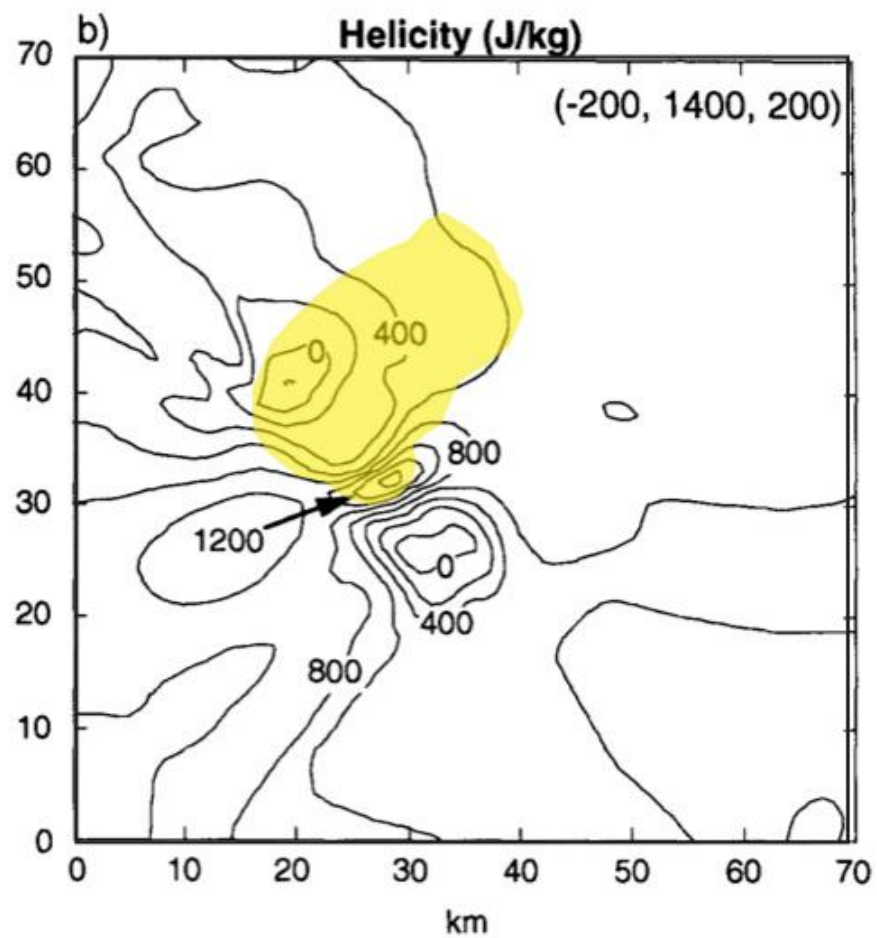
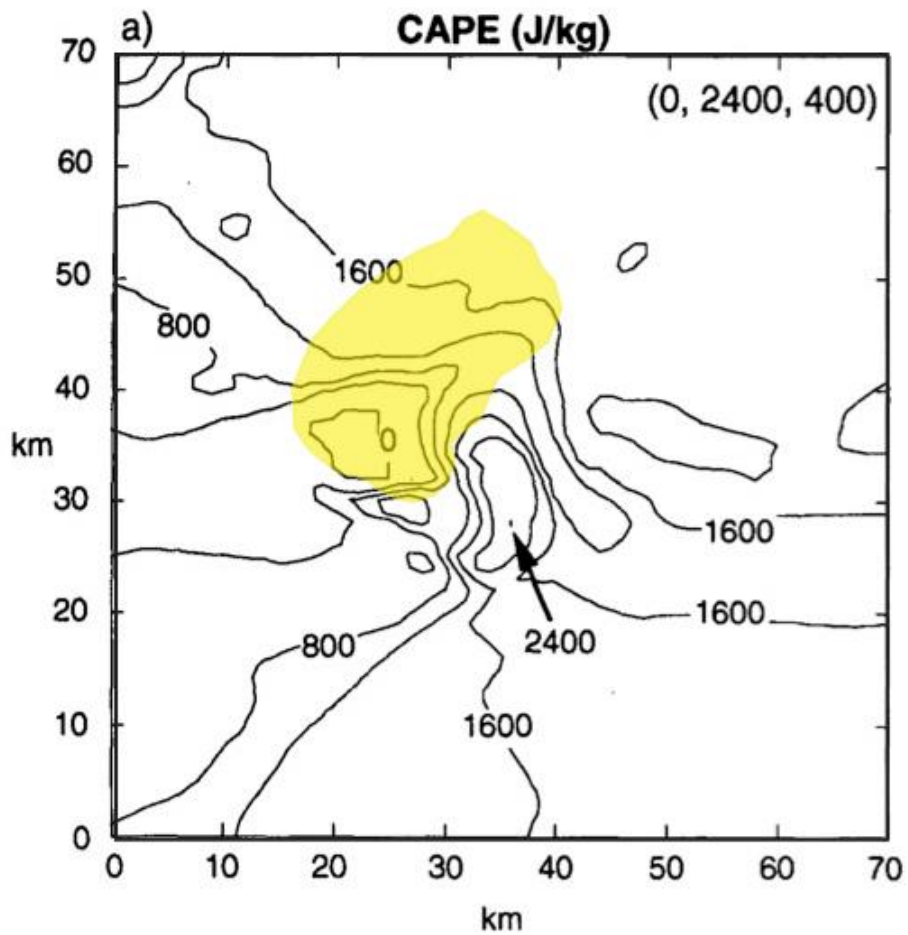


# Ability of Models to Reproduce Upscale Feedback of Deep Convection

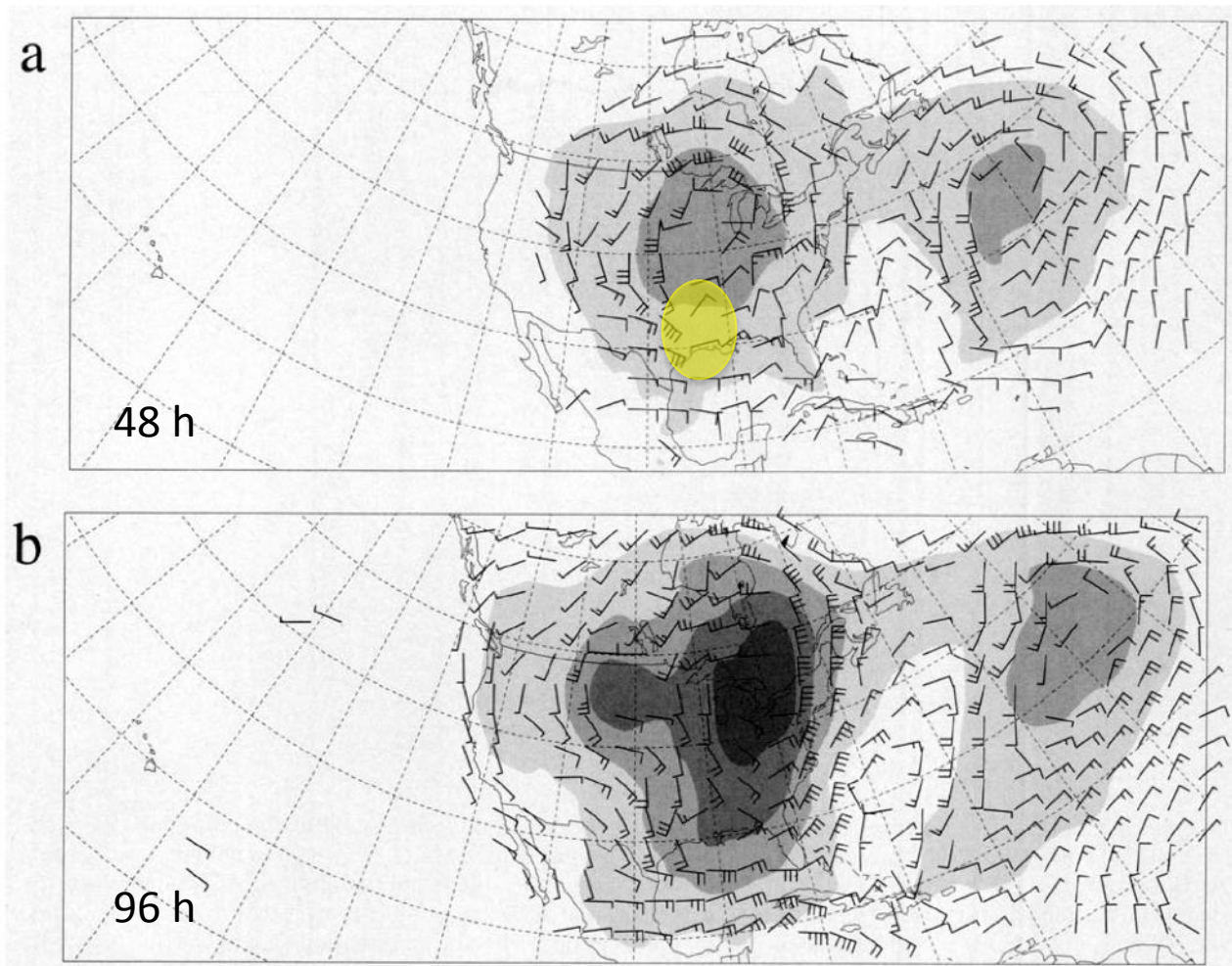
David J. Stensrud

NOAA/National Severe Storms Laboratory



From Brooks et al. (1994 WAF)

200 hPa Wind and Geopotential Height Differences  
Run with deep convection – run without deep convection



From Stensrud (1996 JAS)

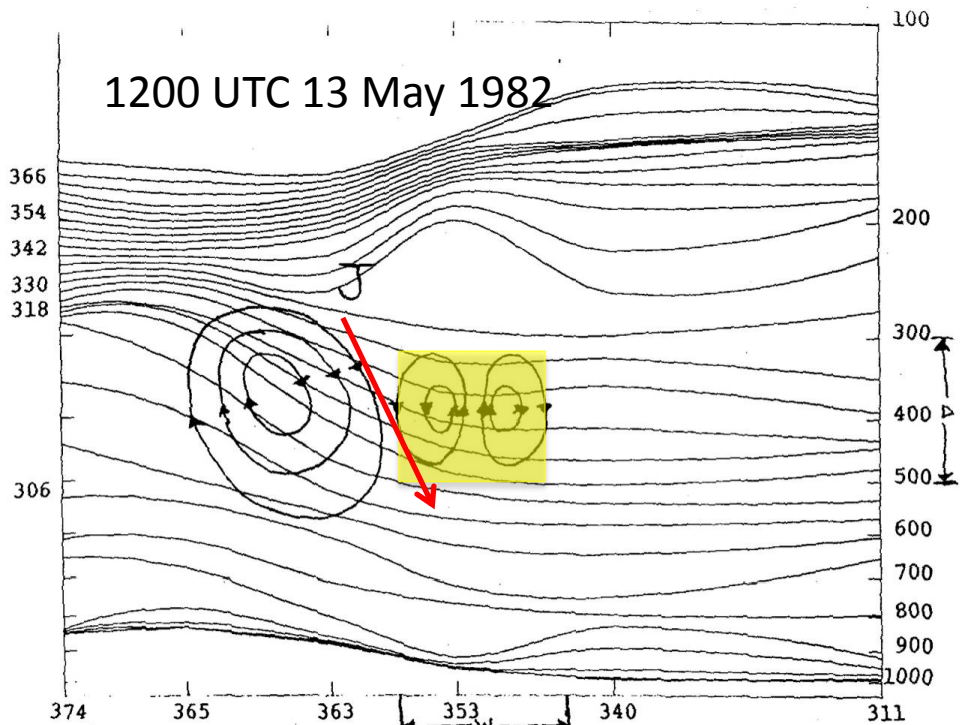


Fig. 10. Cross-section from Athens, Georgia, to Winslow, Arizona, from 1200 UTC 13 May 1982. Isentropes every 3 C. J denotes location of jet core, and solid lines with arrows denote transverse ageostrophic circulations. Hypothesized circulations due to diabatic heating are located in the region indicated by W and D. W is the estimated horizontal width of the anvil region; D is the estimated depth of the anvil. Location of the cross-section is from A-A' in Fig. 2 (d).



Figures from my final paper in Toby Carlson's Advanced Synoptic class. Vertical circulations diagnosed from Sawyer-Eliassen secondary circulation equations (Shapiro 1981). Circulation associated with deep convection would assist in formation of tropopause fold.

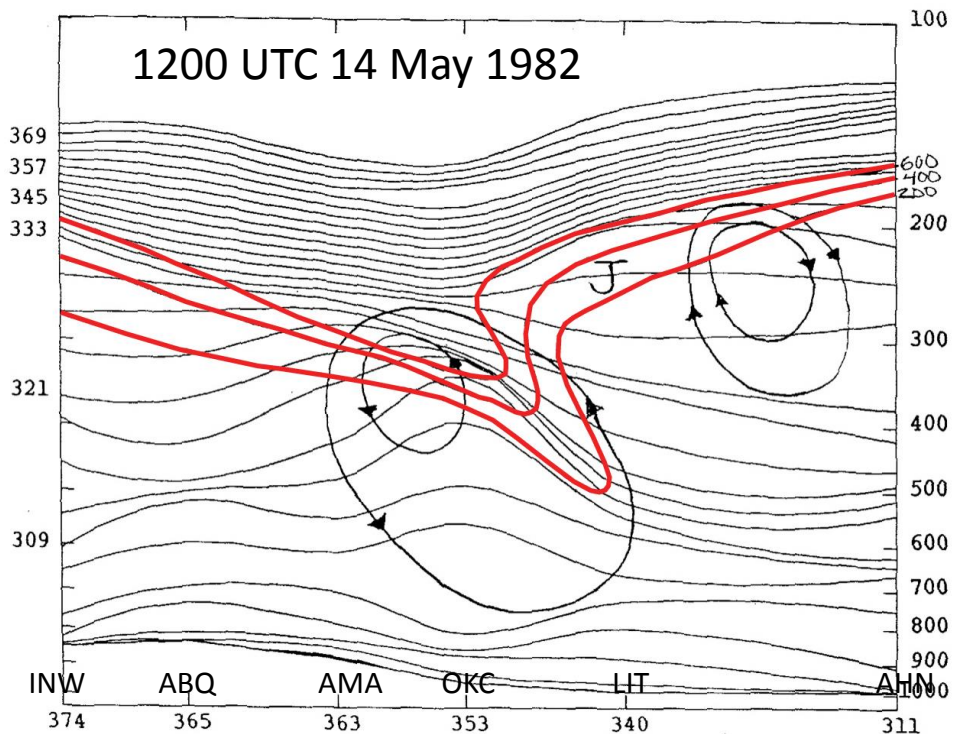
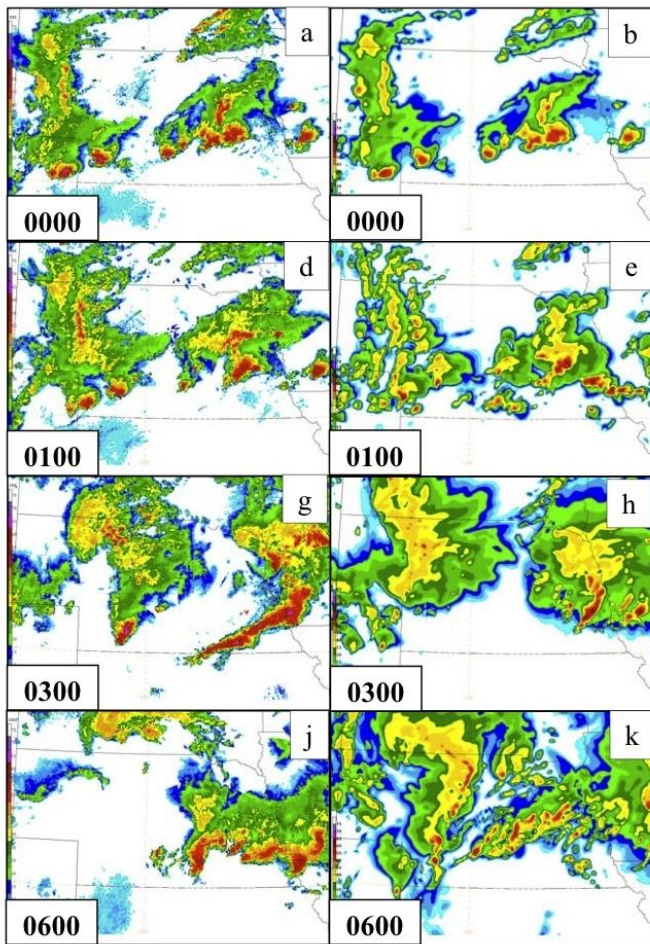


Fig. 3. Cross-section from Athens, Georgia, to Winslow, Arizona, from 1200 UTC 14 May 1982. Isentropes every 3 C. J denotes location of jet core. Dark solid lines show contours of potential vorticity ( $1 \times 10^{-3} \text{ K mb}^{-1} \text{ s}^{-2}$ ), and solid lines with arrows denote diagnosed transverse ageostrophic circulations. Location of cross-section is from A to A' in Fig. 2 (d).

# Challenges in Convective-scale NWP

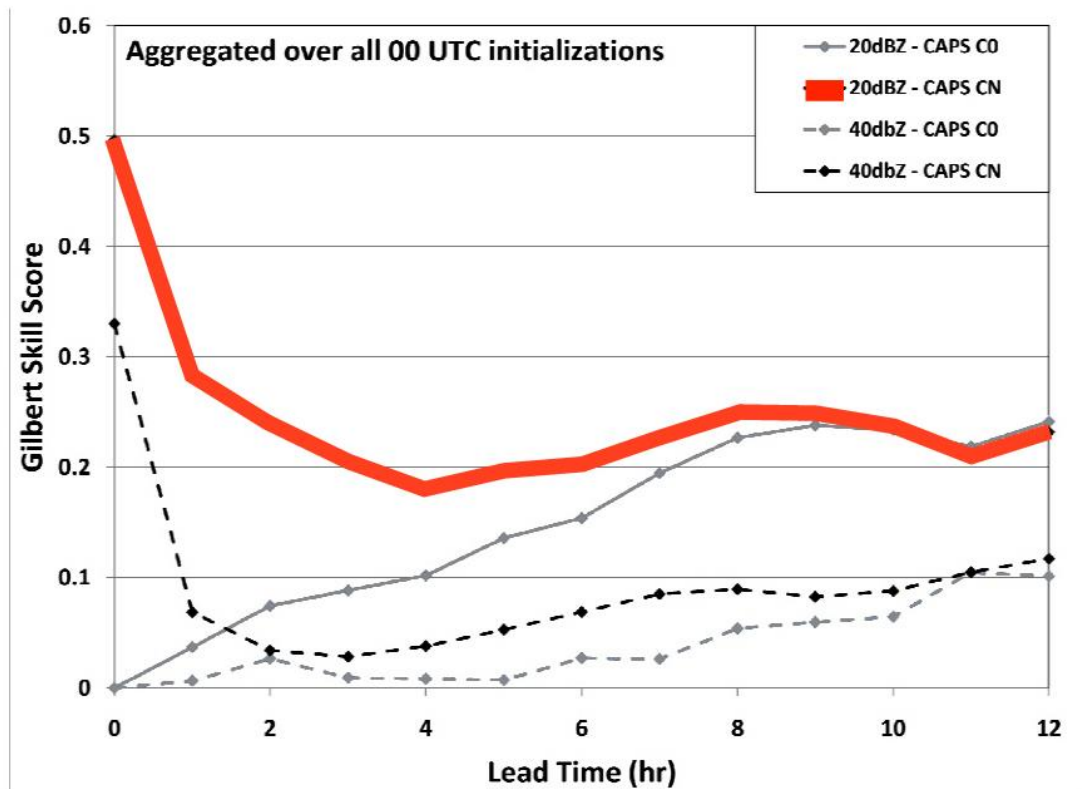
Obs Refl

Sim Refl: CN forecast



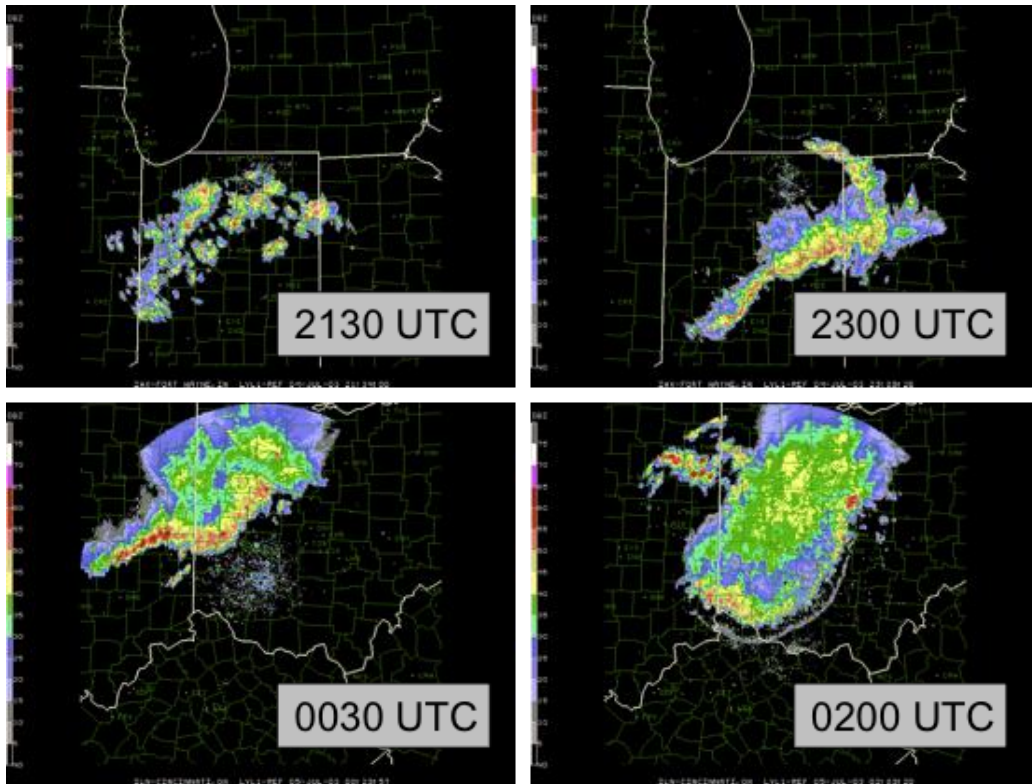
4-km forecasts initialized using radar observations yield improved short-range forecasts of convective activity (Kain et al. 2010).

Particularly helpful for looking at convective mode and evolution.



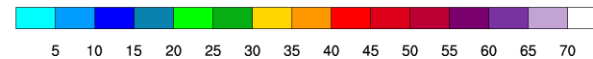
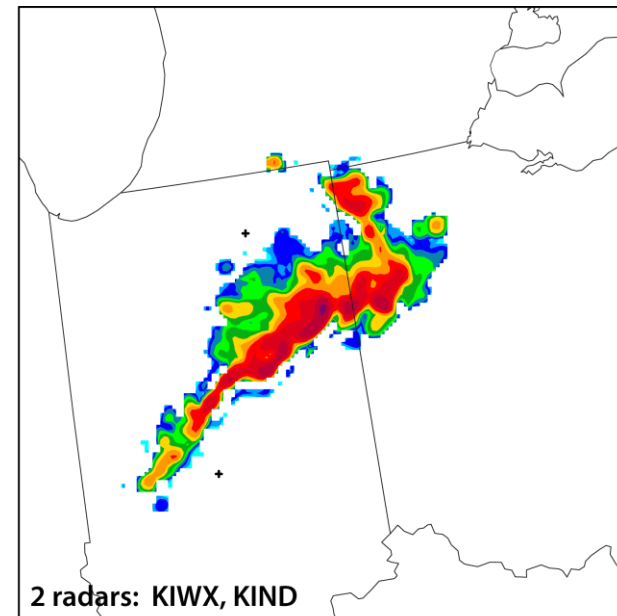
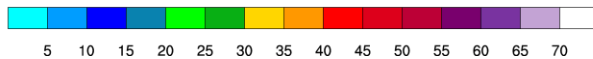
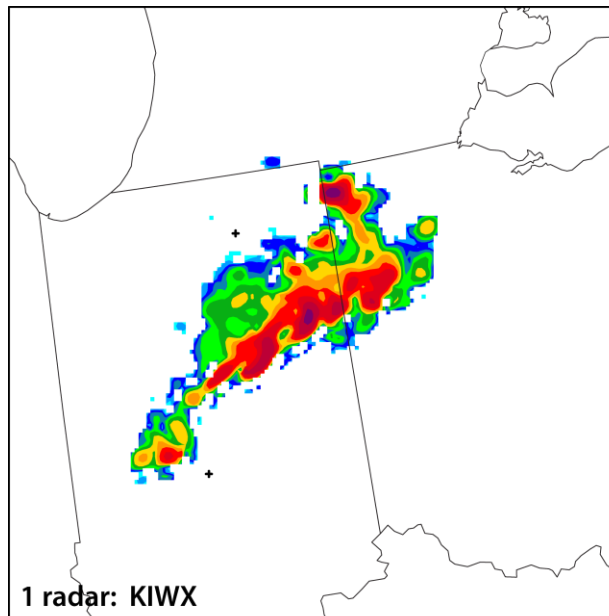
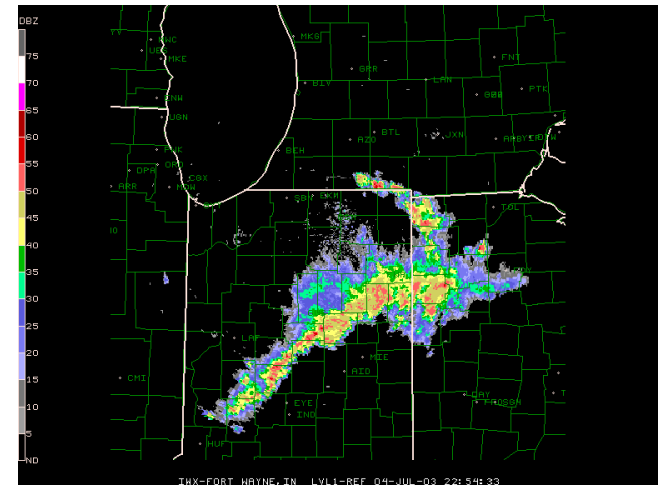
Courtesy Jack Kain

# 4-5 July 2003 MCS damaging-wind and flooding event



- Observed during BAMEX
- Produced 100+ wind reports across Indiana and Ohio
- Contributed to record flooding across north-central Indiana
- Not captured in NWP models of the day

# Ensemble Analyses valid 2300 UTC 4 July

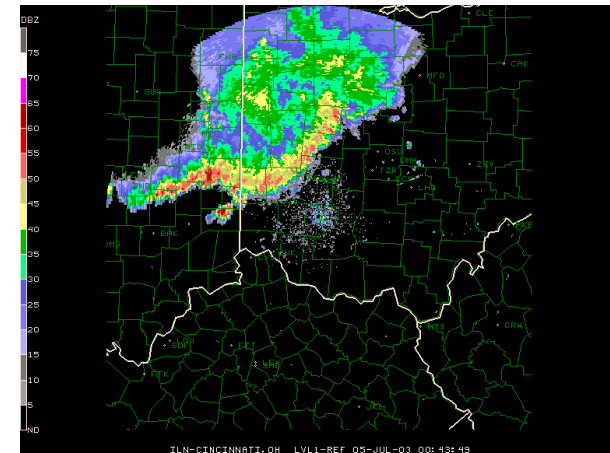


3 km dx, 51 levels, WRF/DART, 45 members, 88D observations

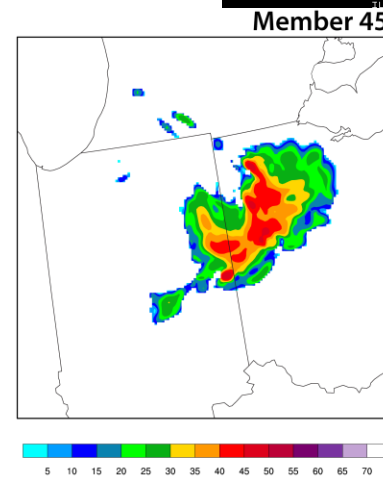
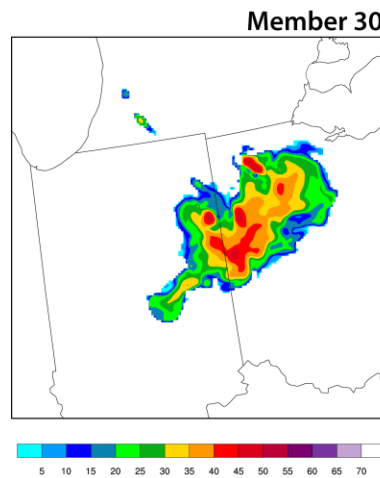
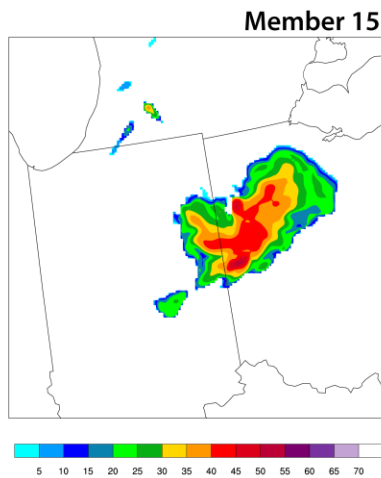
Courtesy Dustan Wheatley



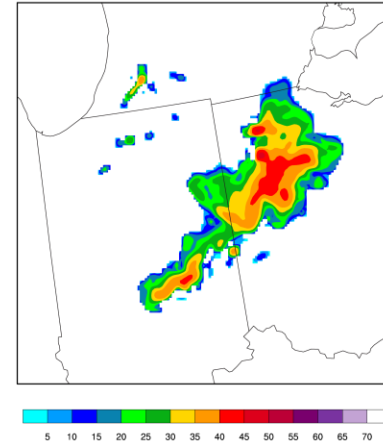
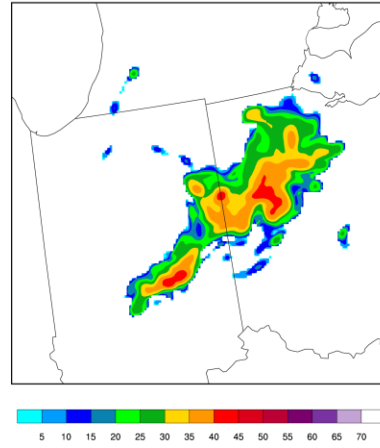
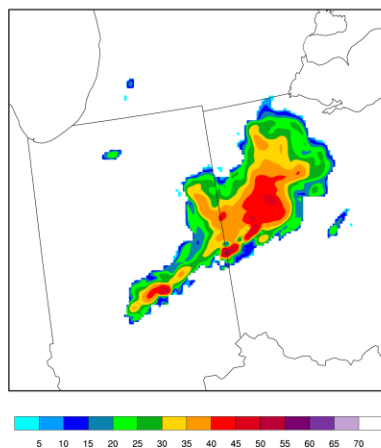
# Ensemble 90-min forecasts Valid 0030 UTC 5 July



1 radar:



2 radars:

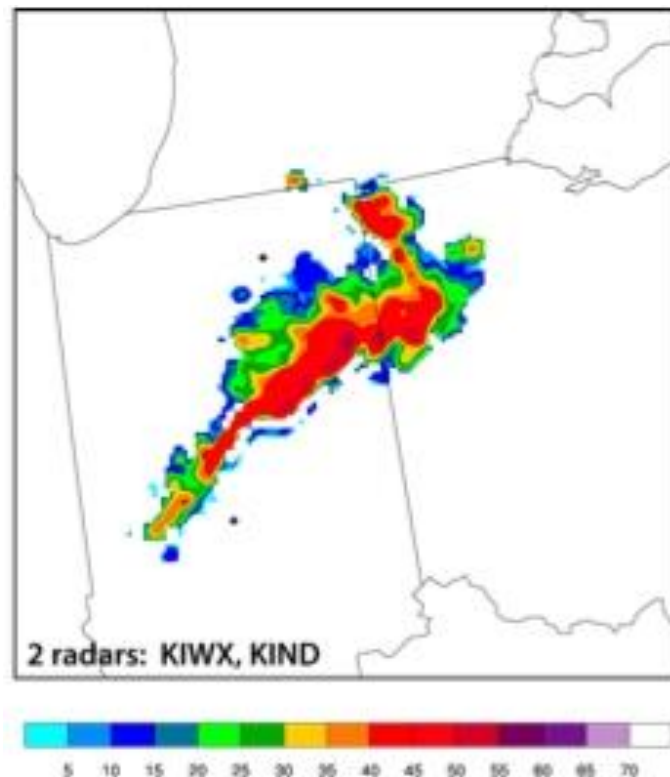


Courtesy  
Dustan Wheatley

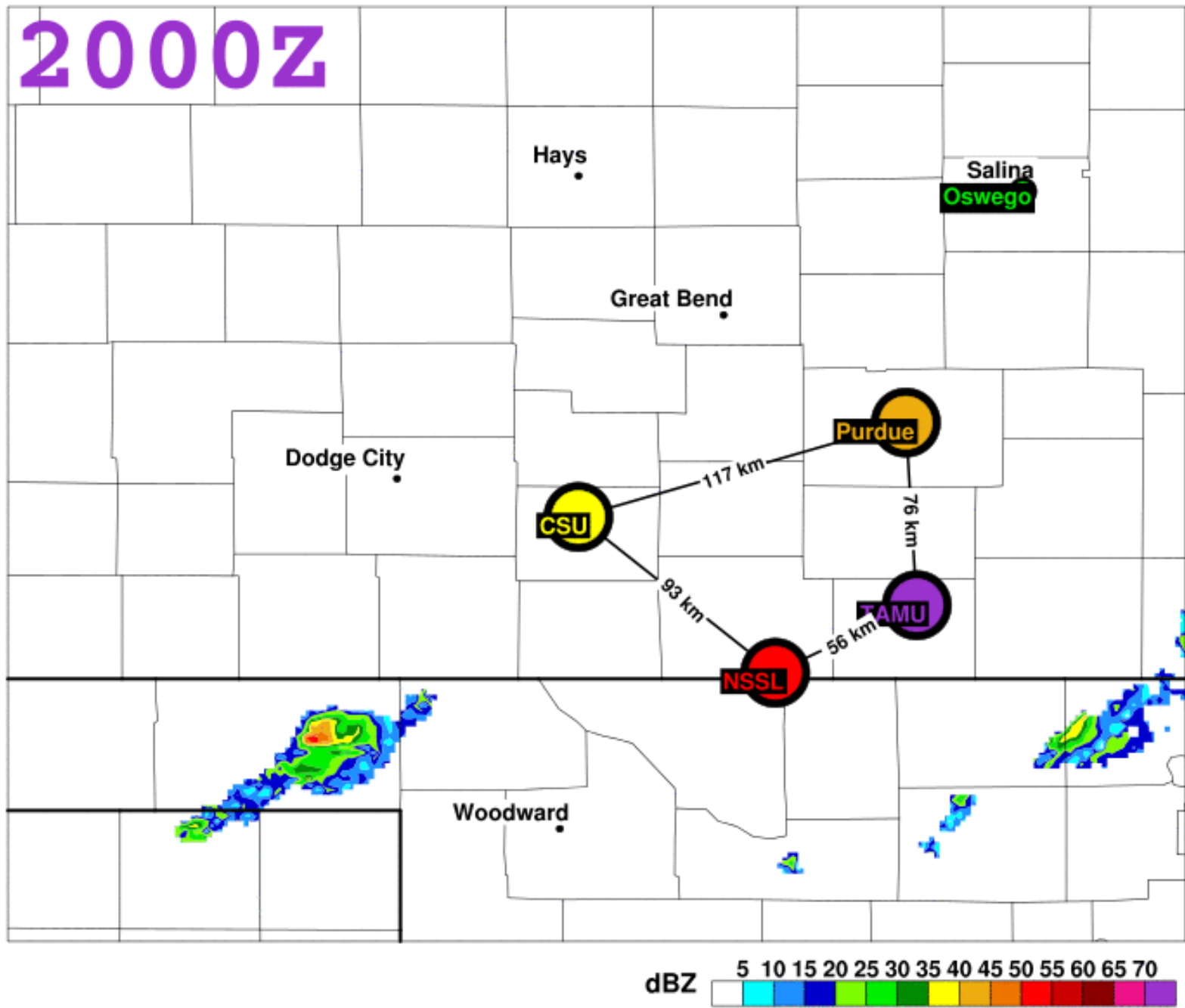
# Research Thoughts

- Current results suggest challenges in predicting organized convection in models, even though analyses appear quite good from radar and surface perspectives.
  - Nearing predictability limit
  - Model error
  - Initial condition error
    - Effects of convective feedbacks?
    - Use 60 minutes worth of data to create storms, yet often they have been in existence for hours.
    - How well do we predict their upscale influence on their nearby environment?

- Use WRF/DART to assimilate radar observations and compare ensemble soundings with MPEX soundings that surround storms/MCSs

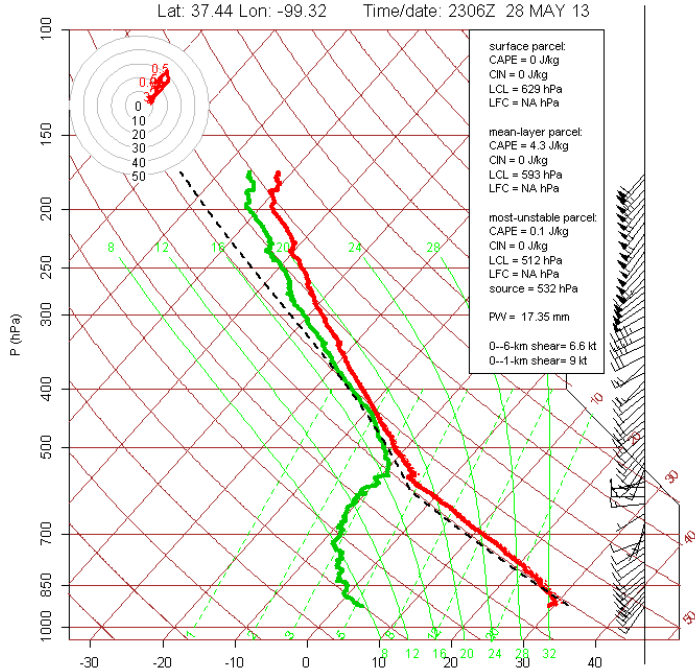


# NSSL NMQ hybrid scan reflectivity valid 2000Z 28 May 2013



### CSU sonde for MPEX

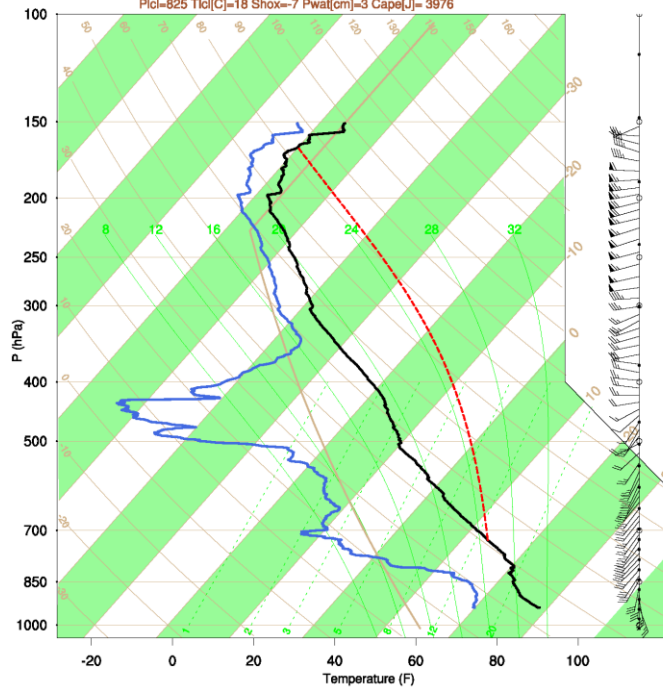
Lat: 37.44 Lon: -99.32 Time/date: 2306Z 28 MAY 13



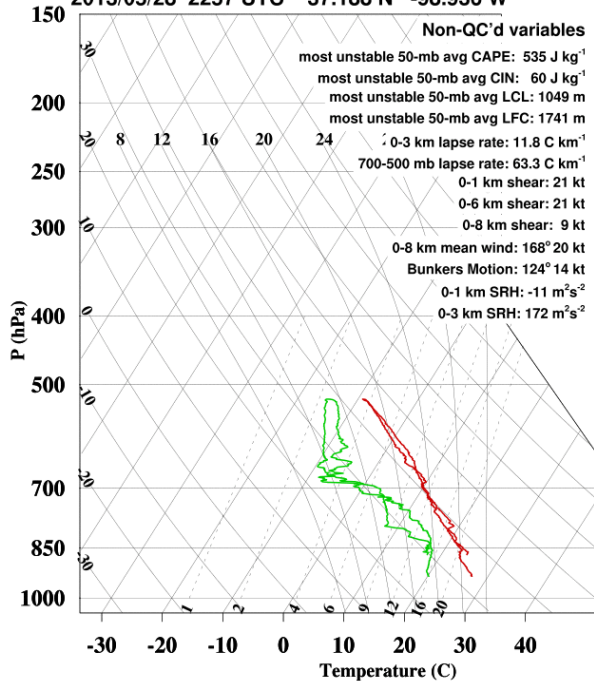
dashed black line shows mean 500-m layer parcel Temperature (C)

### Purdue MPEX 0528 2258 37.7769 -98.5397

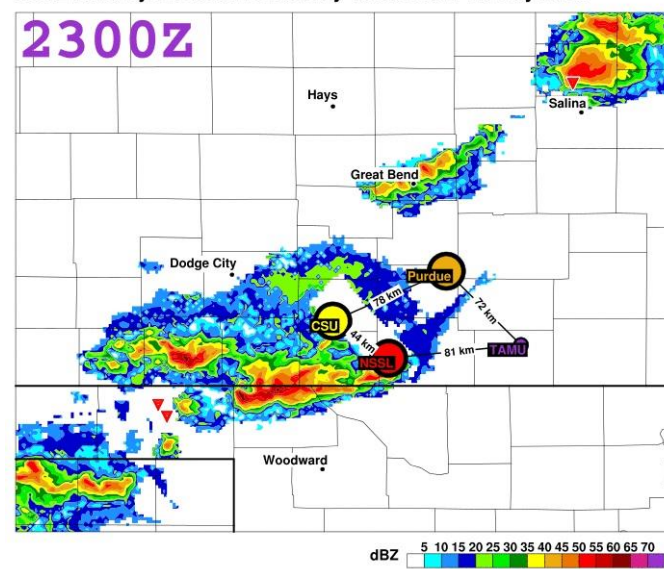
Pfcl=825 Tlo[C]=18 Shox=-7 Pwat[cm]=3 Cape[J]= 3976



2013/05/28 2257 UTC 37.188 N -98.936 W

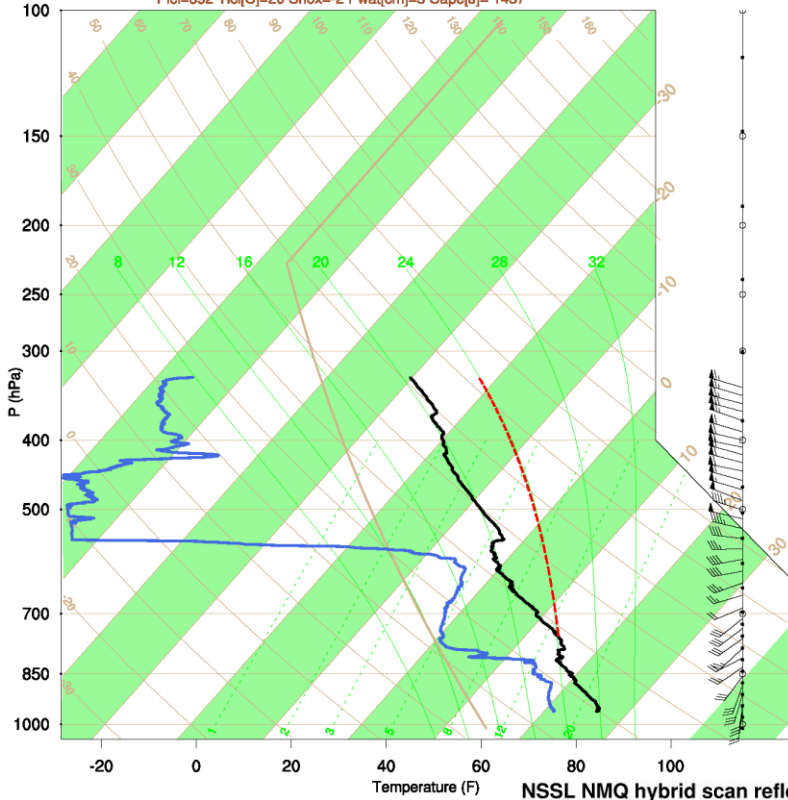


### NSSL NMQ hybrid scan reflectivity valid 2300Z 28 May 2013

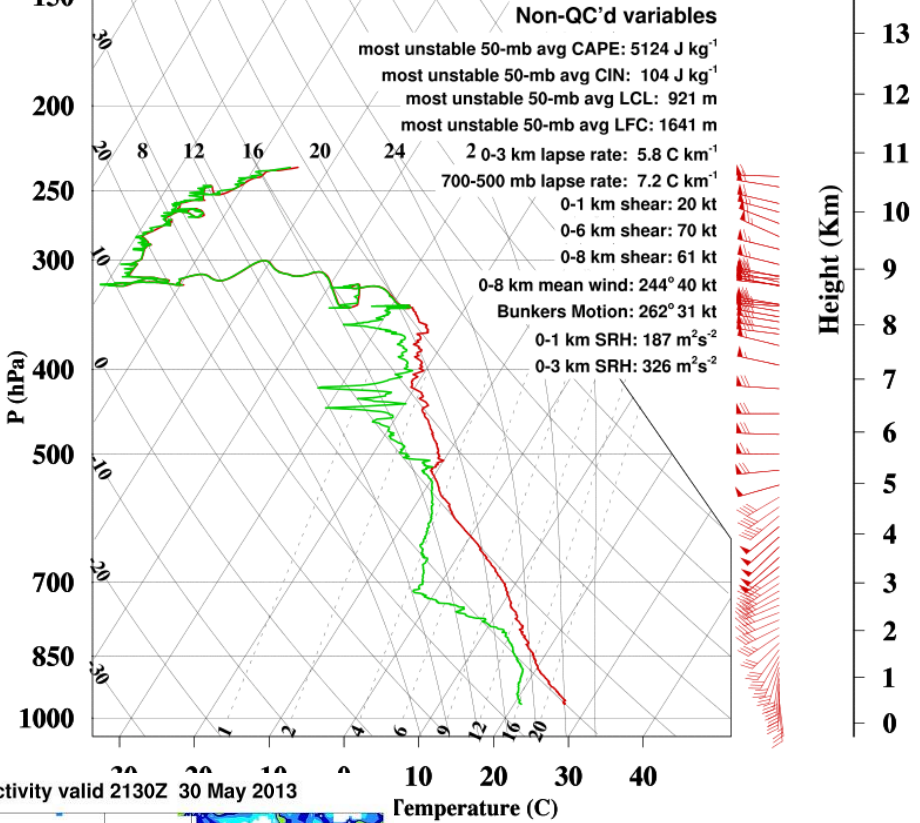


Purdue MPEX 0530 2134 35.0154 -97.5512

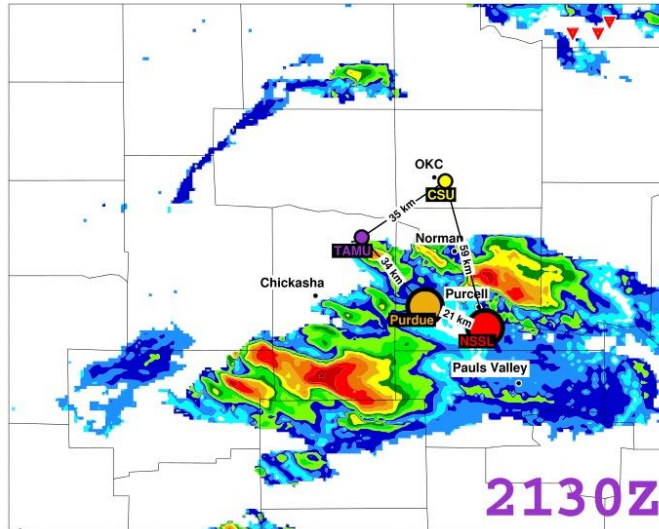
Pfci=892 Tlcl[C]=20 Shox=-2 Pwat[cm]=3 Cape[J]= 1437



2013/05/30 2130 UTC 0.000 N 0.000 W



NSSL NMQ hybrid scan reflectivity valid 2130Z 30 May 2013



dBZ 5 10 15 20 25 30 35 40 45 50 55 60 65 70

2130Z

Questions?