

# Improving the Prediction of Cloud Ice in Operational NWP Models

**Jason Milbrandt**

Environment and Climate Change Canada (RPN-A)

In collaboration with:

**Hugh Morrison**, NCAR (Boulder, USA)

**Zhipeng Qu**, ECCC (Toronto)

# Simplified organizational chart of ECCC

## Environment and Climate Change Canada

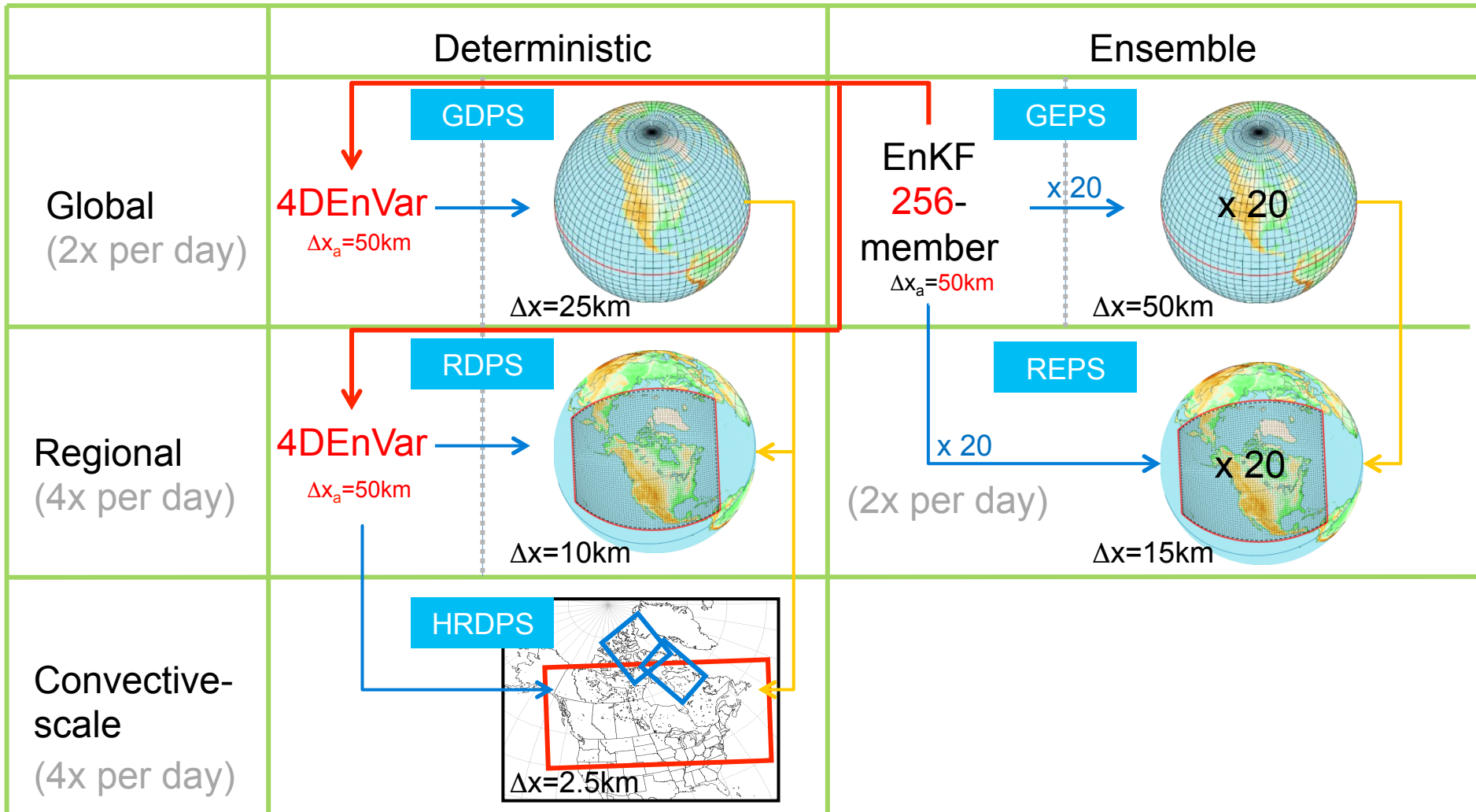
### Science and Technology Branch

**Cloud Physics and Severe Weather** (Toronto)  
- measurements, nowcasting techniques, etc....

**Numerical Weather Prediction Research** (Montreal)  
- development of NWP model and modeling systems

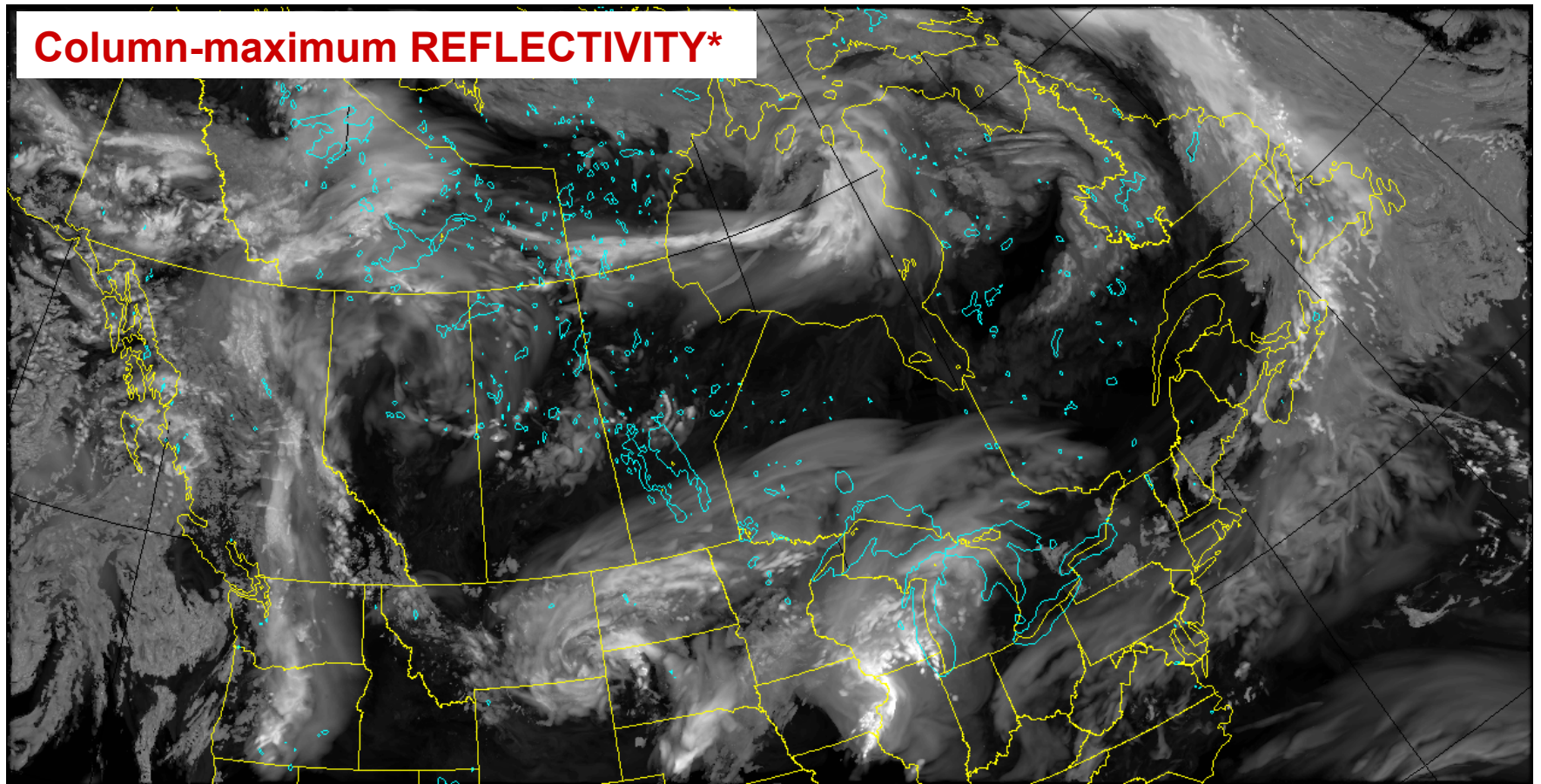
**Meteorological Services of Canada**  
- operational NWP (modeling, forecasting, ...)

# ECCE's current NWP systems



# Simulation with 2.5-km HRDPS

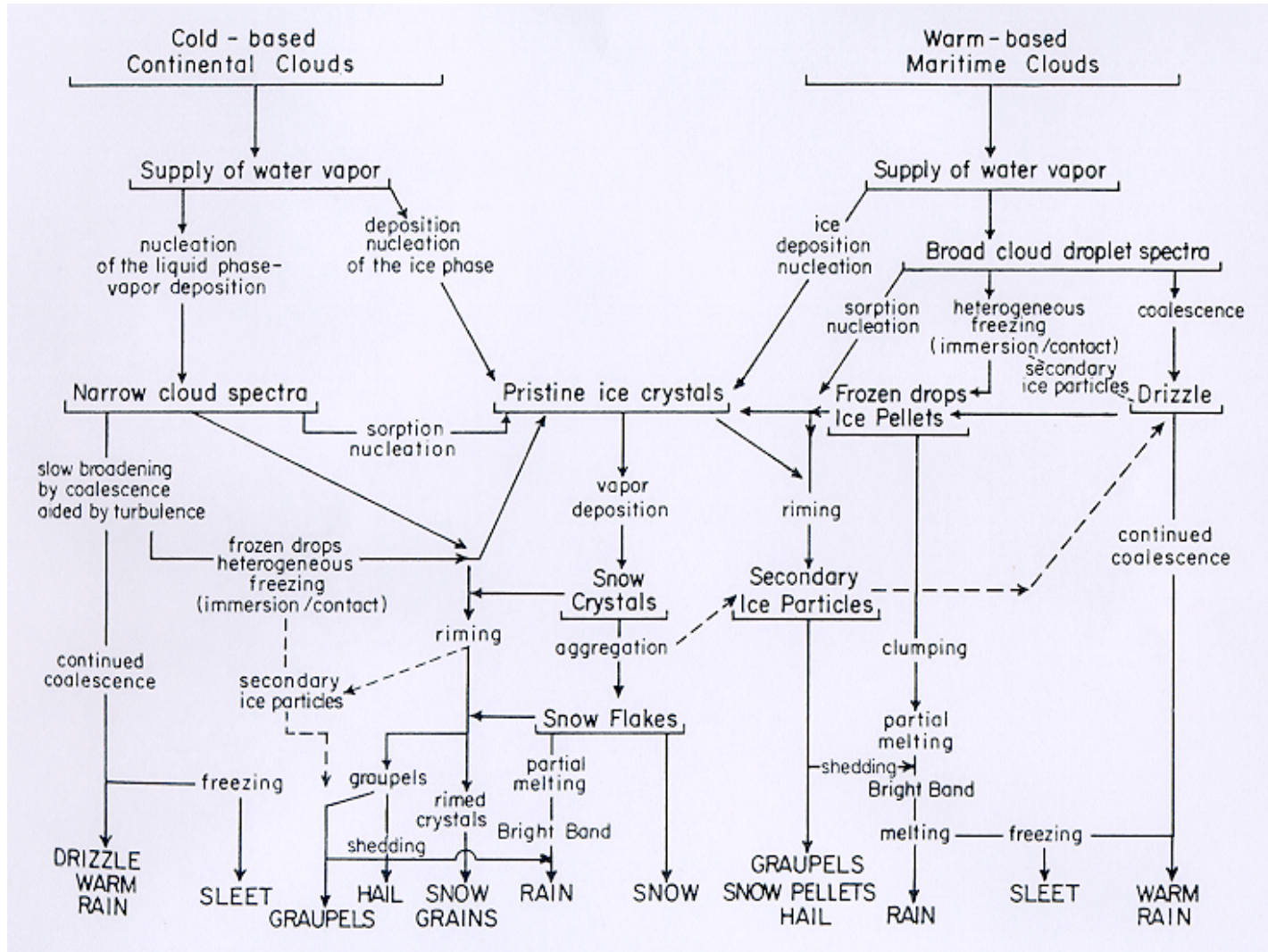
(High Resolution Deterministic Prediction System)



\* Computed from microphysics  
(Milbrandt-Yau 2-moment)

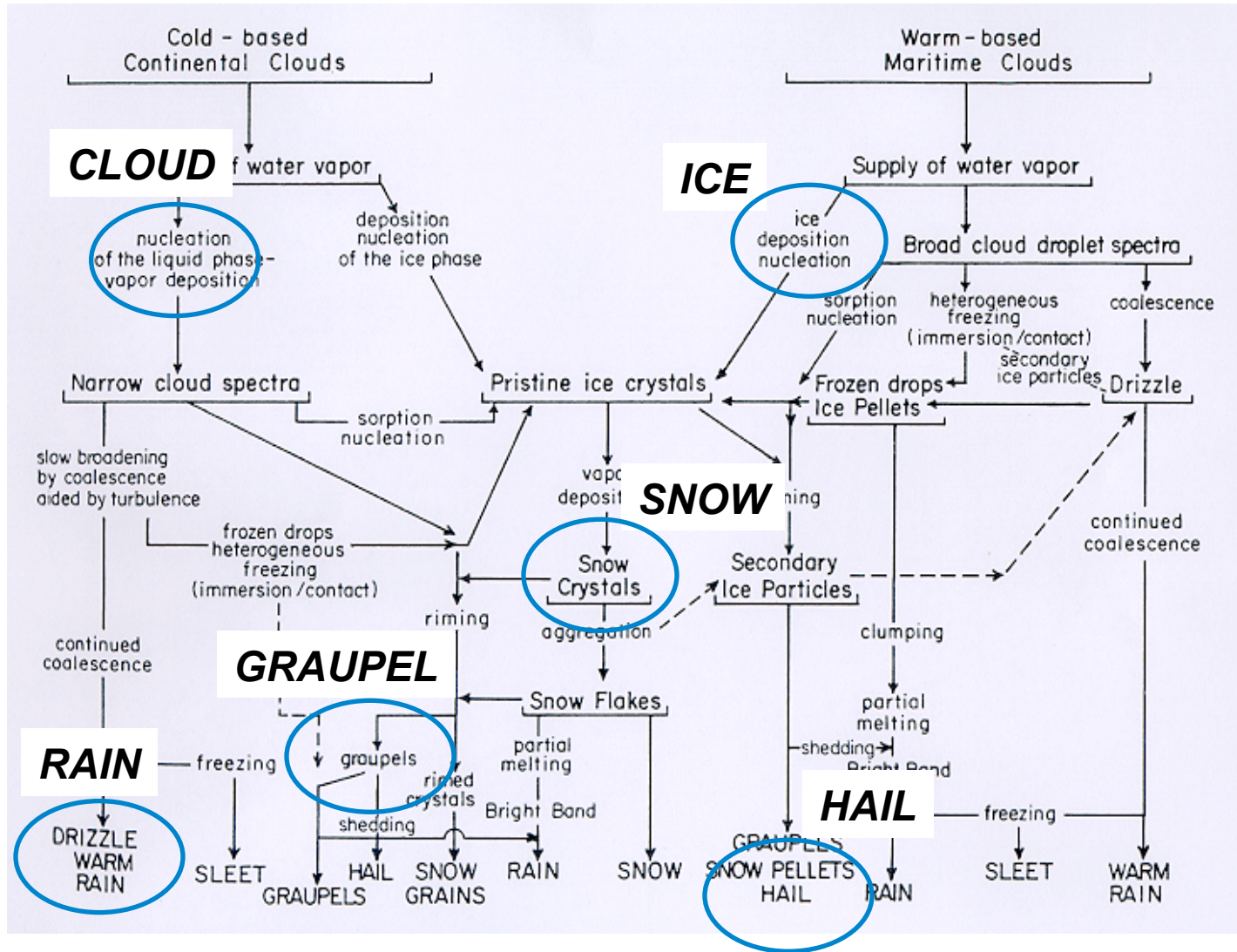


# Cloud Microphysical Processes



# Microphysics Parameterization Schemes

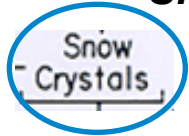
Hydrometeors are traditionally partitioned into categories



# Microphysics Parameterization Schemes

The particle size distributions are modeled

e.g. **SNOW**

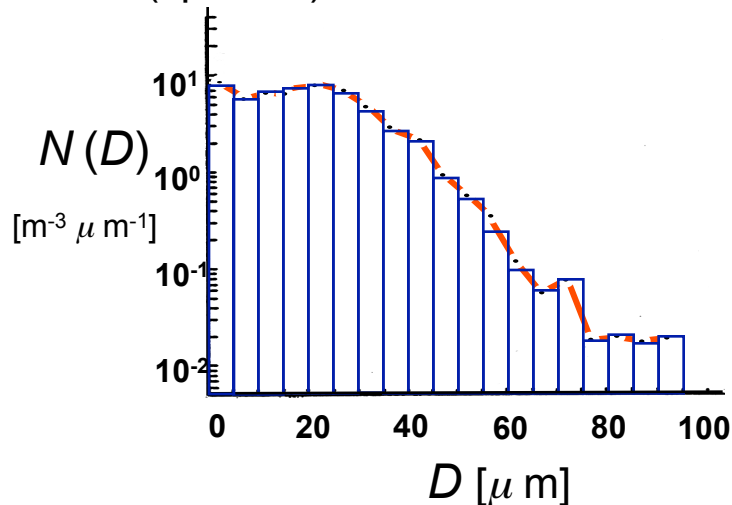


For each category, microphysical processes are parameterized to predict the evolution of the particle size distribution,  $N(D)$

## TYPES of SCHEMES:

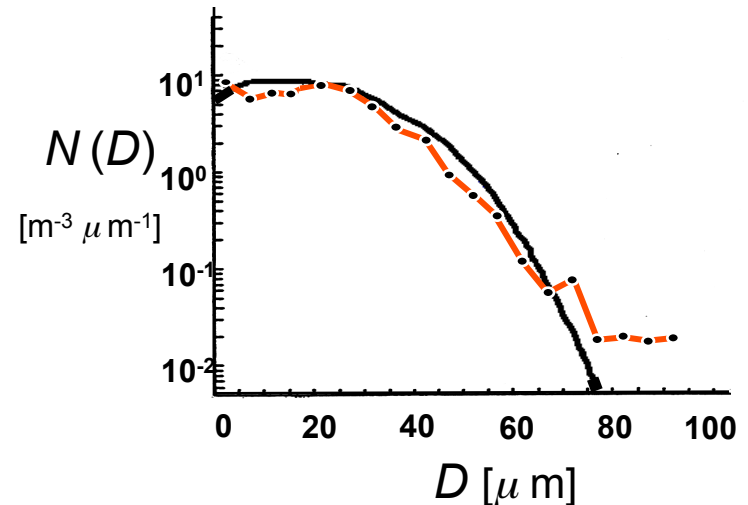
**Bin-resolving:**  
(spectral)

$$N(D) = \sum_{i=1}^I N_i$$

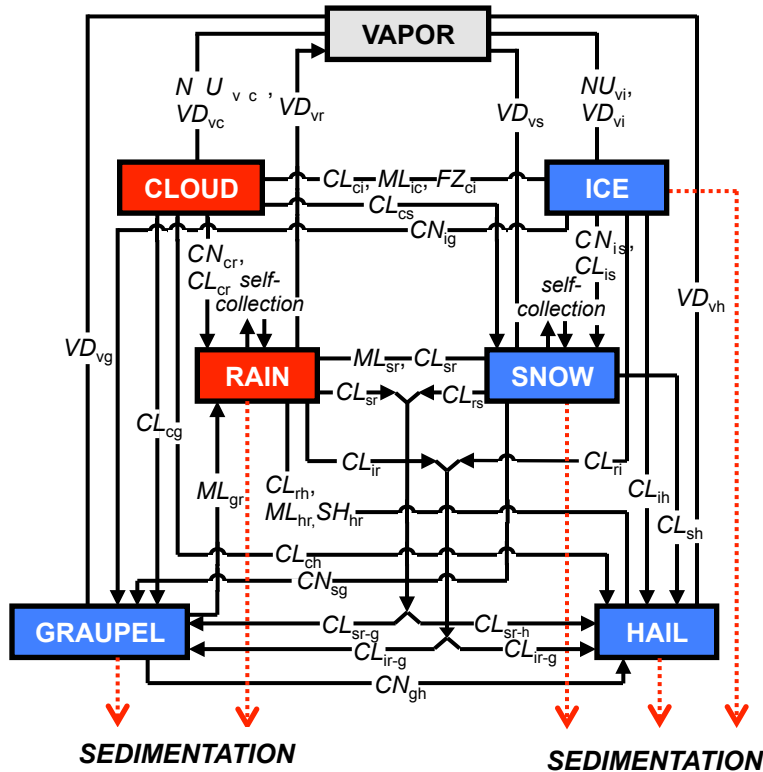


**Bulk:**

$$N(D) = N_0 D^\alpha e^{-\lambda D}$$



# Bulk Microphysics Scheme\* in GEM (HRDPS)



## Six hydrometeor categories:

2 liquid: **cloud, rain**

4 frozen: **ice, snow, graupel, hail**

## For each category $x = c, r, i, s, g, h$ :

$$N_x(D) = N_{0x} D^{\alpha_x} e^{-\lambda_x D} \quad (\text{complete})$$

$$V_x(D) = a_x D^{b_x} \quad (\text{empirical})$$

$$m_x(D) = c_x D^{d_x}$$

## Scheme version

double-moment

## Prognostic variables

$q_x, N_x$  (12)

triple-moment

$q_x, N_x, Z_x$  (17)

\* Milbrandt and Yau (2005)



# Traditional bulk approach for the ice phase

## Problems with pre-defined categories:

1. Real ice particles have complex shapes
2. Physics applied is often inconsistent
3. Conversion between categories is ad-hoc and leads to large, discrete changes in particle properties



### **CLOUD "ICE"**

$$\rho_s = 500 \text{ kg m}^{-3}$$

$$m = (\pi/6 \rho_s) D^3$$

$$V = a_i D^{b_i}$$

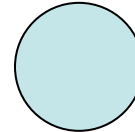


### **"SNOW"**

$$\rho_s = 100 \text{ kg m}^{-3}$$

$$m = c D^2$$

$$V = a_s D^{b_s}$$

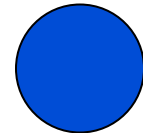


### **GRAUPEL**

$$\rho_g = 400 \text{ kg m}^{-3}$$

$$m = (\pi/6 \rho_g) D^3$$

$$V = a_g D^{b_g}$$



### **HAIL**

$$\rho_h = 900 \text{ kg m}^{-3}$$

$$m = (\pi/6 \rho_h) D^3$$

$$V = a_h D^{b_h}$$

← **abrupt / unphysical conversions**

NOTE: *Bin microphysics schemes have the identical problem*

# *New Bulk Microphysics Parameterization:* **Predicted Particle Properties (P3)\***

*Based on a conceptually different approach to  
parameterize ice-phase microphysics.*

## **NEW CONCEPT**

**“free” category** – predicted properties, thus freely evolving type

vs.

**“fixed” category** – traditional; prescribed properties, pre-determined type

## **Compared to traditional (ice-phase) schemes, P3:**

- avoids some necessary evils (ad-hoc category conversion, fixed properties)
- has self-consistent physics
- is better linked to observations
- Is more computationally efficient

\* Morrison and Milbrandt (2015)  
Milbrandt and Morrison (2016)

## Prognostic Variables: (advected)

**LIQUID PHASE:**    *2 categories, 2-moment:*

$Q_c$  – cloud mass mixing ratio                      [kg kg<sup>-1</sup>]

$Q_r$  – rain mass mixing ratio                            [kg kg<sup>-1</sup>]

$N_c$  – cloud number mixing ratio                      [#kg<sup>-1</sup>]

$N_r$  – rain number mixing ratio                        [#kg<sup>-1</sup>]

**ICE PHASE:**                      *nCat categories, 4 prognostic variables each:*

$Q_{dep}(n)^*$  – deposition ice mass mixing ratio    [kg kg<sup>-1</sup>]

$Q_{rim}(n)$  – rime ice mass mixing ratio                    [kg kg<sup>-1</sup>]

$N_{tot}(n)$  – total ice number mixing ratio                [ # kg<sup>-1</sup>]

$B_{rim}(n)$  – rime ice volume mixing ratio                [m<sup>3</sup> kg<sup>-1</sup>]

\*  $Q_{tot} = Q_{dep} + Q_{rim}$ , total ice mass mixing ratio (actual advected variable)

A given (*free*) category can represent any type of ice-phase hydrometeor

## **Prognostic Variables:**

$Q_{dep}$ – deposition ice mass mixing ratio	[kg kg <sup>-1</sup> ]
$Q_{rim}$ – rime ice mass mixing ratio	[kg kg <sup>-1</sup> ]
$N_{tot}$ – total ice number mixing ratio	[# kg <sup>-1</sup> ]
$B_{rim}$ – rime ice volume mixing ratio	[m <sup>3</sup> kg <sup>-1</sup> ]

## **Predicted Properties:**

$F_{rim}$ – rime mass fraction, $F_{rim} = Q_{rim} / (Q_{rim} + Q_{dep})$	[--]
$\rho_{rim}$ – rime density, $\rho_{rim} = Q_{rim} / B_{rim}$	[kg m <sup>-3</sup> ]
$D_m$ – mean-mass diameter, $D_m \propto Q_{tot} / N_{tot}$	[m]
$V_m$ – mass-weighted fall speed, $V_m = f(D_m, \rho_{rim}, F_{rim})$	[m s <sup>-1</sup> ]
<i>etc.</i>	

## **Diagnostic Particle Types:**

Based on the predicted properties (rather than pre-defined)



## GENERAL (all schemes)

$$Q^+ = Q^0 + \underbrace{\Delta Q|_{PROC\_1} + \Delta Q|_{PROC\_2} + \dots}$$

$$\Delta Q|_{PROC\_1} = \Delta t \cdot \underbrace{\frac{1}{\rho} \int_0^\infty \frac{dm(D)}{dt} |_{PROC\_1} N(D) dD}$$

$\propto M^{(p)}$  (and other moments)

Computing the tendencies for the prognostic variables (i.e. process rates) essentially amounts to computing various moments of  $N(D)$

**Predicting process rates for  $V_x \rightarrow$  computing various  $M_x^{(p)}$**

$V$  = prognostic variable ( $Q, N, \dots$ )  
 $x$  = category (rain, ice, ...)

## Predicting process rates ~ computing $M_x^{(p)}$

### TRADITIONAL SCHEMES (e.g. 2-moment)

$$M^{(p)} \equiv \int_0^\infty D^p N_x(D) dD = N_{0x} \frac{\Gamma(1 + \mu_x + p)}{\lambda_x^{p+1+\mu_x}}$$

Fixed category  $\Rightarrow$  constant  $m$ - $D$  parameters

$$m(D) = \alpha D^\beta$$

$$Q = \frac{1}{\rho} \int_0^\infty m(D) N(D) dD = \frac{1}{\rho} \int_0^\infty \alpha D^\beta N_x(D) dD = \frac{\alpha}{\rho} M^{(\beta)} = \frac{\alpha}{\rho} N_{0x} \frac{\Gamma(1 + \mu_x + \beta)}{\lambda_x^{1+\mu_x+\beta}}$$

$$N = \int_0^\infty N_x(D) dD = M^{(0)} = N_{0x} \frac{\Gamma(1 + \mu_x)}{\lambda_x^{1+\mu_x}}$$

- impose assumption about  $\mu$
- 2 equations, 2 unknowns  $\rightarrow$  solve for  $\lambda$ ,  $N_0$

**$\rightarrow$  Now, any  $M^{(p)}$  can be computed analytically**

# Predicting process rates ~ computing $M_x^{(p)}$

## P3 SCHEME

$$M^{(p)} \equiv \int_0^\infty D^p N_x(D) dD = N_{0x} \frac{\Gamma(1 + \mu_x + p)}{\lambda_x^{p+1+\mu_x}}$$

**Free category**  $\Rightarrow$  variable  $m$ - $D$ ,  $A$ - $D$ , and  $V$ - $D$  parameters

$$Q = \frac{1}{\rho} \int_0^\infty m(D) N(D) dD = \frac{1}{\rho} \int_0^\infty \alpha D^\beta N_x(D) dD = \frac{\alpha}{\rho} M^{(\beta)} = \frac{\alpha}{\rho} N_{0x} \frac{\Gamma(1 + \mu_x + \beta)}{\lambda_x^{1+\mu_x+\beta}}$$

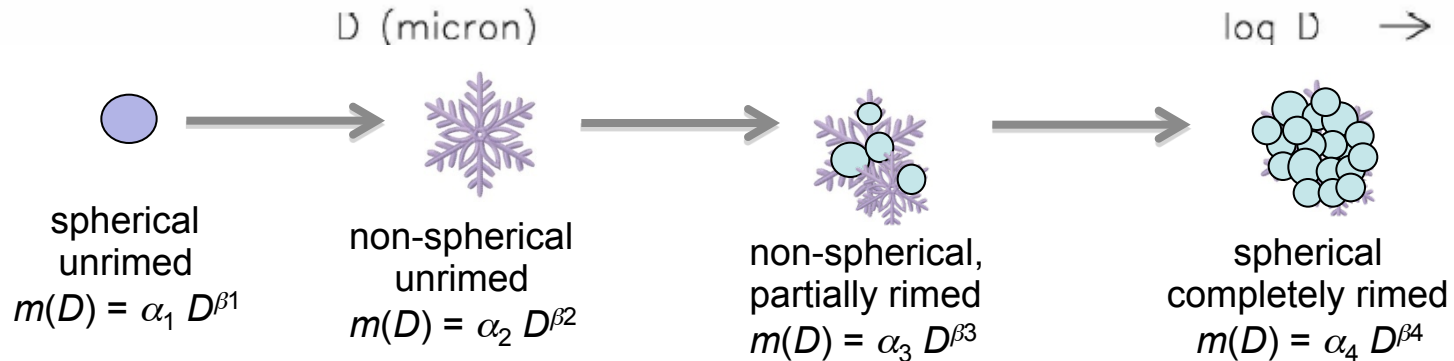
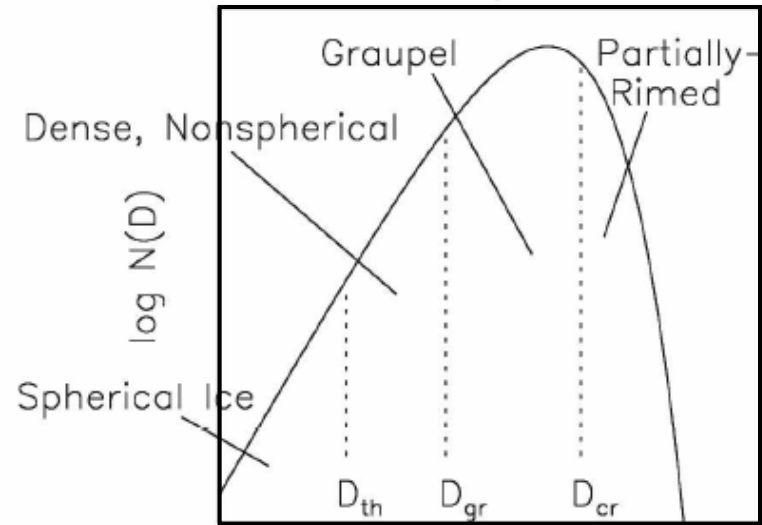
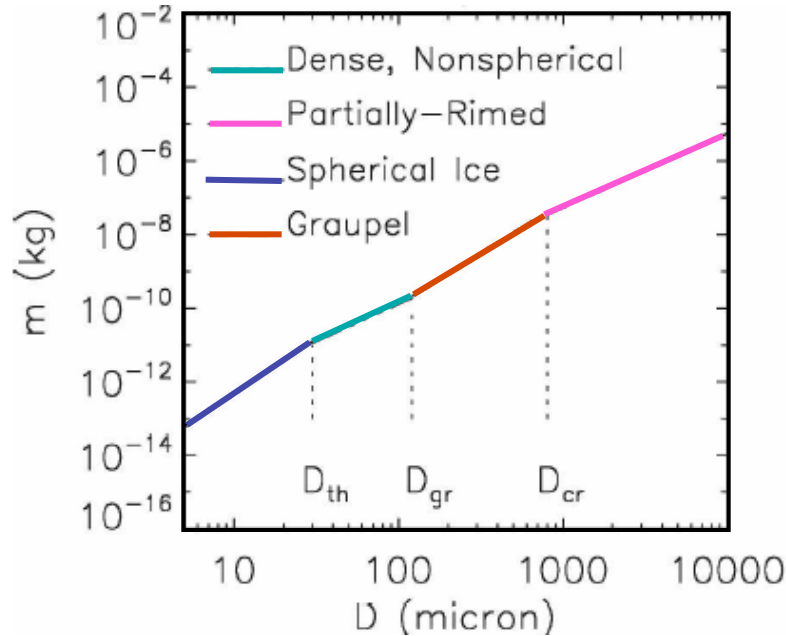
$$N = \int_0^\infty N_x(D) dD = M^{(0)} = N_{0x} \frac{\Gamma(1 + \mu_x)}{\lambda_x^{1+\mu_x}}$$

$\rightarrow$  cannot compute  $Q$  (or any other  $M^{(p)}$ ) analytically

# Predicting process rates ~ computing $M_x^{(p)}$

## P3 SCHEME – Determining $m(D) = \alpha D^\beta$ for regions of $D$ :

General:  $1 > F_{rim} > 0$ ; for a given  $\rho_{rim}$





## Predicting process rates ~ computing $M_x^{(p)}$

### P3 SCHEME – Computing $N(D)$ parameters :

1. Compute properties  $F_{rim} = Q_{rim}/(Q_{dep}+Q_{rim})$ ,  $\rho_{rim} = Q_{rim} / B_{rim}$
2. Determine integral ranges,  $D_{th}$ ,  $D_{gr}$ ,  $D_{cr}$
3. Determine PSD parameters ( $\lambda$ ,  $N_0$ ,  $\mu$ )
  - solved numerically (iteratively; pre-computed and stored in look-up table)

$$Q = \frac{1}{\rho} \left[ \int_0^{D_{th}} \alpha_1 D^{\beta_1+\mu} e^{-\lambda D} dD + \int_{D_{th}}^{D_{gr}} \alpha_2 D^{\beta_2+\mu} e^{-\lambda D} dD + \int_{D_{gr}}^{D_{cr}} \alpha_3 D^{\beta_3+\mu} e^{-\lambda D} dD + \int_{D_{cr}}^{\infty} \alpha_4 D^{\beta_4+\mu} e^{-\lambda D} dD \right]$$

$$N = N_{0x} \frac{\Gamma(1 + \mu_x)}{\lambda_x^{1+\mu_x}}$$

$$\mu = f(\lambda)$$

4. Also, match  $A$ - $D$  parameters to  $m$ - $D$  parameters for the various regions of  $D$ 
  - based on geometric + empirical relations
  - for  $V$ - $D$  (process rates and sedimentation) and  $r_{i\_eff}$  (optical properties)

## Predicting process rates ~ computing $M_x^{(p)}$

### P3 SCHEME – Computing the process rates:

Now, have  $\lambda$ ,  $N_0$ ,  $\mu$ , and integral ranges  $D_{th}$ ,  $D_{gr}$ ,  $D_{cr}$  (plus  $\alpha_{(i)}$ ,  $\beta_{(i)}$ , ...)

$$Q^+ = Q^0 + \underbrace{\Delta Q|_{PROC\_1} + \Delta Q|_{PROC\_2} + \dots}$$

$$\Delta Q|_{PROC\_1} = \Delta t \cdot \underbrace{\frac{1}{\rho} \int_0^\infty \frac{dm(D)}{dt} \Big|_{PROC\_1} N(D) dD}$$

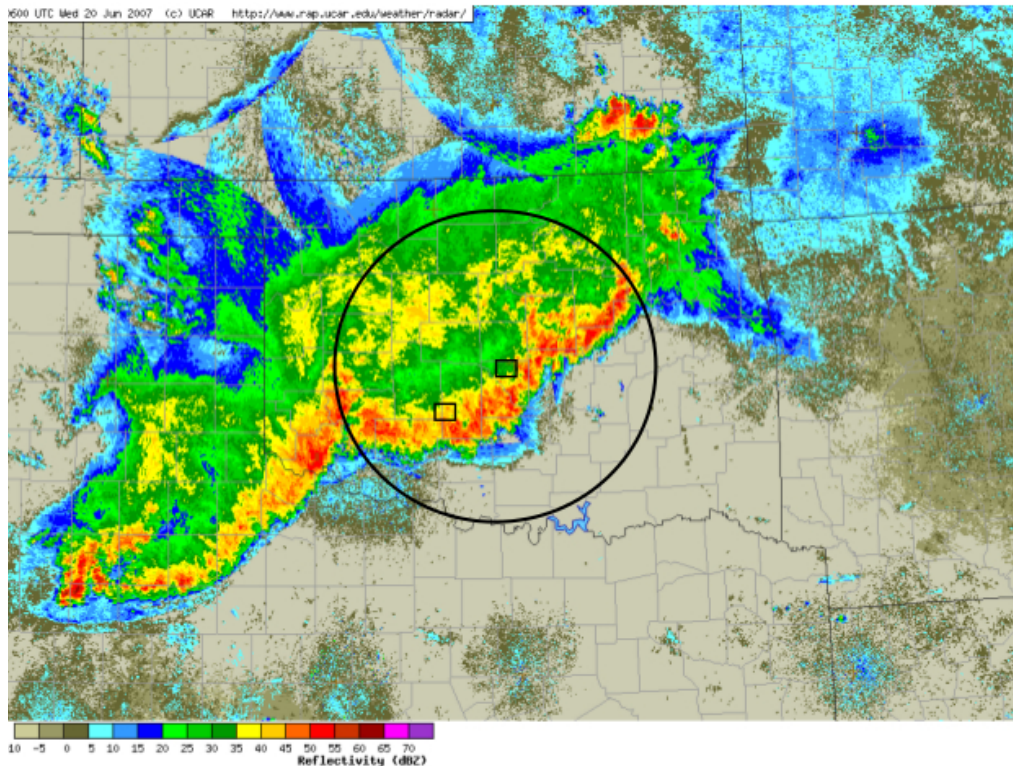
$$\propto X_1 \quad (\text{and } X_2, \dots)$$

$$X_1 = \int_0^{D_{th}} D^a N_0 e^{-\lambda D} f(\alpha_1, \beta_1, \dots) dD + \int_{D_{th}}^{D_{gr}} D^b N_0 e^{-\lambda D} f(\alpha_2, \beta_2, \dots) dD \\ + \int_{D_{gr}}^{D_{cr}} D^c N_0 e^{-\lambda D} f(\alpha_3, \beta_3, \dots) dD + \int_{D_{cr}}^\infty D^d N_0 e^{-\lambda D} f(\alpha_4, \beta_4, \dots) dD$$

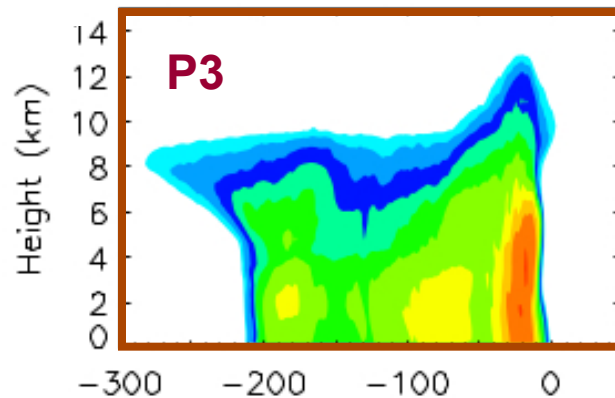
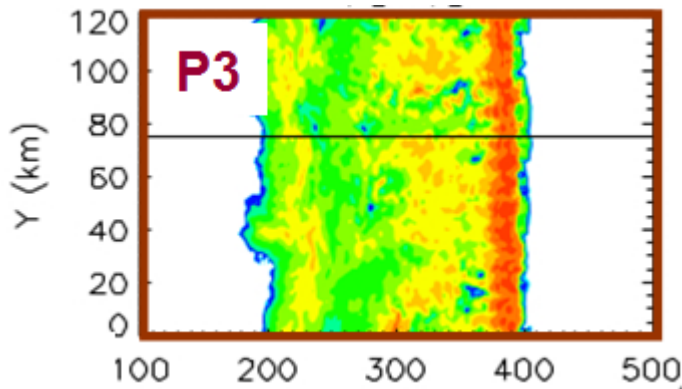
**Predicting process rates  $\rightarrow$  computing sums ( $X_n$ ) of partial moments**

# 3D Squall Line case: (June 20, 2007 central Oklahoma)

- WRF\_v3.4.1,  $\Delta x = 1$  km,  $\Delta z \sim 250$ -300 m, 112 x 612 x 24 km domain
- initial sounding from observations
- convection initiated by  $u$ -convergence
- no radiation, surface fluxes

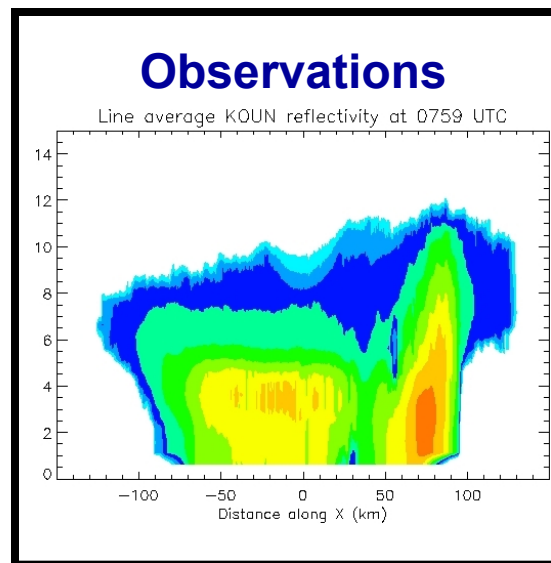
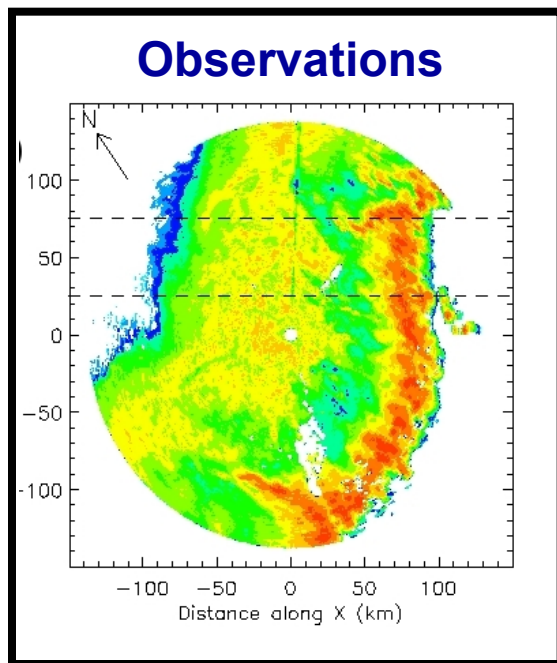
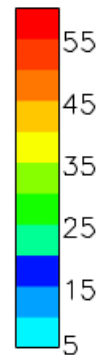


# 1-km WRF Simulations with P3 microphysics (1 category):



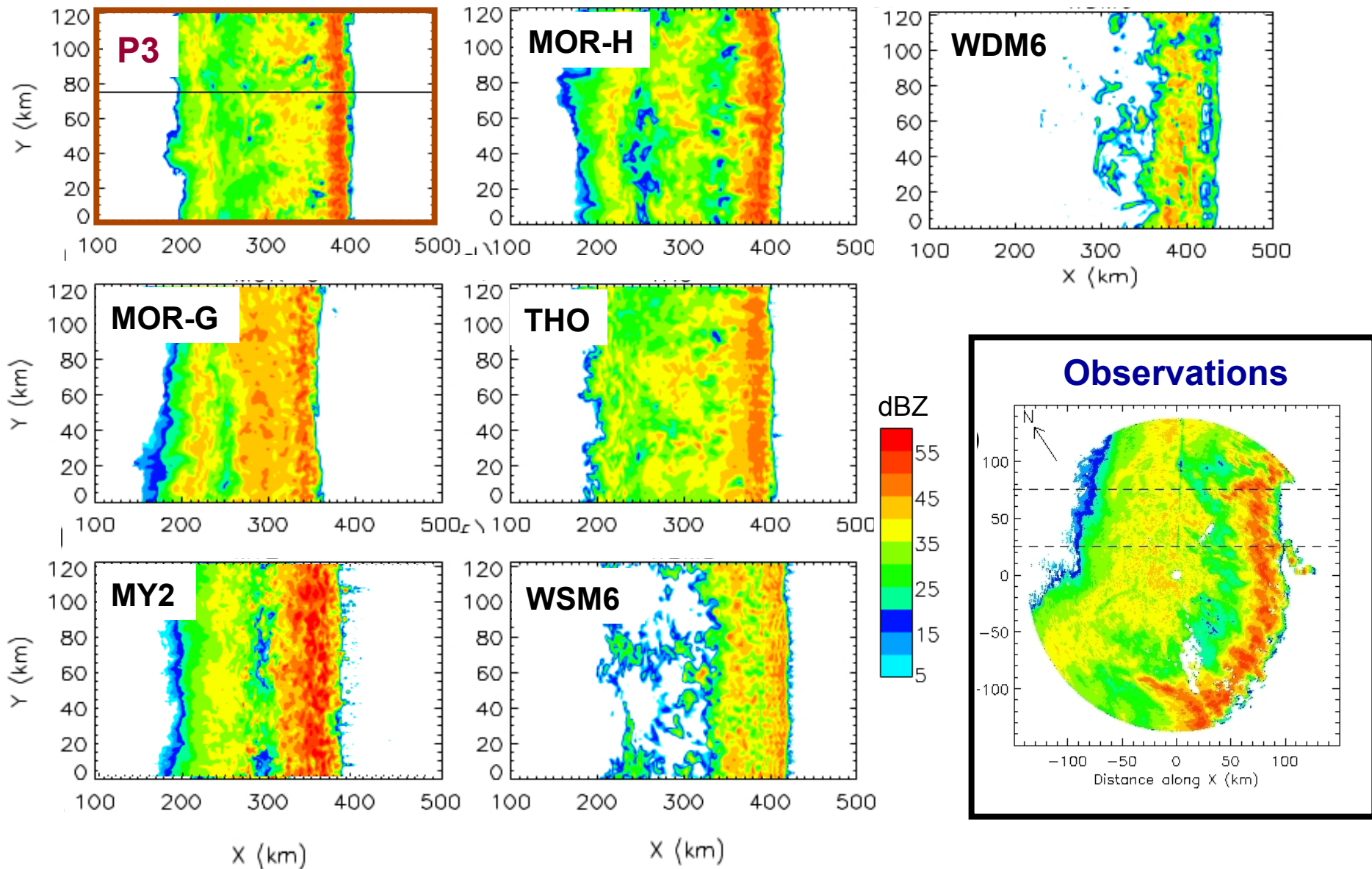
Reflectivity

dBZ

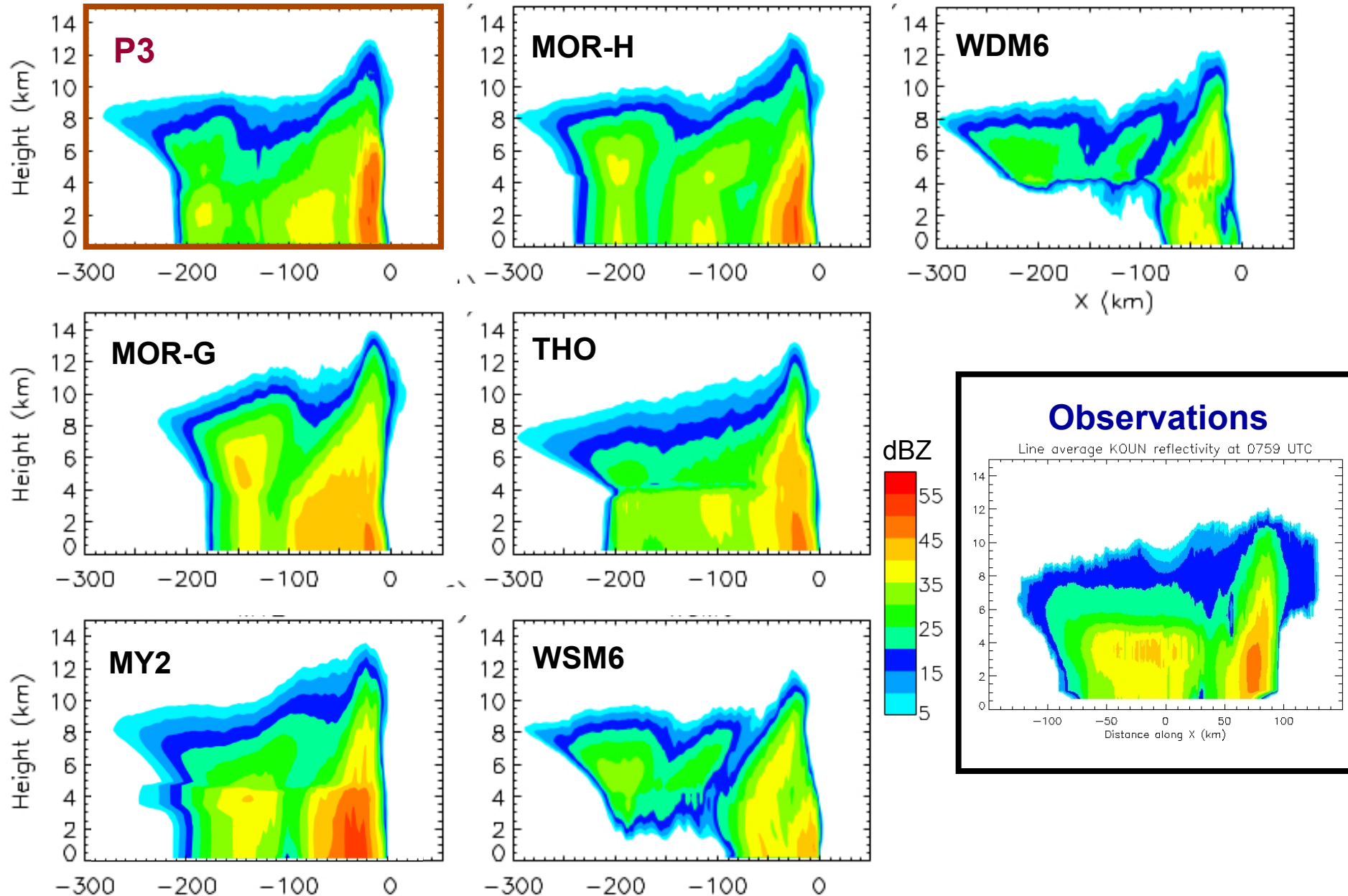




# WRF Results: Base Reflectivity (1 km AGL, t = 6 h)



# WRF Results: Line-averaged Reflectivity (t = 6 h)



## Vertical cross section of model fields ( $t = 6$ h)

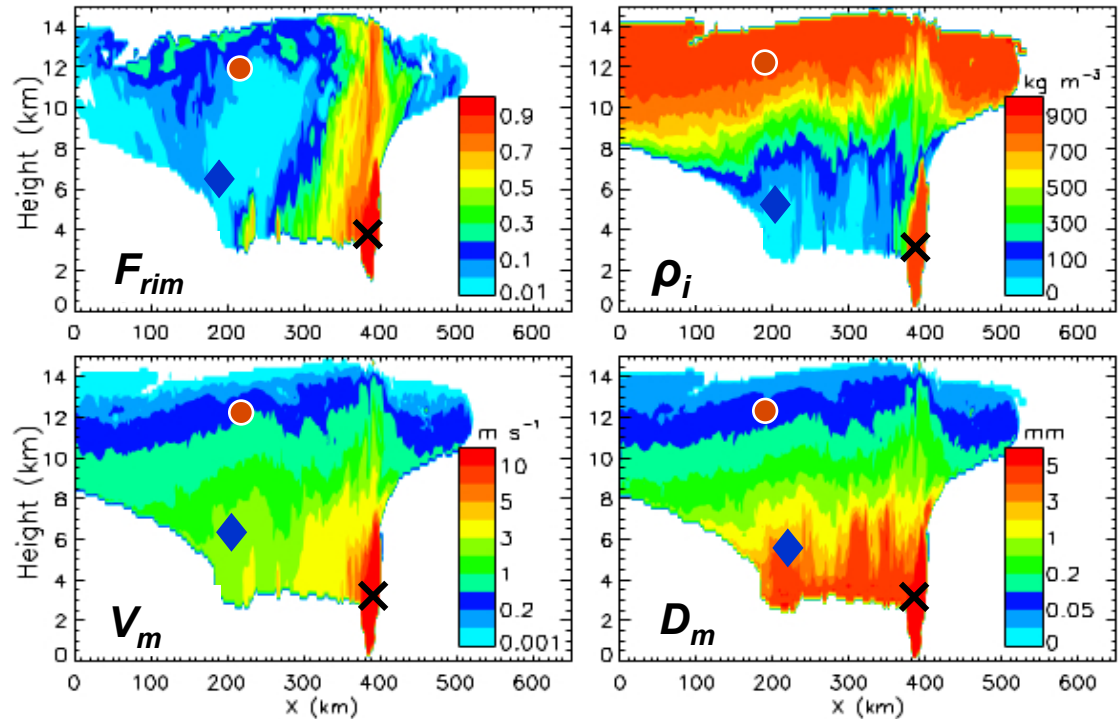
### Ice Particle Properties:

$F_r \sim 0-0.1$   
 $\rho \sim 900 \text{ kg m}^{-3}$   
 $V \sim 0.3 \text{ m s}^{-1}$   
 $D_m \sim 100 \text{ }\mu\text{m}$   
 $\rightarrow$  *small crystals*

$F_r \sim 0$   
 $\rho \sim 50 \text{ kg m}^{-3}$   
 $V \sim 1 \text{ m s}^{-1}$   
 $D_m \sim 3 \text{ mm}$   
 $\rightarrow$  *aggregates*

$F_r \sim 1$   
 $\rho \sim 900 \text{ kg m}^{-3}$   
 $V > 10 \text{ m s}^{-1}$   
 $D_m > 5 \text{ mm}$   
 $\rightarrow$  *hail*

*etc.*



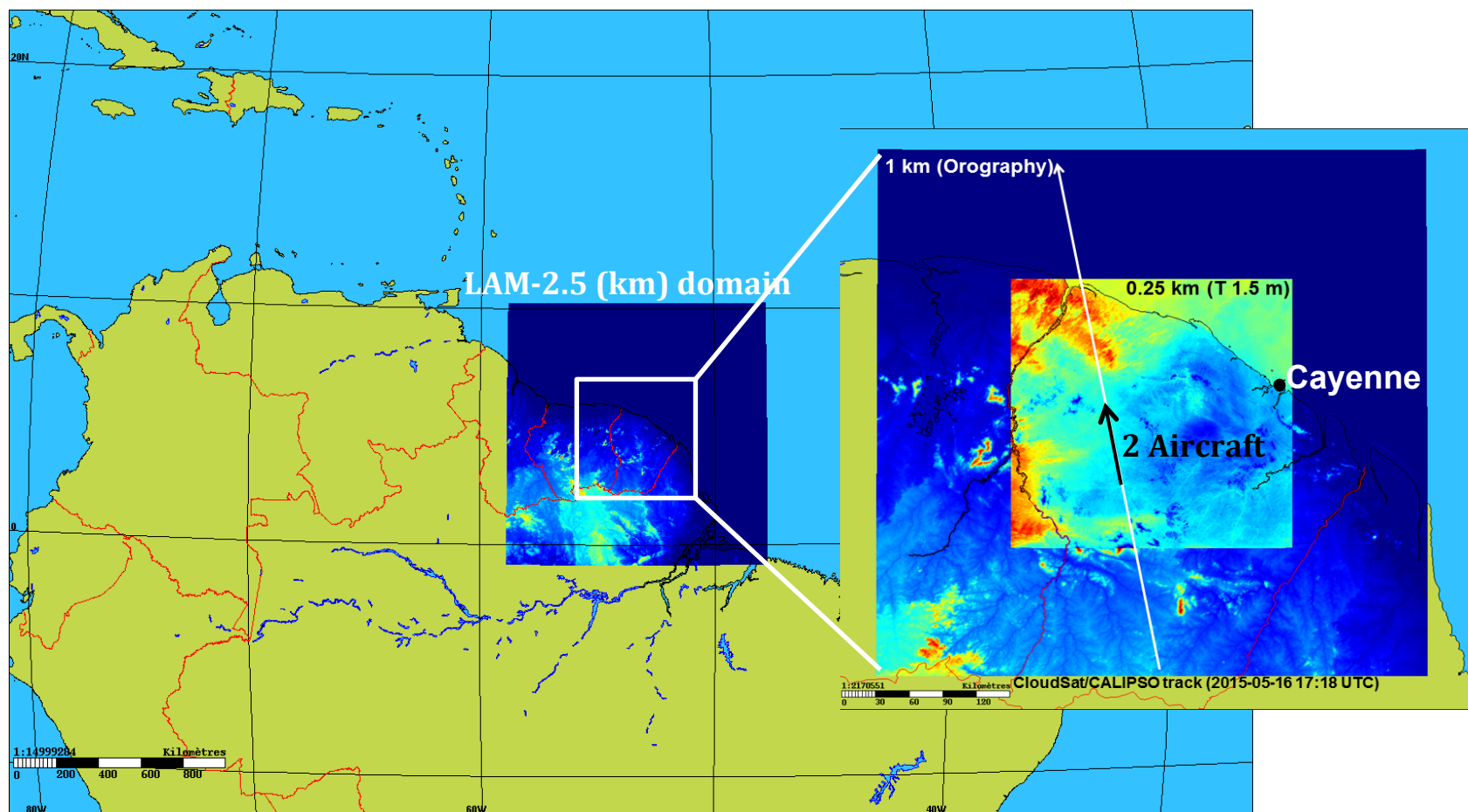
Note – only one (free) category

# 1. High Resolution NWP model at Environment Canada

Objective 2: Assessment of the hi-res NWP model.

Case: Cayenne, French Guiana (May 16, 2015)

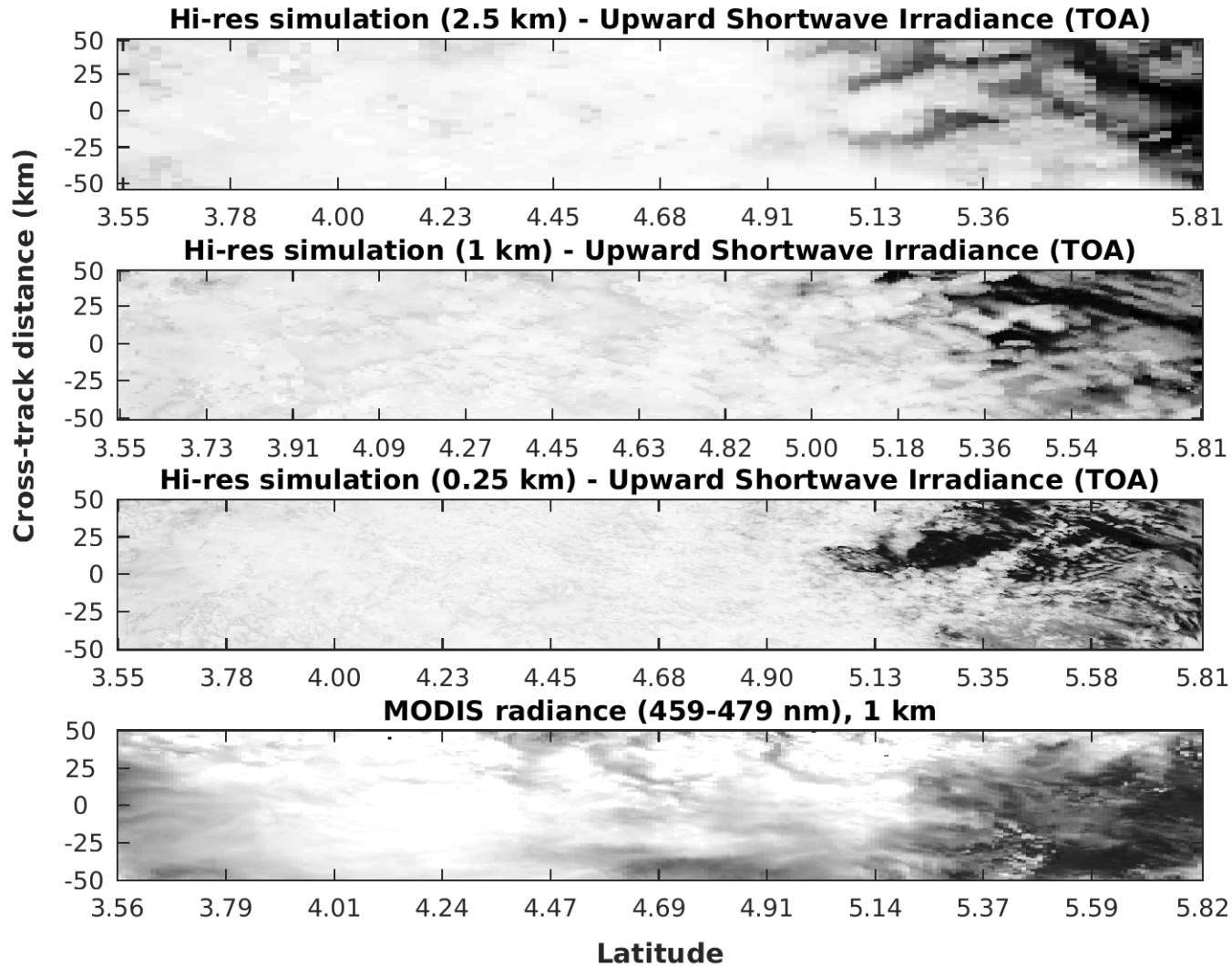
(Aircraft *in situ* measure, A-train overpasses, tropical deep convective cloud)

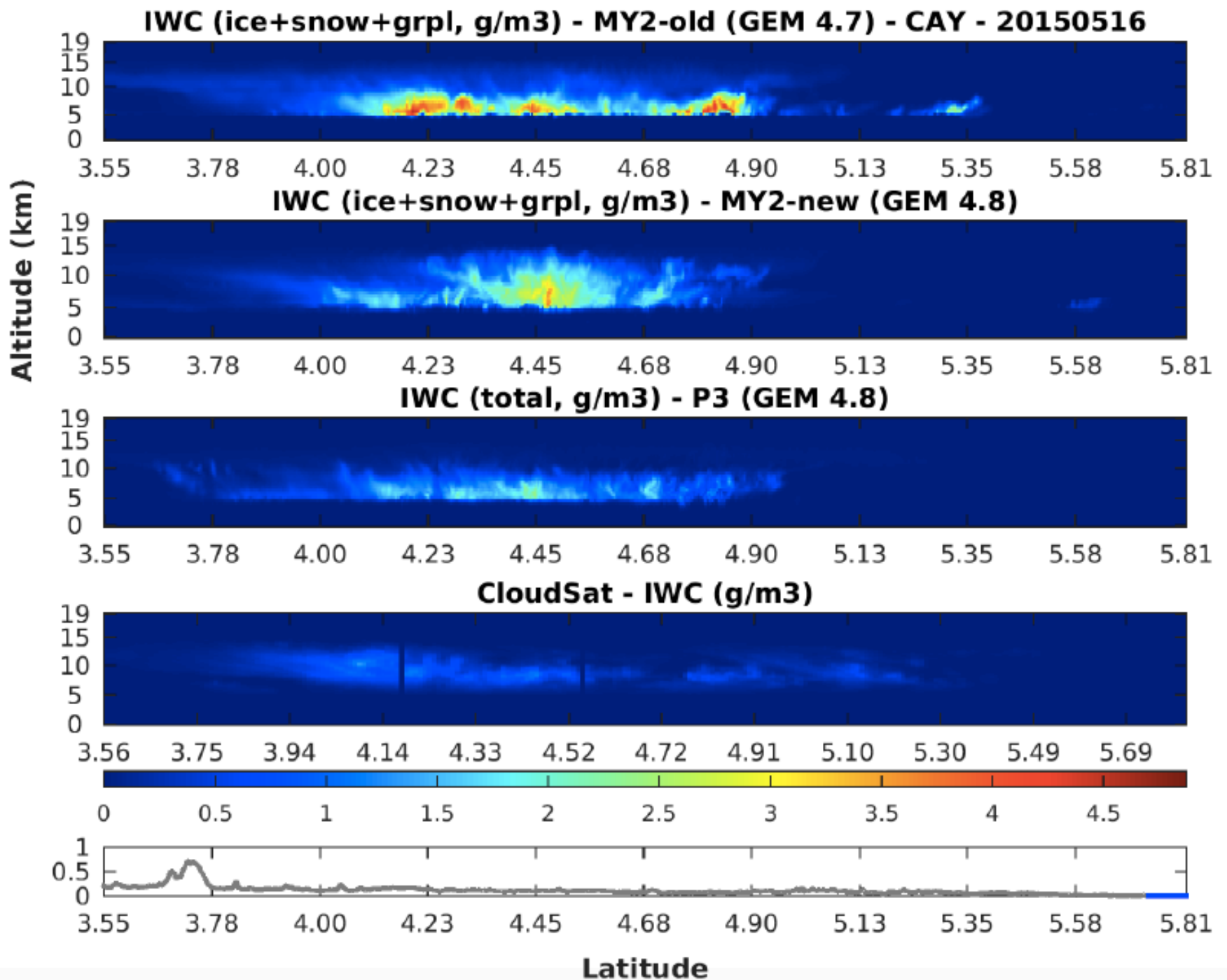




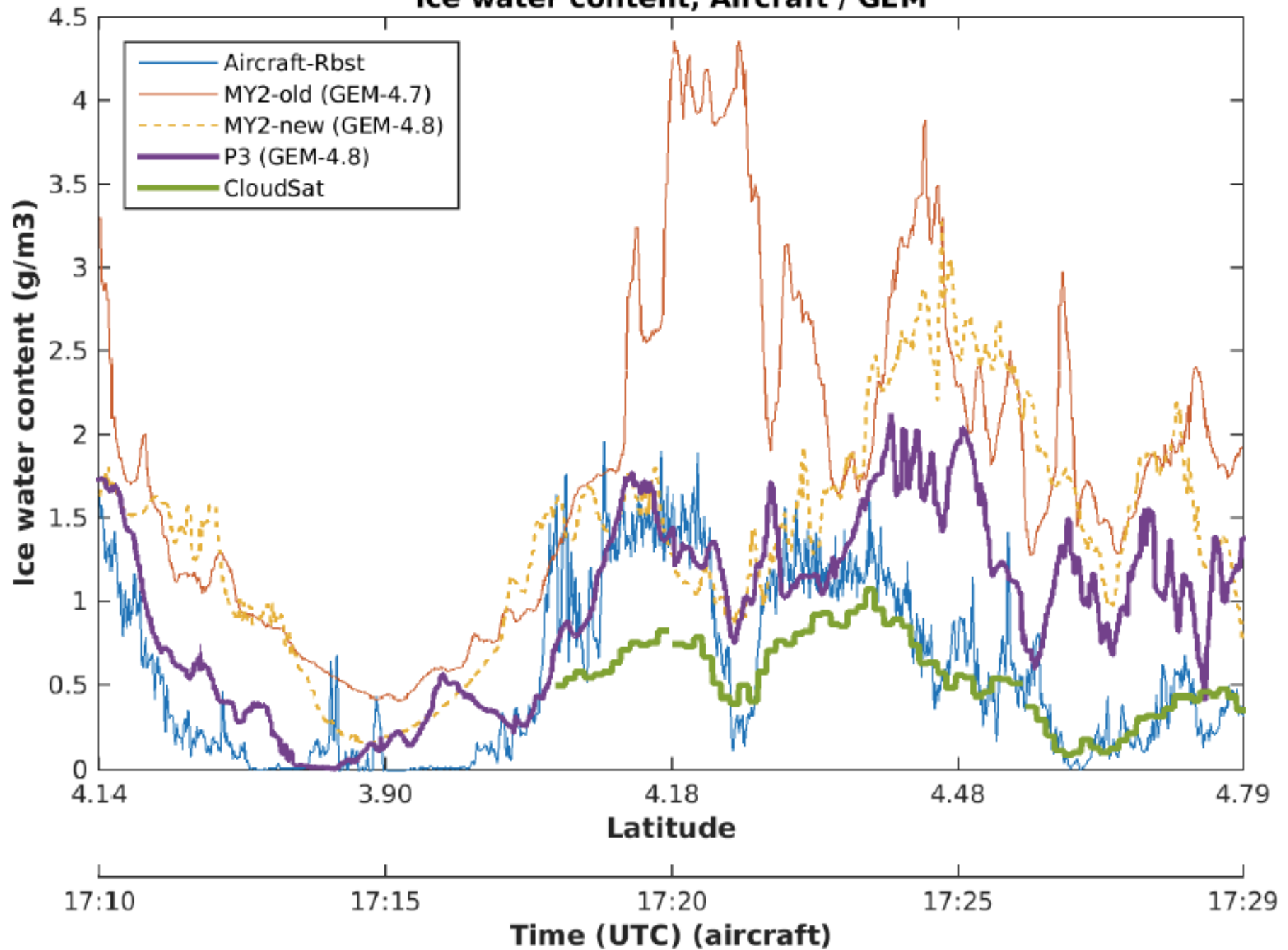
### 3. Case study

Model simulation for 17:20 UTC

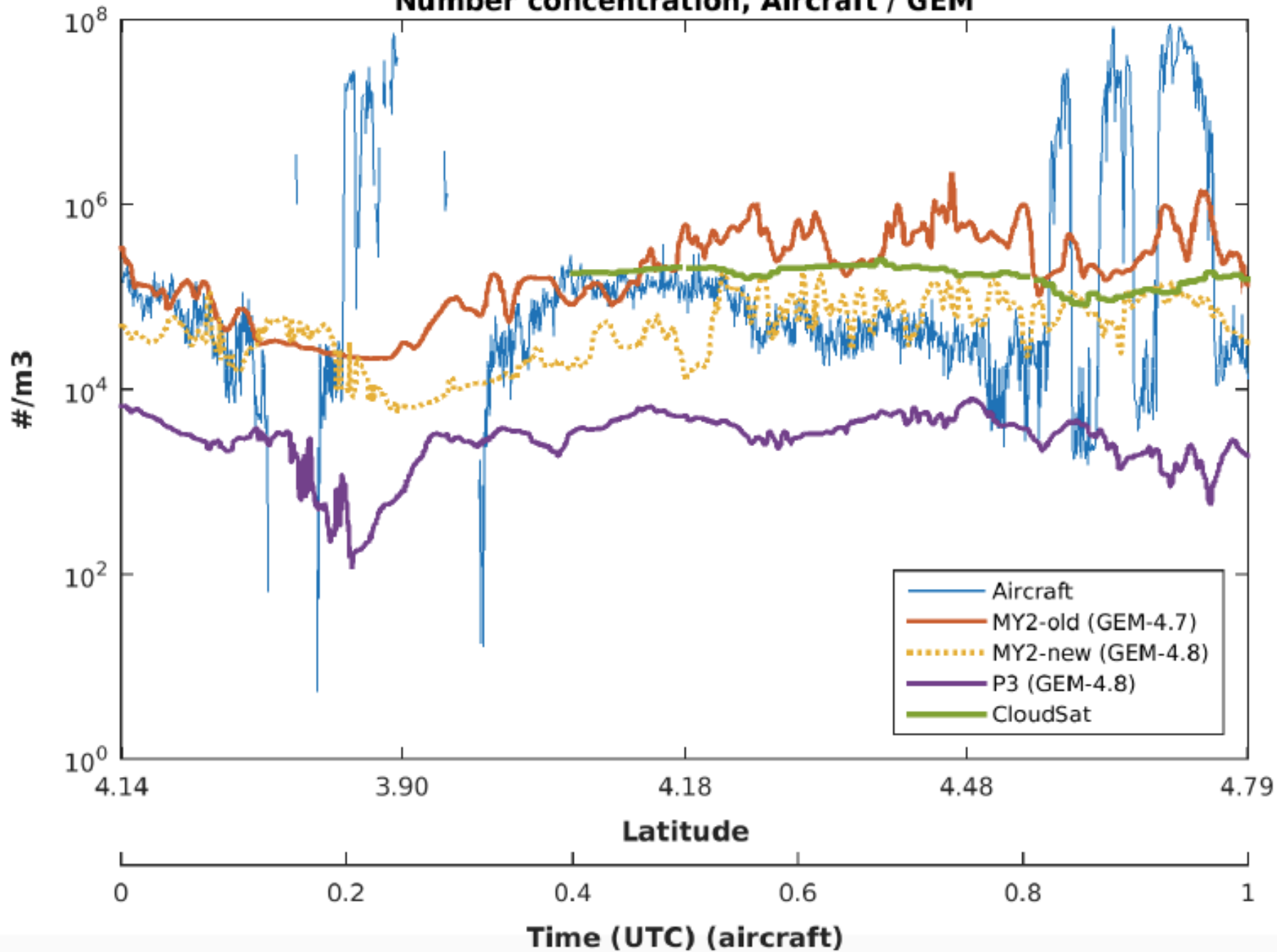




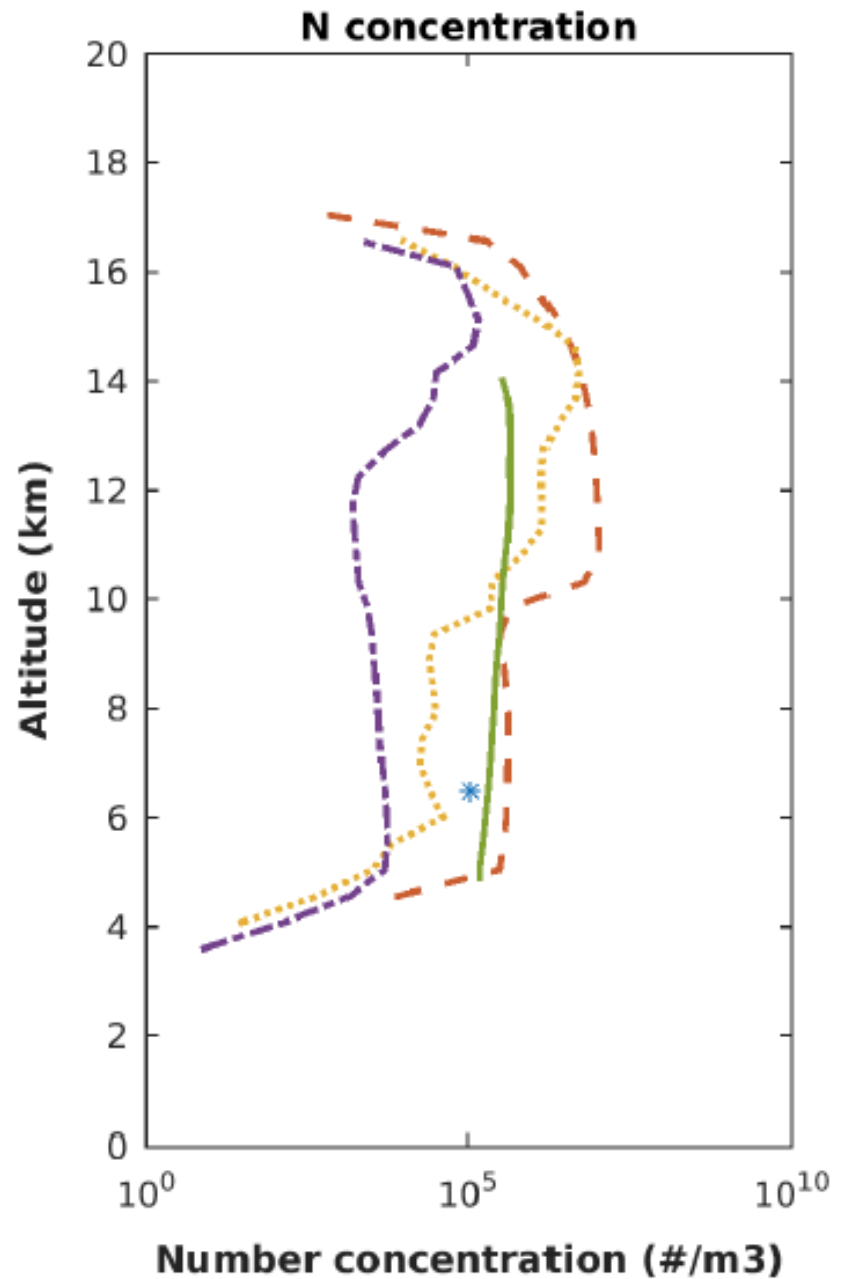
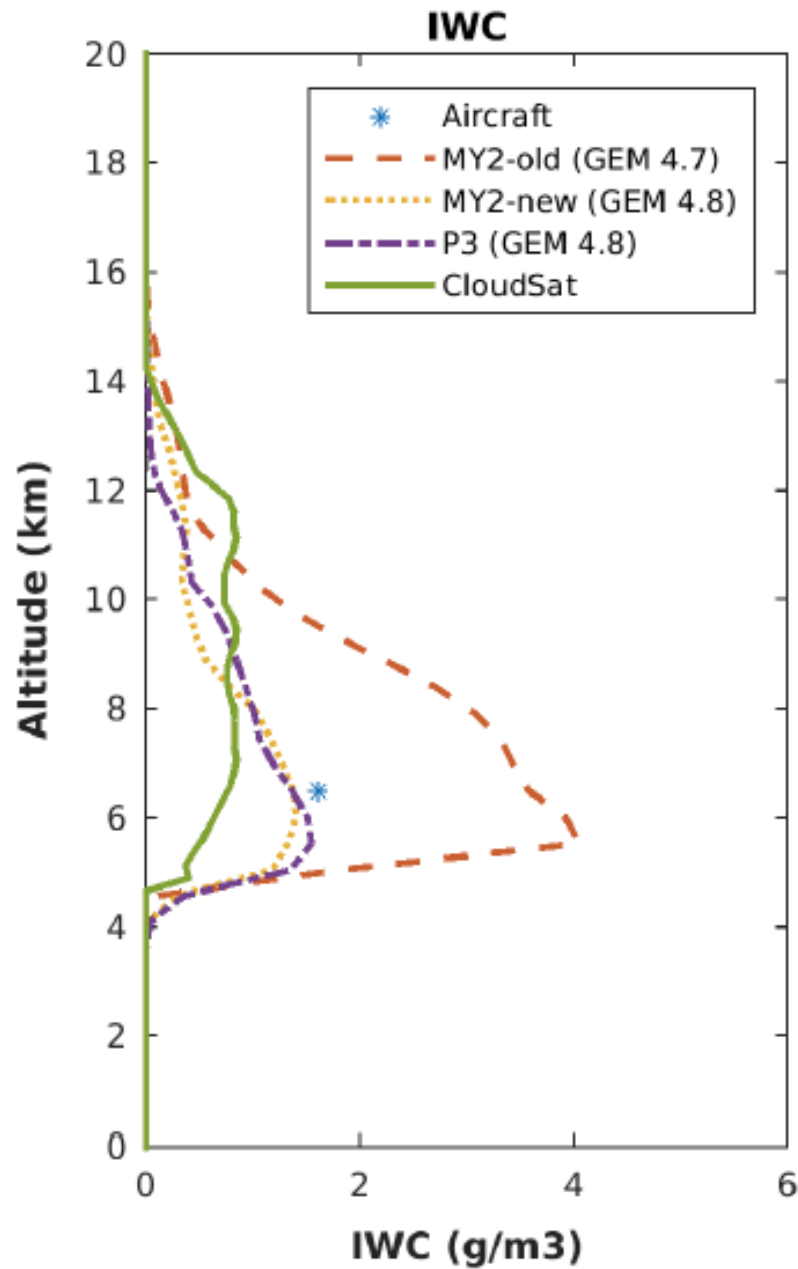
### Ice water content, Aircraft / GEM



# Number concentration, Aircraft / GEM



# Vertical profile at Lat: 4.17, Lon: -53.52



# Concluding comments

Operational NWP models are now at the convective scale ( $dx = 1-3$  km) which permits (requires) detailed bulk microphysics parameterizations (BMPs) and thus detailed treatment of cloud ice

New techniques in BMPs – such as the P3 approach to the representation of ice – show promise for the improvement in numerical guidance of fields related to ice-phase microphysics

For the development and improvement of BMPs – for research and operational NWP – field campaigns such as HAIC-HIWC and the related research are essential