



NRC research update

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HAIC-HIWC Science Team Meeting, 16-19 May 2016



National Research Conseil national Council Canada de recherches Canada



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W/X-band Dual – Frequency Analysis



NAWX - IWC vs. Z_e



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NAWX - Z (23-May-2015)



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Polarimetric method for HIWC identification and estimation



Background



 $Z\downarrow h, v = \lambda \uparrow 4 /\pi \uparrow 5 K\uparrow 2 \int \uparrow m \sigma \downarrow h, v N(D)$

$K \downarrow dp = 180 \lambda / \pi \int f m Re(f \downarrow h - f \downarrow v) N(h)$

$$\sigma \downarrow h - \sigma \downarrow v / Re(f \downarrow h - f \downarrow v) = 2\pi \uparrow 2$$

Conventional estimation of specific differential phase K_{dp}

 $\Psi \downarrow dp = arg\{\langle H \downarrow i \downarrow \uparrow * V \downarrow i \rangle\}$ $\Psi \downarrow dp = \Phi \downarrow dp + \delta \downarrow dp$ $K \downarrow dp (r) = 1/2 d$ $\Phi \downarrow dp (r)/dr$



The NAWX system





W band: pale yellow

	X-band	W-band
RF output frequency	9.41 GHz +/- 30 MHz	94.05 GHz
Peak transmit power	25 kW magnetron split between two ports	1.7 kW typical
Transmit polarization	H and V	H or V
Maximum Pulse Repetition Rate	5 kHz	15 kHz
Receivers	2	2
Receiver polarization	simultaneous H and V	co and cross-polarization
Processing	Pulse pair and FFT	Pulse pair, FFT, PDPP and raw (IQ)
Antennas	Side: 1 x 26" dual-pol Nadir/Zenith: 2 x 18" single-pol	Side/Aft: 2 x 12" dual-pol Nadir/Zenith: 1 x 12" single-pol



Method for estimation of K_{dp}





Α





Case Study: May 23



Case Study: May 23





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Case Study: May 23

IWC retrieval using Kdp(X)



Map Z(X) and est.



Future work

- IWC Kdp, Zdr, T relationship (X-band)
- Above + polarimetric W-band
- Retrieval algorithm using polarimetric X-band, W-band and lidar/radiometer







Putting it all together...



23-May; Increasing extinction

20-May; Decreasing extinction





Fine-scale structure

May 15, 2015



Future tasks

- Klett inversions in the nadir direction
- Klett sensitivity to the initial conditions: reference range, reference extinction/backscatter value, lidar ratio.

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- Multiple-scattering correction
- Depolarization ratio (absolute values)
- Time synchronization
- Zenith and nadir return power coherency
- More test case analysis

Simulations of 183 Ghz GVR Microwave Brightness Temperatures for Ice Clouds Over Ocean Water



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G-band (183 GHz) water Vapor Radiometer (GVR)

 Developed by ProSensing Inc. and first airborne installation on Convair in 2007 (Pazmany & Wolde, 2009)

Measures brightness temperature at 183.31 ±1, ±3, ±7 and ± 14 GHz

Neural Network Retrievals of PWV and LWP from GVR brightness temperature (Pazmany, 2009; Cadeddu et. al., 2009)





Example of GVR responses to the HIWC environemnt:17-09:55:41 UTC



Objective

Scalar radiative transfer simulations of brightness temperatures were compared to those mmeasured by the GVR radiometer to ascertain:

- the sensitivity of the latter to thermodynamic/optical properties of the cloud and atmosphere as well as the surface albedo
- to determine the consistency of the measured radiometric data.



Microwave scattering properties of water and ice particles



Optical depth as a function of frequency for the US standard atmosphere with no moisture and with 100% relative humidity. At the GVR frequencies, the relative humidity introduces significant absorption. Thus the accurate computation of the latter requires knowledge of the vertical distribution of water vapor.



Variation of the real and imaginary components of the index of refraction of ice with frequency and temperature. The index of refraction is used to calculate the Mie scattering properties of the spherical ice particles.

Sensitivity Analysis



Model cloud, comprised of ice particles, with upper and lower boundaries deduced from radar range gates at time 10:49:22.

Sensitivity of brightness temperatures at the 8 GVR frequencies to biases in pressure, temperature, relative humidity, particle number concentration and surface albedo.



Effect of Column Optical Depth Averaging on Brightness Temperature Calculations

The purpose of this analysis was to compare radiances calculated without (vertical) spatial averaging of the optical properties to cases where such averaging was employed. The four cases analyzed in detail were selected manually at the times indicated by the red vertical lines in the figure. At these times, the cloud exhibits plane parallel characteristics and contains mainly ice particles. The layers altitude, pressure, temperature and relative humidity) are shown below. Results are illustrated In the next slide.



Cloud consisting of 438 layers.



Selected vertical cloud distributions exhibiting plane parallel characteristics and containing mainly ice particles

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Simulated Brightness Temperatures Along the Flight Path



Averaged cloud, consisting of 29 layers.



As above (right panel) but with the cloud optical depth multiplied everywhere by a factor of five.



Brightness temperature measurements compared to simulations. Time interval is 6.0 seconds, or 239 measurements.

Results of radiative transfer simulations are In broad qualitative agreement with observations, considering the absence of the thermodynamic structure of the atmosphere and the PSD within the cloud.

