



Determining Best Method for Estimating Observed Level of Maximum Convective Detrainment based on Radar Reflectivity

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1. Introduction and Background

Convective mass transport is the transport of mass from near the surface up to the upper troposphere and lower stratosphere (UTLS) by a deep convective updraft. This transport can alter the chemical makeup and water vapor balance of the UTLS, which can affect cloud formation and the radiative properties of the atmosphere. It is therefore important to understand the exact altitudes at which mass is detrained from convection. These detrainment altitudes are also important for constraining deep convective transport in chemical transport models and climate models. According to parcel theory, the maximum detrainment in a storm should be at the level of neutral buoyancy (LNB) since that is the level where an air parcel's upward acceleration ceases. However, the LNB is an idealized variable that does not account for entrainment, so in practice the maximum detrainment would occur below the LNB (Mullendore et al., 2009; Takahashi and Luo, 2012; Mullendore et al., (2013). Mullendore et al. (2009) used radar reflectivity as a direct observer of vertical transport for a tropical squall line. This study builds upon this previous work by testing three methods with a variety of storm types from the Severe Thunderstorm Electrification and Precipitation Study (STEPS) and from the Colorado State CHILL and Pawnee Doppler radar network. The purpose of this study is to develop a method for estimating the level of maximum detrainment within convection using data from individual radars. Such an approach would maximize the spatial and temporal coverage of convective mass-detrainment estimates.

2. Three Methods

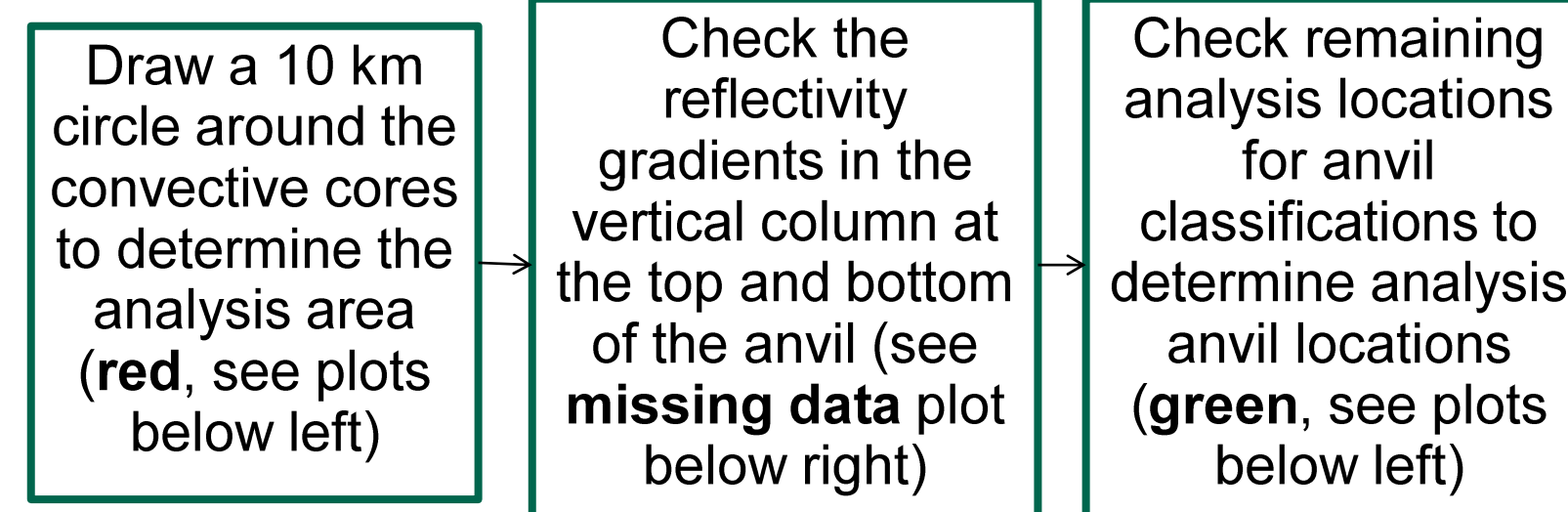
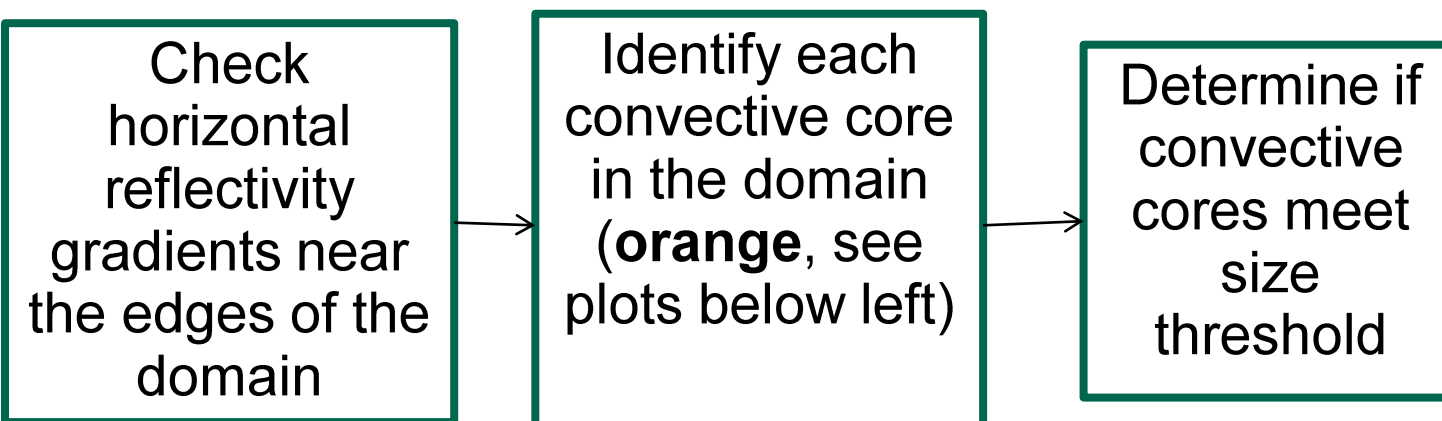
Method Name	Convective	Anvil
Steiner	Steiner et al. (1995) radar reflectivity based	simple anvil
Dual-Doppler (DDA)	6 m/s average updraft	simple anvil
Convective Stratiform Anvil (CSA)	Feng et al. (2011) radar reflectivity based	Feng et al. (2011) radar reflectivity based

- Convective and anvil classifications needed because only hydrometeors in a storm's convectively generated anvil were found to be a proxy in Mullendore et al. (2009).
- Anvil locations were determined by "simple anvil" from Mullendore et al. (2009) and an echo layer identification method from Feng et al. (2011).
- Methods were chosen because: Steiner is similar to previous methods, DDA makes use of vertical velocities to find updrafts, and CSA has a more robust anvil identification technique.
- Results of all three methods were compared against vertical mass divergence based LMDs which were considered "truth".

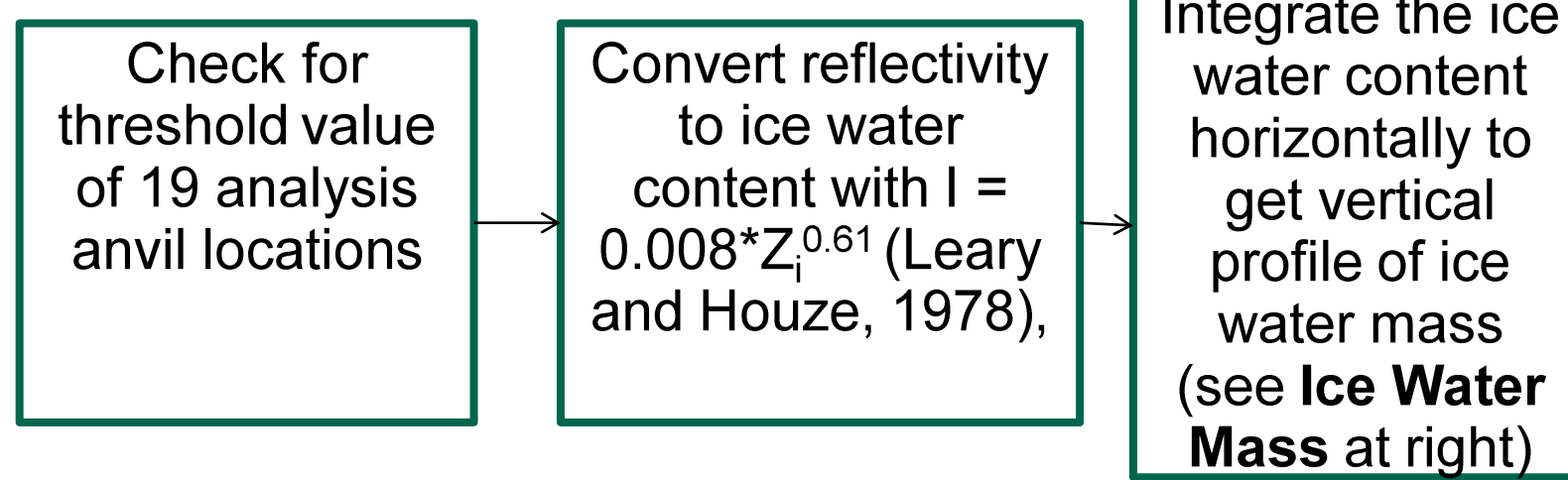
3. Process

Entire Domain

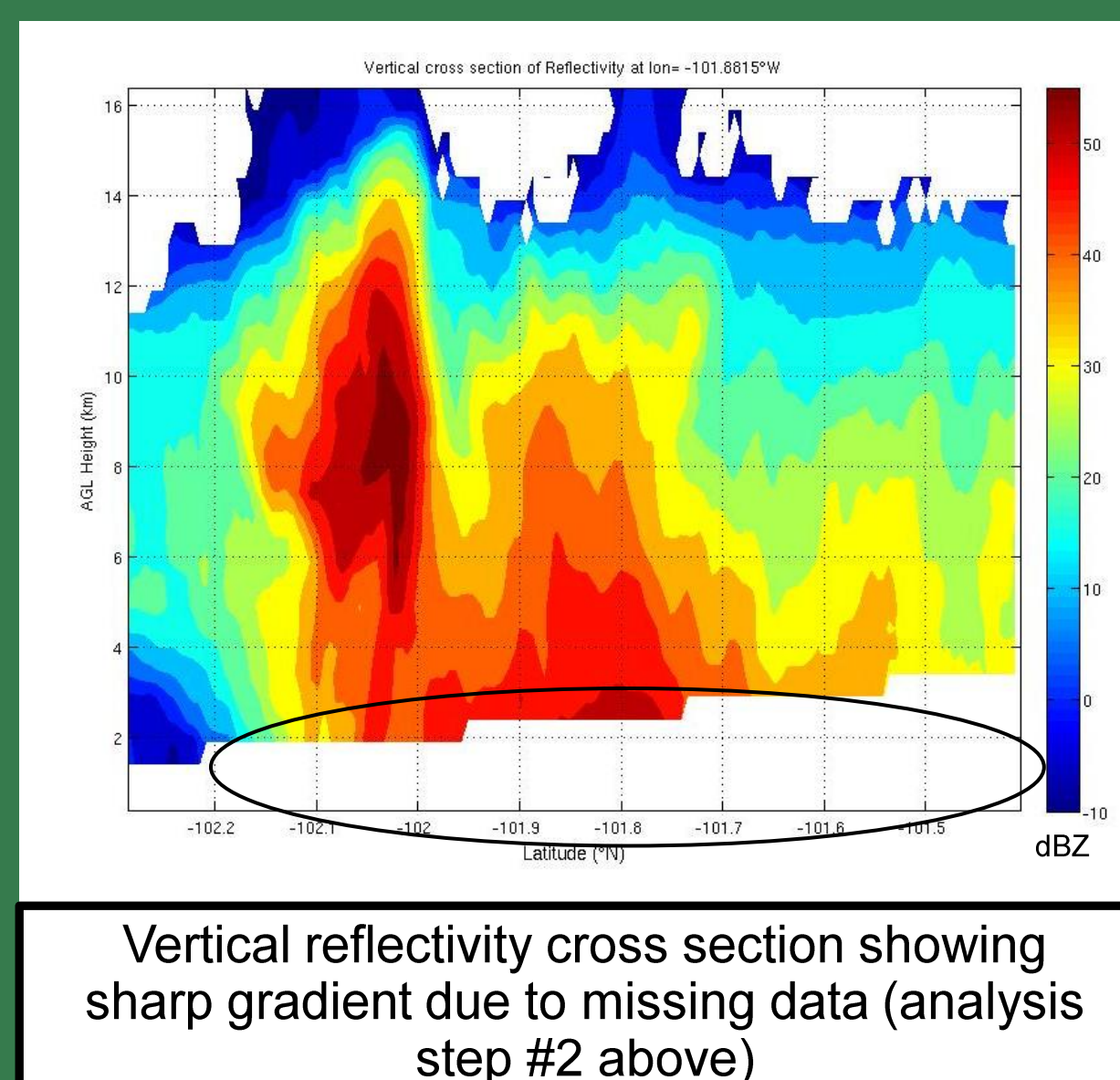
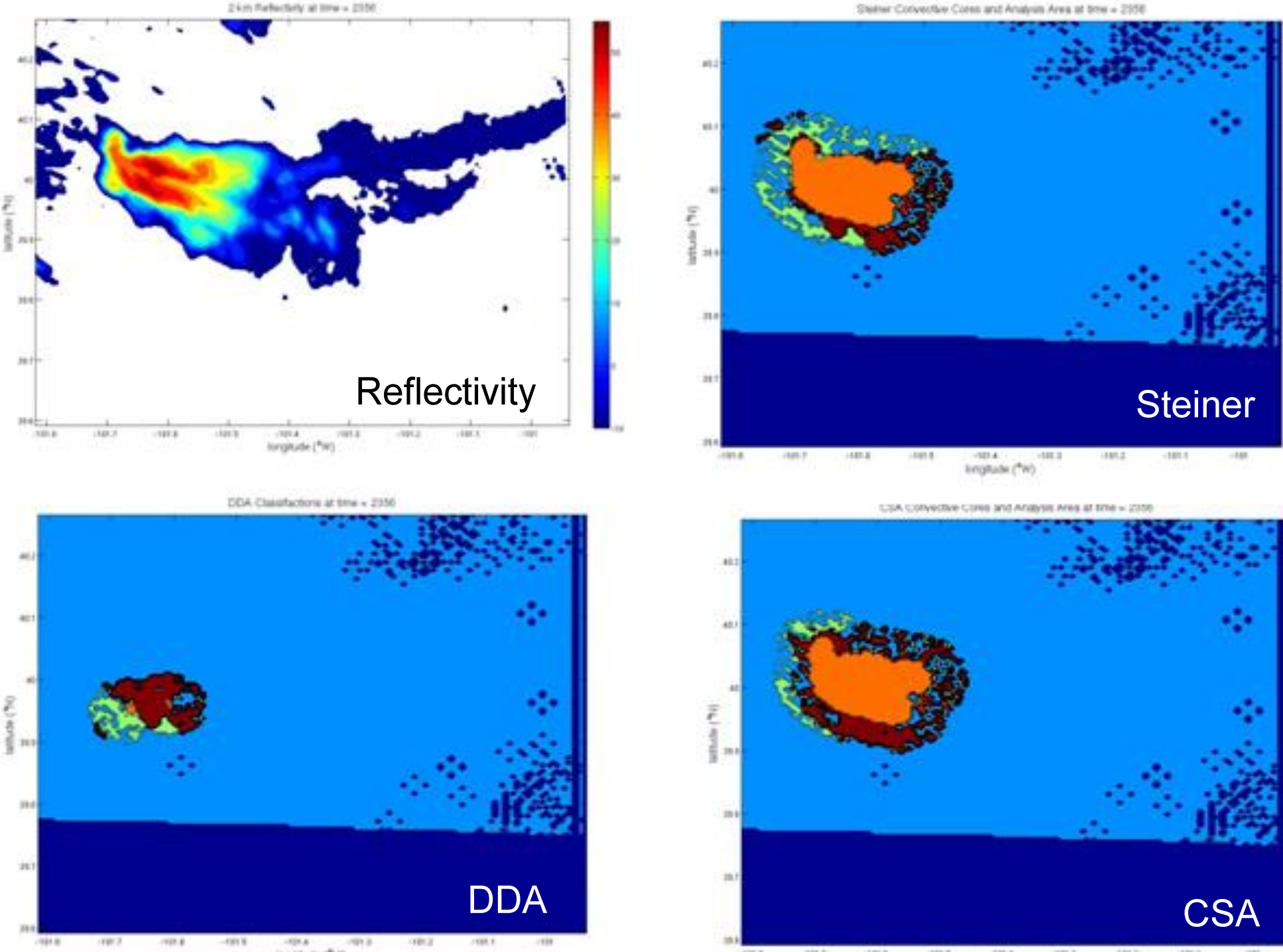
Analysis Area



Ice Mass Calculation



Convective and Anvil Locations for all Three Methods (colors defined above)



4. Method Comparison Results

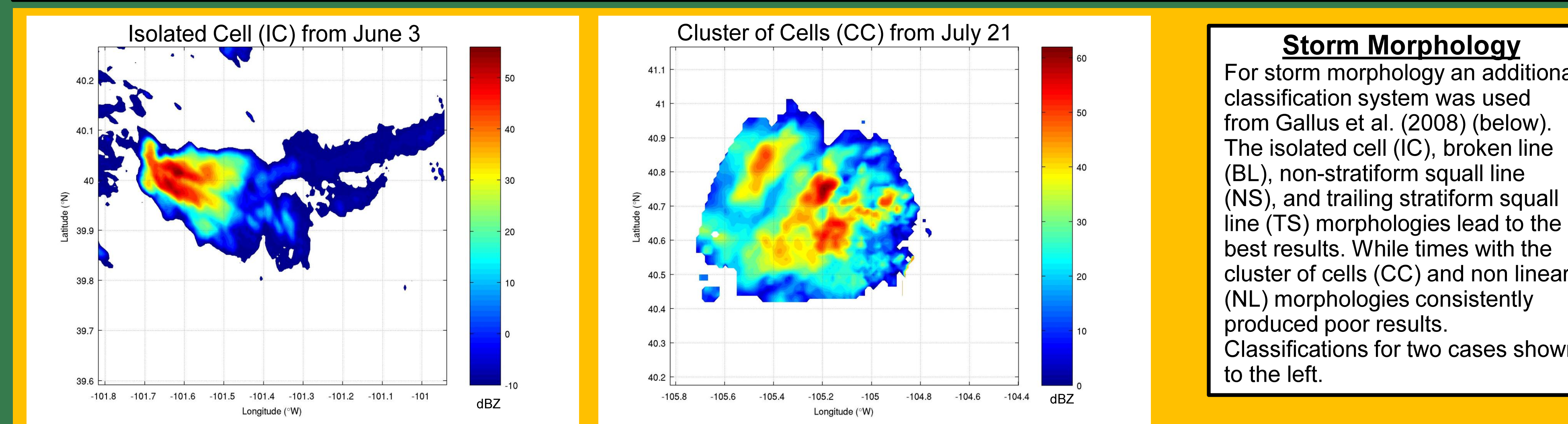
Case	Div-DDA (km)	Div-Steiner (km)	Div-CSA (km)	Date	Morphology	Rep Time (UTC)	Div-DDA (km)	Div-Steiner (km)	Div-CSA (km)
June 3 Single Cell	0.98 (43%)	0.82 (43%)	0.51 (100%)	June 3	IC	2356	0.5	0.5	0
June 11 Convective Line	0.58 (43%)	1.33 (43%)	2.00 (29%)	June 11	TS	0025	-	0.2	-0.3
June 19 Convective Line	- (0%)	- (0%)	0.25 (29%)	June 19	NL	0148	-	-	-
June 22 Convective Line	2.09 (100%)	2.05 (100%)	1.31 (91%)	June 22	TS	0054	5.0	2.5	4
June 23 Multicell	3.25 (14%)	2.00 (14%)	0.93 (100%)	June 23	BL	2232	-	-	0.5
June 29 Supercell	2.61 (100%)	2.75 (100%)	1.25 (100%)	June 29	IC	0010	3.0	3.5	-1.0*
July 2 Supercell	3.50 (21%)	3.50 (21%)	1.64 (100%)	July 2	IC	0135	-	-	1.5
July 15 Convective Line	2.75 (50%)	2.64 (50%)	1.40 (36%)	July 15	NS, CC	0028	3.0	2.5	2.5
July 21 Multicell	4.75 (14%)	4.75 (14%)	3.90 (36%)	July 21	CC	2126	4.5	4.5	4.5

Difference between the three proxy methods' LMDs and the vertical mass divergence LMD
CSA method outperforms DDA and Steiner methods (yellow boxes)

- For the table above the difference between the average vertical mass divergence and the LMD at each time was taken. This was done for all the times that passed the tests and then for each case an average was determined.
- In parenthesis is the percentage of times for each case that passed all of the tests.

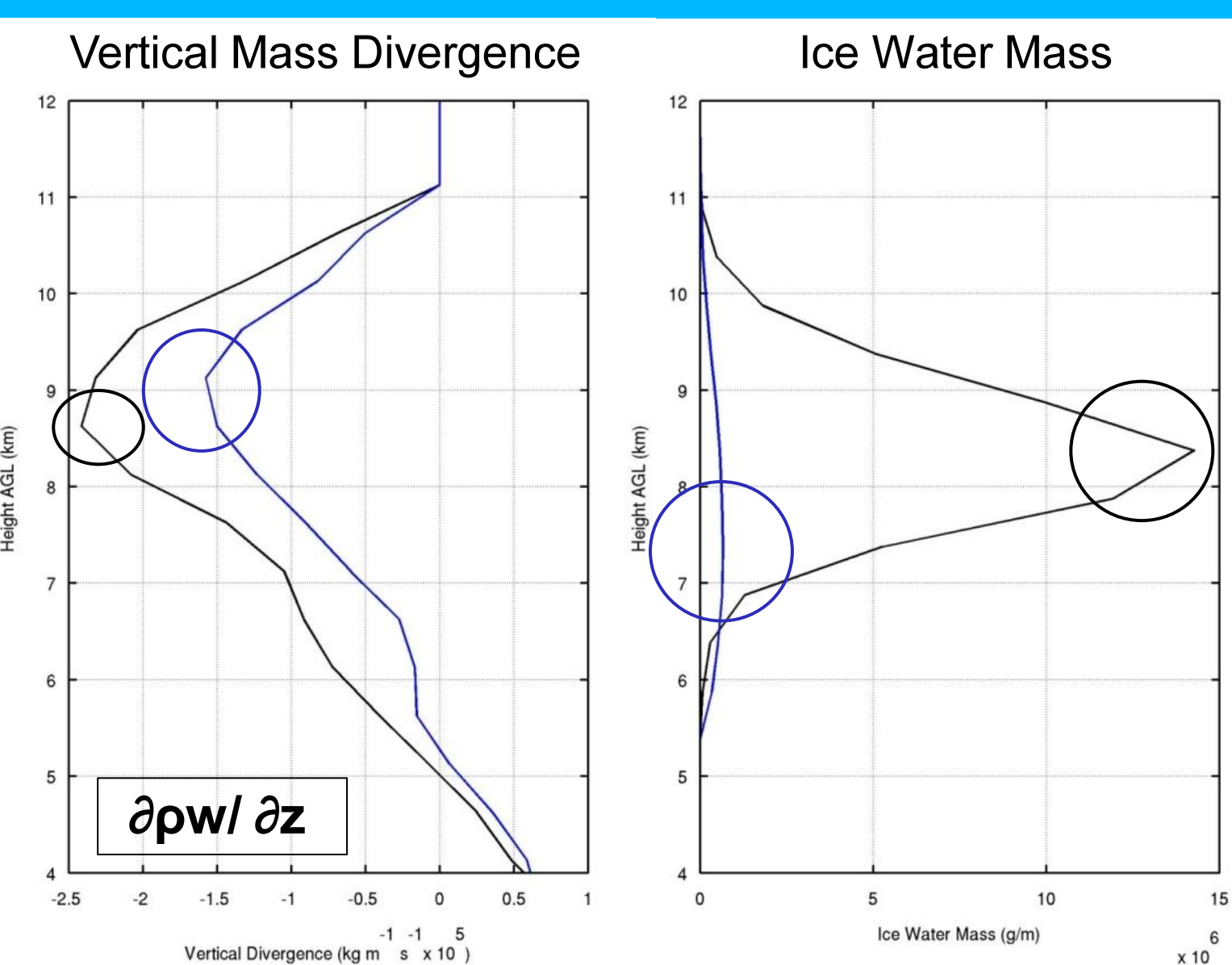
- For the table above a different approach was taken. Based on later analysis of when the methods performed best a most representative time, late mature time, was determined for each case (see storm maturity identification below).
- At this most representative time the difference between the divergence LMD and LMD from the proxy methods was taken.
- Green morphologies were found to lead to satisfactory results while red morphologies were found to lead to unsatisfactory results.
- The time with an * focuses on the supercell and ignores the developing cell.

5. Improving Application of Proxy Method



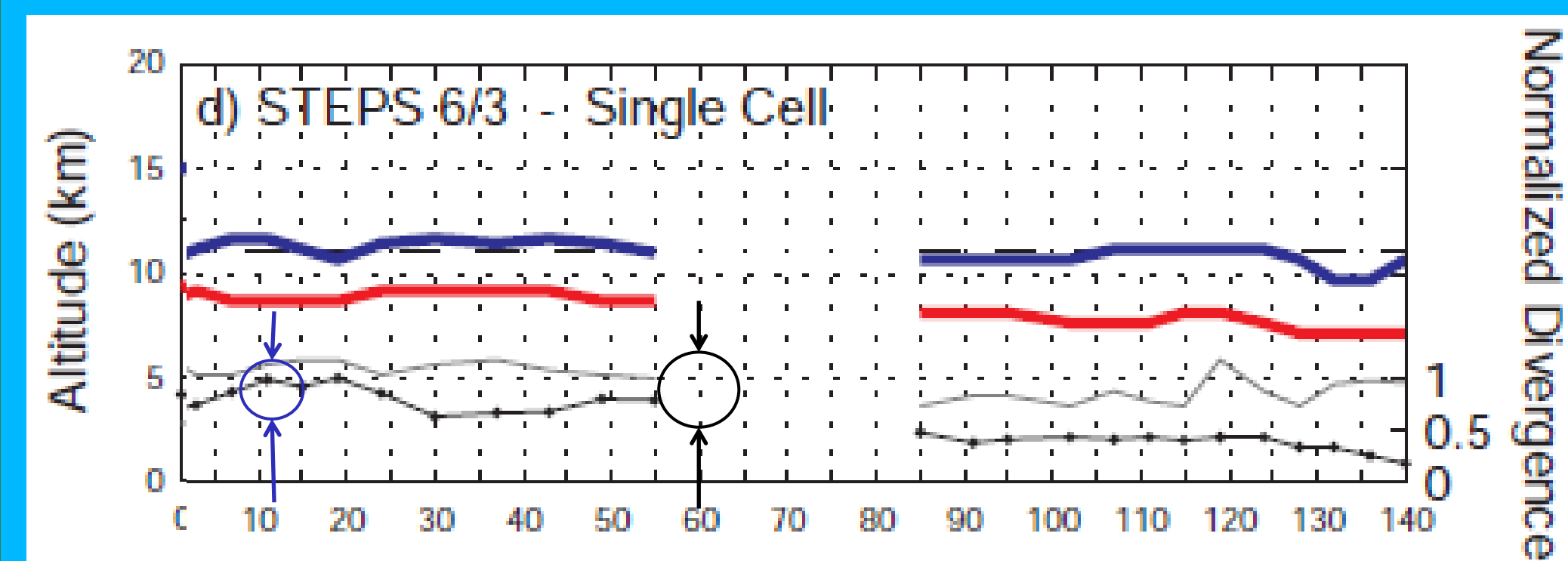
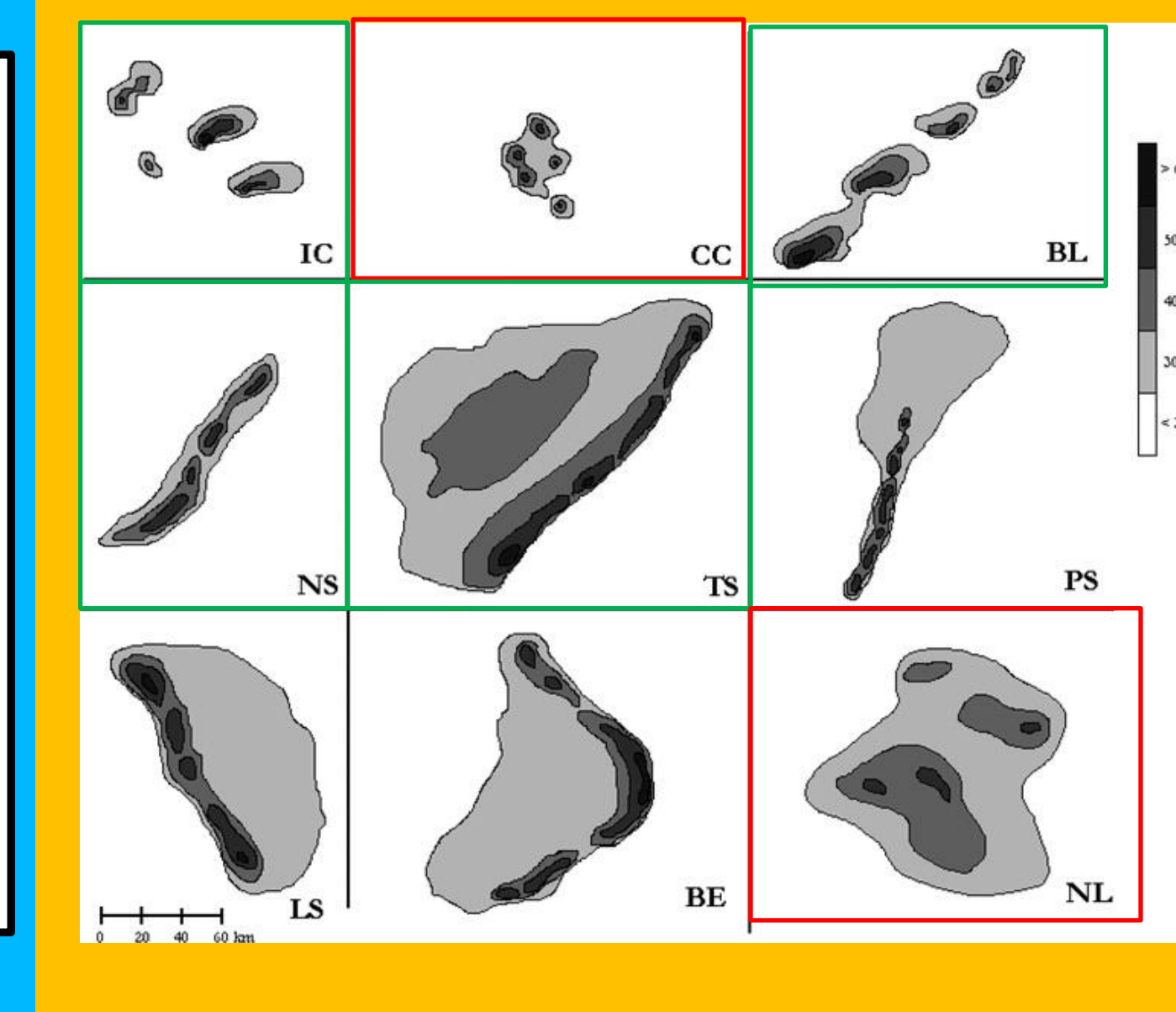
Storm Morphology

For storm morphology an additional classification system was used from Gallus et al. (2008) (below). The isolated cell (IC), broken line (BL), non-stratiform squall line (NS), and trailing stratiform squall line (TS) morphologies lead to the best results. While times with the cluster of cells (CC) and non linear (NL) morphologies consistently produced poor results. Classifications for two cases shown to the left.



Storm Maturity

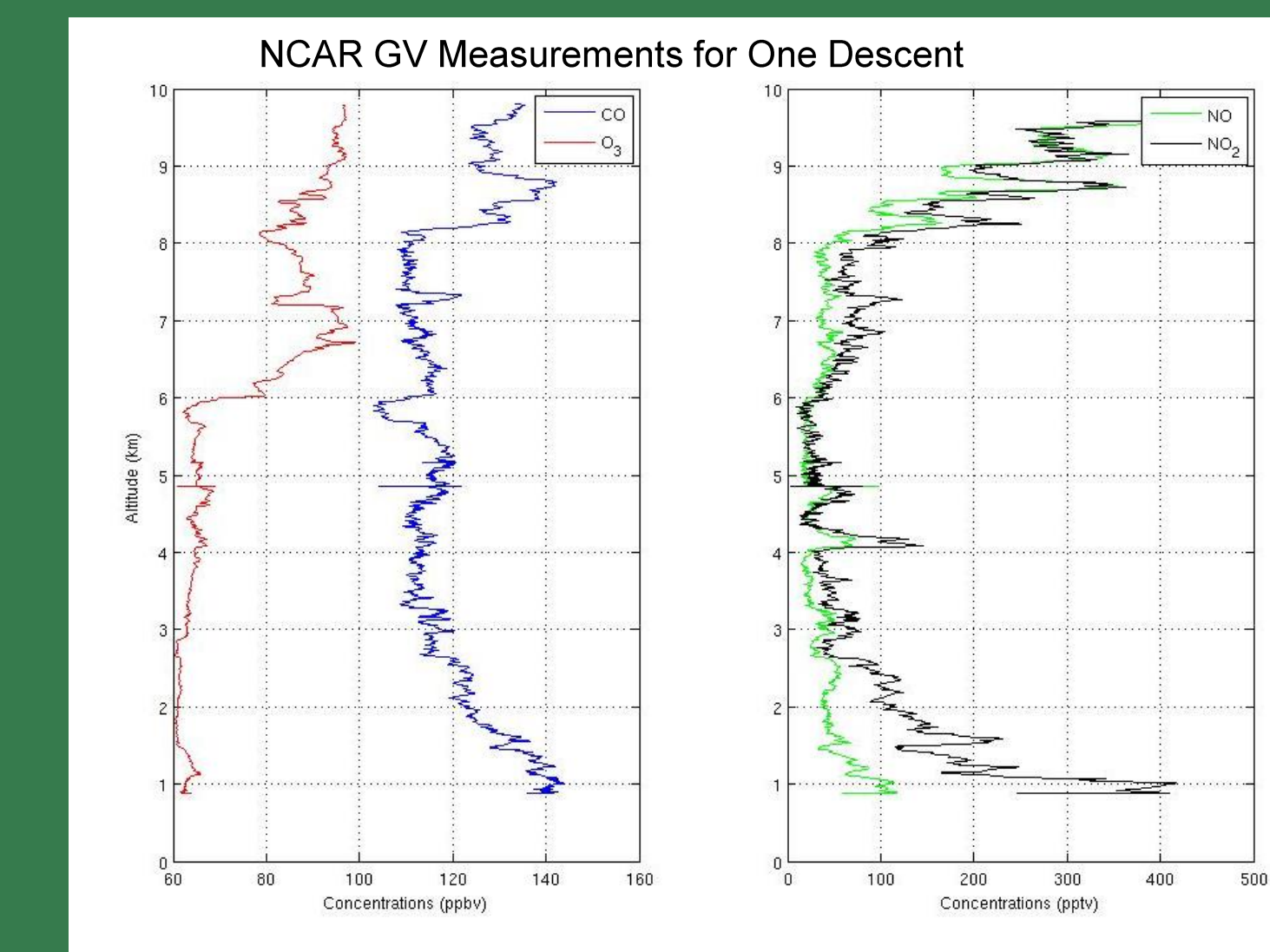
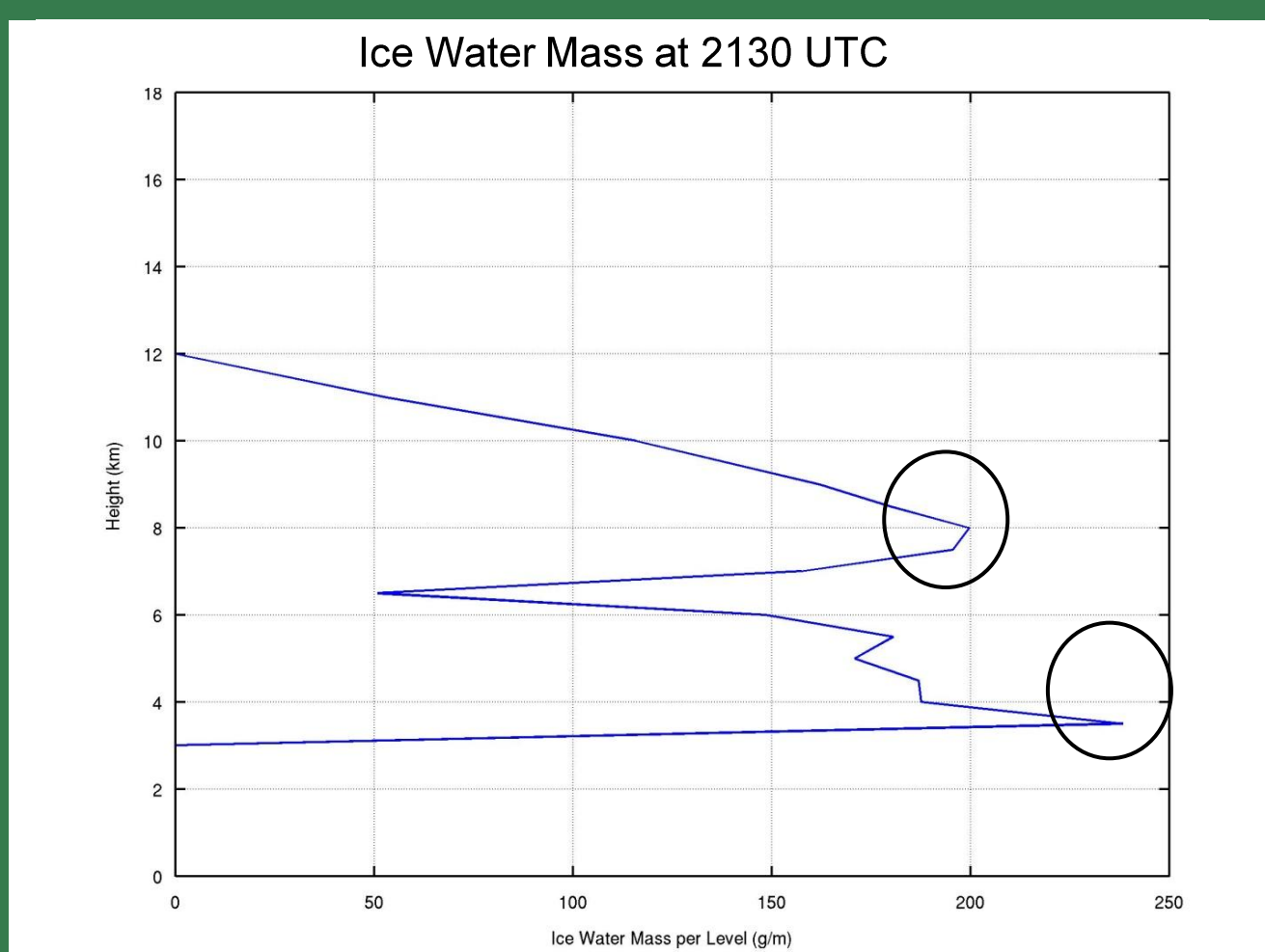
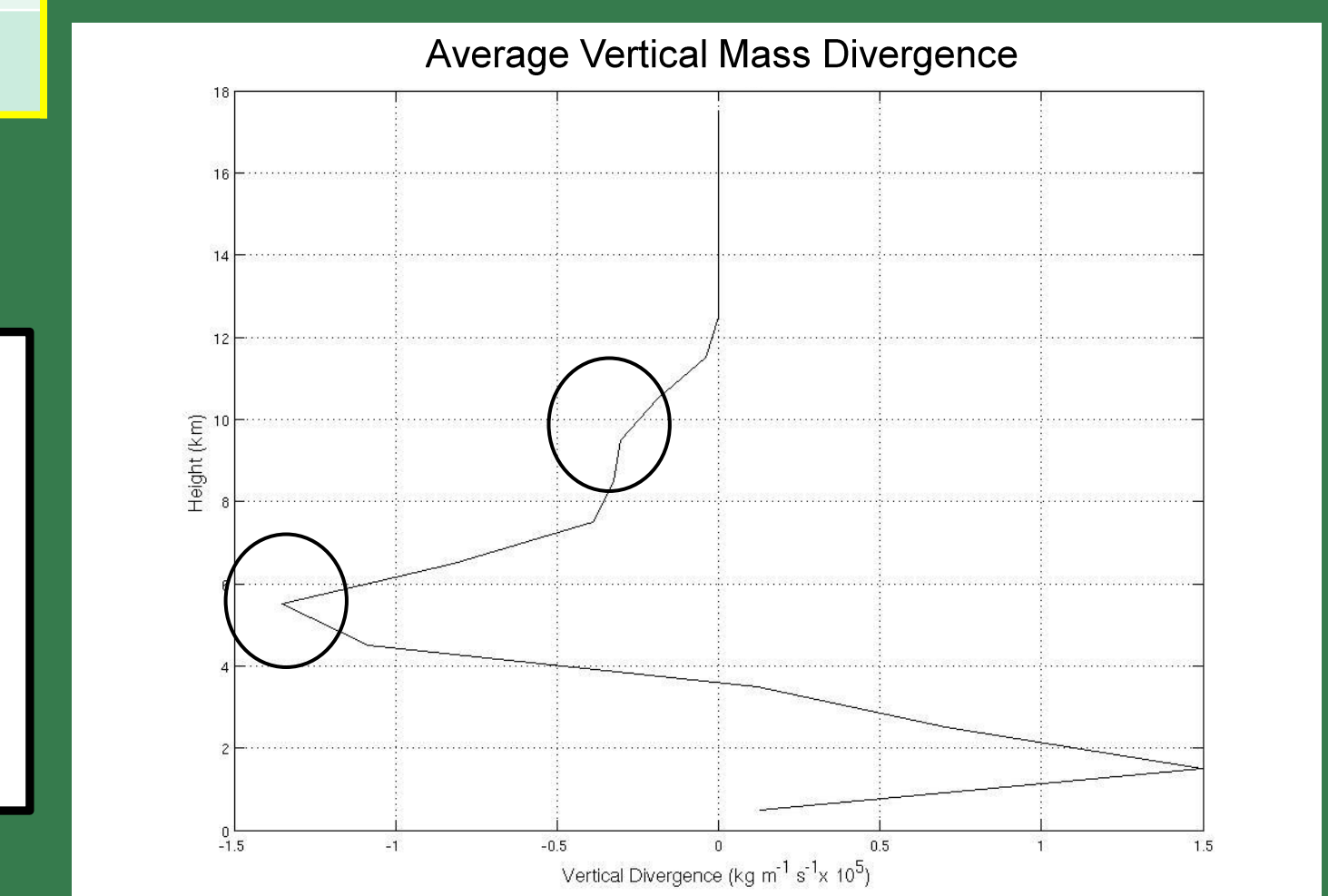
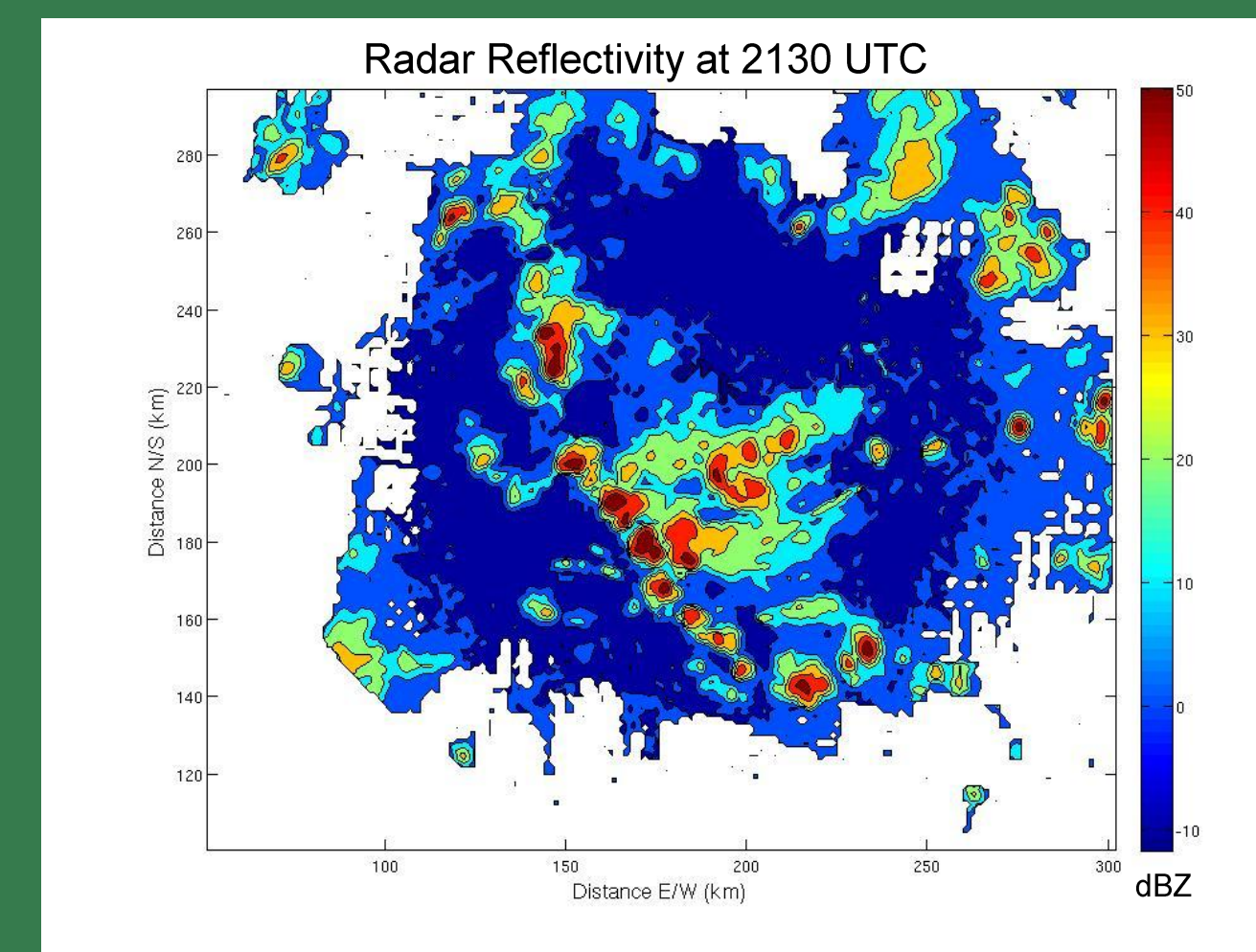
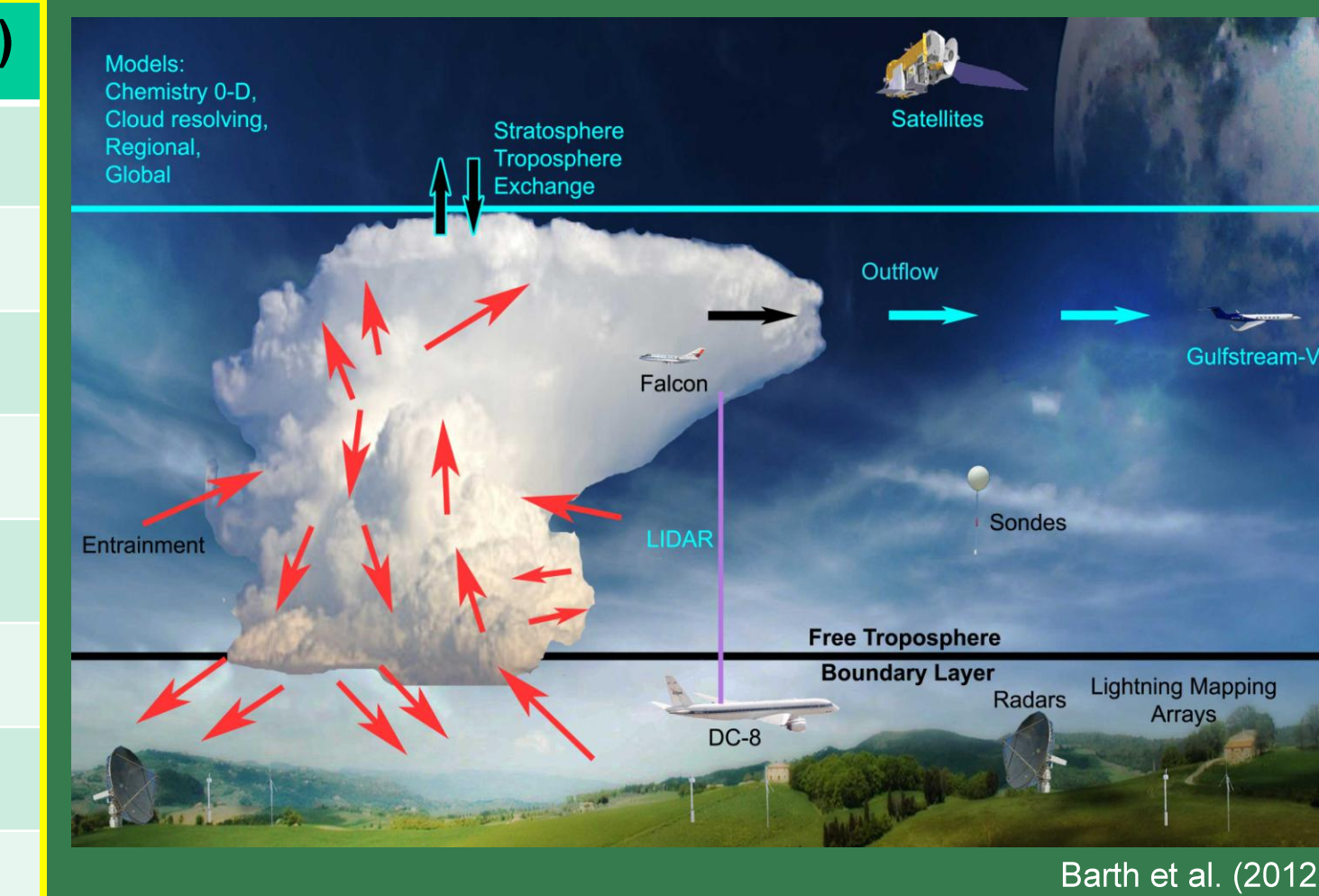
For storm maturity late mature cells produced the best results. The figure to the left shows two times an early mature (blue lines) and a late mature (black lines). The black lines have closer LMDs (circles) and there is a significant magnitude difference. This is because the ice mass LMD is dependent on the anvil which takes time to grow while the vertical mass divergence LMDs are not dependent on well developed convective anvils.



Storm Maturity Identification

To identify storm maturity, for the most representative time, the magnitude of the divergence was used. To the left is a plot from Mullendore et al. (2013) that includes a normalized divergence line (black line) for one of the cases used in this study. The time used for the June 3 storm is circled in black on the plot. The time circled in blue is the early mature time in the figure above and to the left. The black circled time is the last mature time because after that time the magnitude is lower and never rebounds.

6. May 21, 2012 Alabama DC3 Case Study



On May 21, 2012 isolated and scattered storms fired along and ahead of a cold front in northern Alabama. As part of the DC3 experiment both dual-Doppler data and aircraft chemical measurements were taken. The CSA method produced LMDs at both 4 and 8 km while a dual-Doppler based vertical mass divergence calculation found a LMD at 5.5 km with another local LMD at 9.5 km. A plot of chemical measurements for CO, O₃, NO, and NO₂ is to the left. The tropopause was at 11.8 km at this time

7. Conclusions

- All methods work well under optimal conditions:
 - At late maturity
 - For isolated cells and linear morphologies
- All methods perform well under optimal conditions but CSA method performs slightly better.
- CSA proxy method most closely predicts the divergence LMD, even when conditions not optimal.
- CSA method is an improvement on previously published proxy methods.

8. References

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9. Acknowledgements

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