# **Parameterization of Langmuir turbulence**

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#### **Methods**

1) Modeling upper ocean turbulence with Large Eddy Simulation (LES) techniques

 Comparing with turbulence observations from Lagrangian floats in wind & wave – driven mixed layers, below & near hurricanes & typhoons.

 Parameterizing the effects of surface waves on upper ocean turbulence and mixed layer evolution in KPP and in second moment q<sup>2</sup>-q<sup>2</sup>l type closures (e.g. M-Y 2.5, in NCOM)



$$\begin{aligned} \frac{\mathbf{D}}{\mathbf{D}t}(q^2) &- \frac{\partial}{\partial z} \left[ q\ell S_q \frac{\partial}{\partial z}(q^2) \right] = -2\overline{u}\overline{w} \left( \frac{\partial U}{\partial z} + \frac{\partial u_{\mathrm{S}}}{\partial z} \right) - 2\overline{v}\overline{w} \left( \frac{\partial V}{\partial z} + \frac{\partial v_{\mathrm{S}}}{\partial z} \right) + 2\beta g\overline{w}\overline{\theta} - 2\frac{q^3}{B_1\ell} \\ \frac{\mathbf{D}}{\mathbf{D}t}(q^2\ell) &- \frac{\partial}{\partial z} \left[ q\ell S_1 \frac{\partial}{\partial z}(q^2\ell) \right] = E_1\ell \left( -\overline{u}\overline{w}\frac{\partial U}{\partial z} - \overline{v}\overline{w}\frac{\partial V}{\partial z} \right) + E_6\ell \left[ \left( -\overline{u}\overline{w}\frac{\partial u_{\mathrm{S}}}{\partial z} - \overline{v}\overline{w}\frac{\partial v_{\mathrm{S}}}{\partial z} \right) \right] \\ &+ E_3\ell_j(\beta g\overline{w}\overline{\theta}) - E_2\frac{q^3}{B_1} \left[ 1 + E_4\left(\frac{\ell}{\kappa\ell_w}\right)^2 \right] + E_5(2\Omega)q^2\ell \end{aligned}$$

Stokes drift **u**<sup>s</sup> drives '<u>vortex force</u>' production in TKE (q<sup>2</sup>) (McWilliams et al., 1997) and length scale (q<sup>2</sup>I) equations (Kantha & Clayson, 2004)

## Steady State LES of Langmuir turbulence





#### Scaling of bulk ML Vertical TKE:

Harcourt & D'Asaro (2008): <w<sup>2</sup> > /u<sup>\*2</sup>~ La<sub>SI</sub> -<sup>4/3</sup>

Uses a 'Surface Layer Langmuir Number  $La_{SL}$  using near-surface mean  $\langle u^s \rangle$ .

Consistent Lagrangian float <w<sup>2</sup>>observations.

Scaling entrainment power:

Energy rate to entrainment  $< w_e^3 > /u^{*3} \sim La_{SL}^{-2}$ 

LES of wide variations in wind & wave forcing

BUT not much inertial shear.

## Kantha & Clayson (2004) model:

$$\frac{Dq^2}{Dt} - \frac{\partial}{\partial x_k} \left[ ql S_q \frac{\partial q^2}{\partial x_k} \right] = 2 \left[ -\overline{u_i' u_j'} \left( \frac{\partial \overline{u_i}}{\partial x_j} + \frac{\partial u_i^S}{\partial x_i} \right) - \beta g_j \overline{u_j' \theta'} - \frac{q^2}{B_l l} \right]$$

Vortex force production added to M-Y of Kantha & Clayson (1994)

$$\begin{split} \frac{Dq^{2}l}{Dt} &- \frac{\partial}{\partial x_{k}} \left[ qlS_{l} \frac{\partial q^{2}l}{\partial x_{k}} \right] = -E_{1}l\overline{u_{l}'u_{j}'} \left( \frac{\partial \overline{u}_{i}}{\partial x_{j}} + \frac{\partial u_{i}^{S}}{\partial x_{i}} \right) - E_{3}l\boldsymbol{\beta}g_{j}\overline{u_{j}'\boldsymbol{\theta}'} - \frac{q^{3}}{B_{1}} \left( 1 + E_{2} \left( \frac{l}{\boldsymbol{\kappa}_{w}} \right)^{2} \right) \\ \overline{w'}\boldsymbol{\theta}' &= -\frac{\mathbf{K}_{H}}{S_{H}} \frac{\partial \overline{\boldsymbol{\theta}}}{\partial z}, \quad \overline{v'w'} = -\frac{\mathbf{K}_{M}}{S_{M}} \frac{\partial \overline{v}}{\partial z} \\ S_{H} &= \frac{A_{2}(1 - 6A_{1}/B_{1})}{1 - 3A_{2}G_{H}(6A_{1} + B_{2}(1 - C_{3}))} \\ S_{M} &= A_{1} \frac{(1 - 6A_{1}/B_{1} - 3C_{1}) + 9(2A_{1} + A_{2}(1 - C_{2}))S_{H}G_{H}}{1 - 9A_{1}A_{2}G_{H}} \end{split}$$

$$S_{H} = -\frac{l^{2}}{q^{2}} \mathbf{\beta}g \frac{\partial \mathbf{\theta}}{\partial z}, \quad G_{M} &= \frac{l^{2}}{q^{2}} \left| \frac{\partial \overline{u}}{\partial z} \right|^{2}, \quad S_{l} = S_{q} = S_{H}, \quad A_{l}, B_{l}, C_{l}, E_{l} = const \end{split}$$

## Problem: Stability functions do not depend on CL forcing.

Equilibrium model, dropping all transport divergence & Coriolis terms:

$$\overline{u'^{2}} = q^{2} (1 - 6A_{1} / B_{1}) / 3 - 6A_{1} l q^{-1} \overline{u'w'} \partial_{z} \overline{u}, \quad \overline{v'^{2}} = q^{2} (1 - 6A_{1} / B_{1}) / 3 - 6A_{1} l q^{-1} \overline{v'w'} \partial_{z} \overline{v}$$
$$\overline{w'^{2}} = q^{2} (1 - 6A_{1} / B_{1}) / 3 + 6A_{1} l q^{-1} \left[ \beta_{z} \overline{w'\theta'} - \overline{u'w'} \partial_{z} u^{z} - \overline{v'w'} \partial_{z} v^{z} \right]$$

$$\overline{v'w'} = -3A_1 lq^{-1} \left[ \overline{w'^2} - C_1 q^2 \right] \partial_z \overline{v} - \beta g \overline{v'\theta'} + \overline{v'^2} \partial_z v^s + \overline{u'v'} \partial_z u^s$$

Solution requires momentum flux down Stokes gradient!

$$\overline{u'w'} = -lq \Big( S_M \partial_z \overline{u} + S_M^S \partial_z u^S \Big), \qquad \overline{v'w'} = -lq \Big( S_M \partial_z \overline{v} + S_M^S \partial_z v^S \Big), \qquad \overline{w' \theta'} = -lq S_H \partial_z \overline{\theta}$$



Solution: 
$$\overline{u'w'} = -lq(S_M\partial_z\overline{u} + S_M^S\partial_zu^S), \quad \overline{v'w'} = -lq(S_M\partial_z\overline{v} + S_M^S\partial_zv^S), \quad \overline{w'\theta'} = -lqS_H\partial_z\overline{\theta}$$

$$\begin{split} S_{H} &= \frac{D_{H0}D_{M1}D_{S1} + D_{H2}D_{M0}D_{S1} + (D_{H3}D_{M1} + D_{M3}D_{H2})D_{S0}}{D_{H1}D_{M1}D_{S1} - D_{H2}D_{M2}D_{S1} - (D_{H3}D_{M1} + D_{M3}D_{H2})D_{S2}} \\ S_{M}^{S} &= (D_{S0} + D_{S2}S_{H})/D_{S1}, \quad S_{M} = (D_{M0} + D_{M2}S_{H} + D_{M3}S_{M}^{S})/D_{M1} \\ \hline \textbf{Stability functions: Much more complicated!!} \\ D_{S0} &= A_{1}(1 - 6A_{1}/B_{1} - 3C_{1}^{S}), \quad D_{S1} = 1 - 9A_{1}(A_{2}G_{H} + A_{1}G_{V}), \quad D_{S2} = 9A_{1}A_{2}C_{2}^{S}G_{H} \\ D_{M0} &= A_{1}(1 - 6A_{1}/B_{1} - 3C_{1}), \quad D_{M1} = 1 - 9A_{1}(A_{2}G_{H} + 4A_{1}G_{V}), \\ D_{M2} &= 9A_{1}(2A_{1} + A_{2}(1 - C_{2}))G_{H}, \quad D_{M3} = 27A_{1}^{2}G_{S} \\ D_{H0} &= A_{2}(1 - 6A_{1}/B_{1}), \quad D_{H1} = 1 - 3A_{2} \Big[ (6A_{1} + B_{2}(1 - C_{3}))G_{H} + 3A_{2}(1 - C_{2})G_{V} - 3A_{2}C_{2}^{S}G_{S}) \Big], \\ D_{H2} &= 9A_{2}G_{V}(2A_{1} + A_{2}), \quad D_{H3} = 9A_{2}G_{S}(2A_{1} + A_{2}) \end{split}$$

<u>New n.d. forcing functions</u>  $G_{V} = \frac{l^{2}}{q^{2}} \frac{\partial \vec{u}}{\partial z} \cdot \frac{\partial \vec{u}^{S}}{\partial z}, \qquad G_{S} = \frac{l^{2}}{q^{2}} \left| \frac{\partial \vec{u}^{S}}{\partial z} \right|^{2}$ 

$$G_{H} = -\frac{l^{2}}{q^{2}} \beta g \frac{\partial \theta}{\partial z}, \qquad G_{M} = \frac{l^{2}}{q^{2}} \left| \frac{\partial \vec{u}}{\partial z} \right|^{2},$$





### Entrainment power scaling: LES vs. closure:

Corresponding scaling of work rate with La<sub>SL</sub> --Closure model entrainment higher overall.

Similar response to angle between wind & waves -- skewness from Coriolis stronger.

# LES vs 2<sup>nd</sup> moment M-Y closure (Kantha & Clayson,1994), i.e. NCOM Wave forcing from WW3 model (by II-Ju Moon)

<u>LES</u>

# Kantha & Clayson (1994)









Eulerian |U| (m/s)





# LES vs recent new 2<sup>nd</sup> moment closure

<u>LES</u>







# New 2<sup>nd</sup> moment closure v.B4







#### Entrainment physics across cyclone:

Final wake formation dominated by mixing driven by inertial shear.

Intermediate state, timing of wake below cyclone shifted by Stokes drift.

Below maximum winds, cooling is 50% more: reduced 1 C vs 2/3 C .

Mixing enhanced ahead of cyclone, damped to rear.

Combined effects cancel and are also much less than shear-driven mixing.



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Wave-modified turbulence closure and LES (P. Sullivan too) need wind & wave forcing from Typhoons to simulate ITOP turbulence observations. Best: archived wave model data (WW3, SWAN) verified against remote & in-situ observations.

Seeking collaboration to put into modified closure in wave-ocean coupled model.

Lots of SST, Ocean Color imagery prepared for field cruise operations in small geographic domains still available at itop.org/data/harcourt . Includes some daily AVISO maps -- AVISO archives weekly maps. Other original satellite data archived at NODC & others, so .... should we archive any of this imagery?

