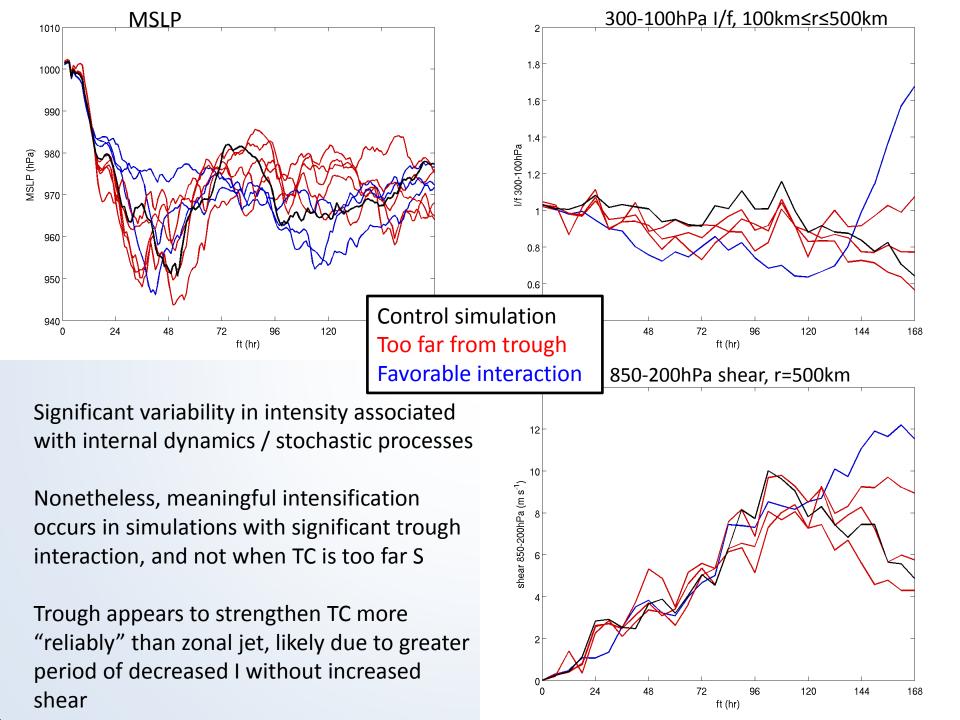
Favorable TC-trough interaction

200hPa wind

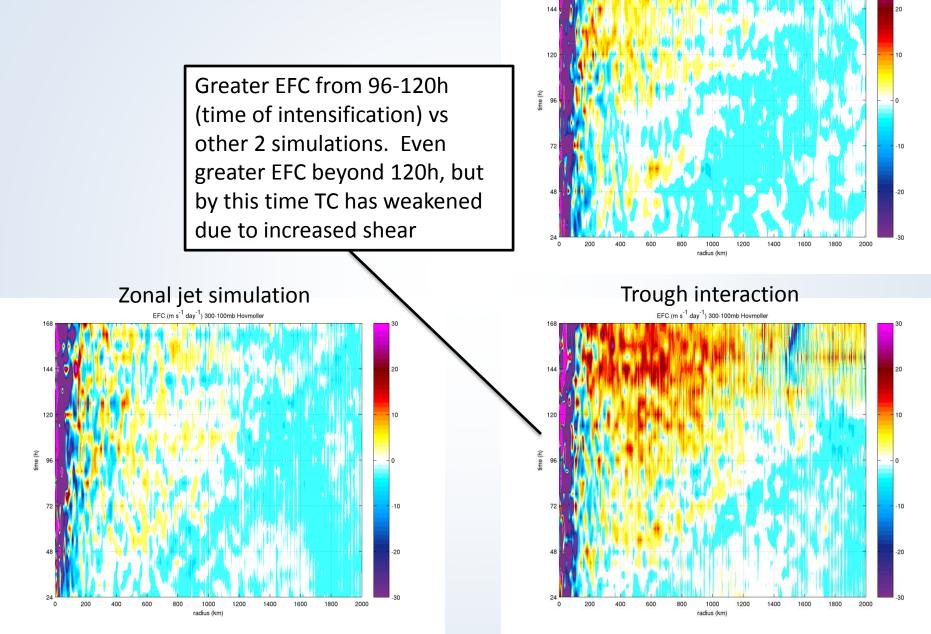


300-100hPa avg inertial stability (vortex relative)





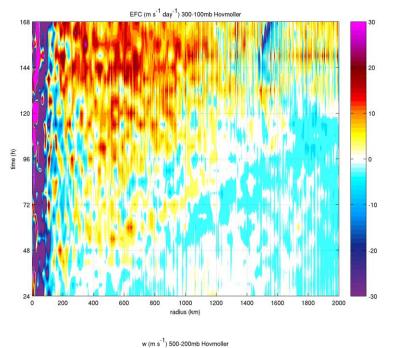
Hovmollers of EFC 300-100mb mean, r = $0 \rightarrow 2000$ km

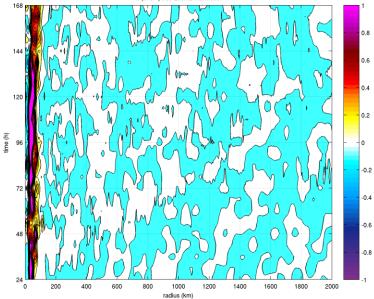


No jet simulation (control)

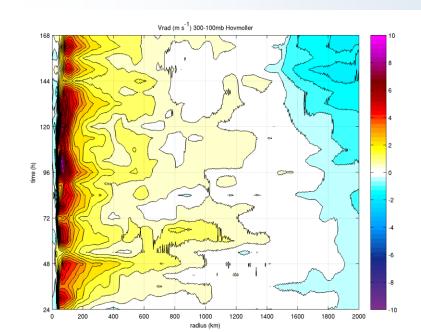
EFC (m s⁻¹ day⁻¹) 300-100mb Hovmoller

Does large +EFC trigger inward-propagating maximum (either local or absolute) in v_r or w?





At least in these simulations, it doesn't appear so



Summary

- We have examined the interaction between a trough/jet and TC:
- If the TC is too far from the trough, minimal interaction occurs and the result is similar to a control run (no trough)
- If the TC is close enough to the trough, favorable interaction occurs from 96-120h during the period of maximum reduction in inertial stability
- Beyond 120h, the TC drifts N (due to beta effect and southerly flow ahead of trough) and weakens under stronger shear
- Trough appears to strengthen TC more "reliably" than zonal jet, likely due to greater period of decreased I without increased shear
- If EFC is indeed contributing to TC intensification, mechanism is still somewhat unclear
 - Neither an inward-propagating maximum in v_r nor w is observed
- Currently running additional sensitivity tests, testing for robustness of results