

# Exploring the Differences in Deep Convective Transport Characteristics Between Quasi-Isolated Strong Convection and Mesoscale Convective Systems Using Seasonal WRF Simulations

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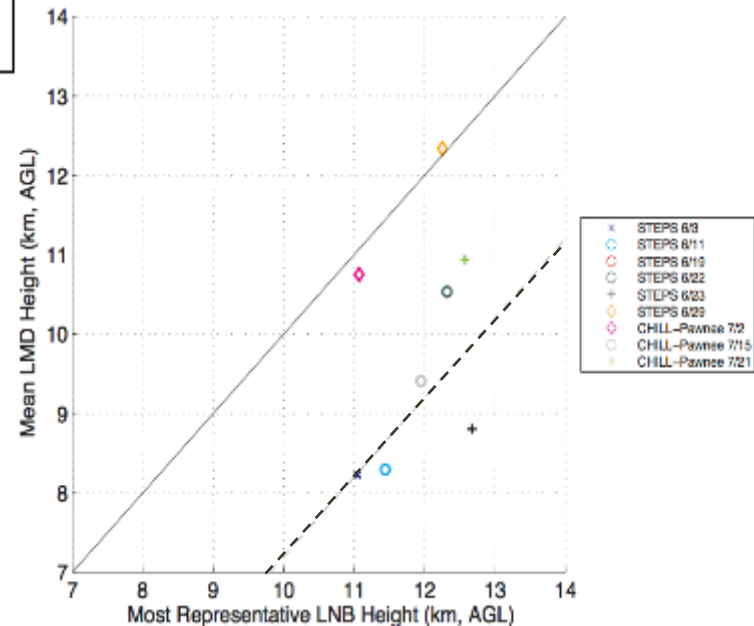
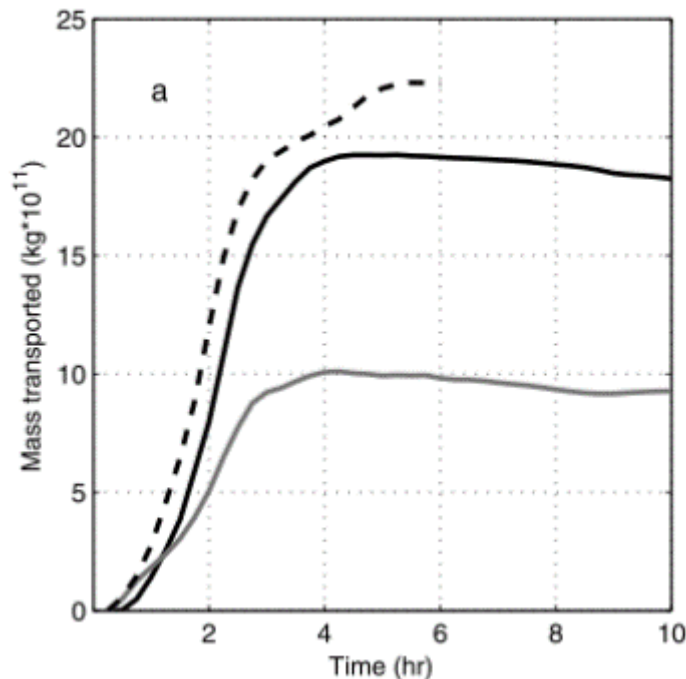


# Motivations

- Mass Transport
  - Convection is an efficient and important mechanism by which the transport of chemical constituents from the planetary boundary layer (PBL) to the upper troposphere/lower stratosphere (UTLS) region occurs (e.g. Dickerson et al. 1987; Mullendore et al. 2005; Barth et al. 2007; Lawrence and Salzmann 2008)
- Previous Studies
  - Focus of study was on either a single type of convection, or general cumulus convection (e.g. Thompson et al. 1994; Stenchikov et al. 1996; Barth et al. 2007; Halland et al. 2009)

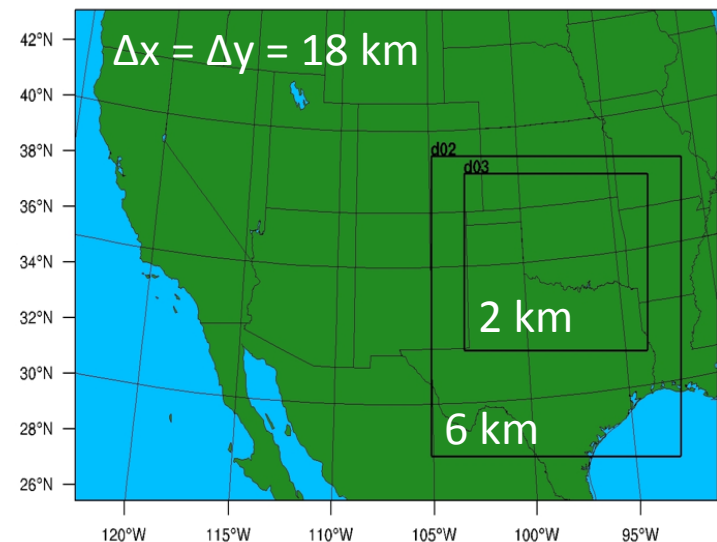
# Basis and Purpose of Study

- Differences in transport characteristics between a supercell and a multicell have been recognized (Mullendore et al. 2005; Mullendore et al. 2013).
- In terms of Magnitude...
  - Mullendore et al. (2005)
- And altitude...
  - Mullendore et al. (2013)
  - Supercells = ◆



# Model Setup

- Model
  - Weather Research and Forecasting model with chemistry (WRF-CHEM)
- Analysis Period
  - 15 – 31 May 2007 (17 days)
    - Shear zone convection
  - 01 – 13 July 2007 (13 days)
    - Sub-tropical warm sector convection
- Discretization
  - Vertical
    - Added resolution in UTLS
      - ~9-13 km
      - Equal 250 m spacing



# Thunderstorm Classification

**Green** = Agrees with Schoen and Ashley (2011)

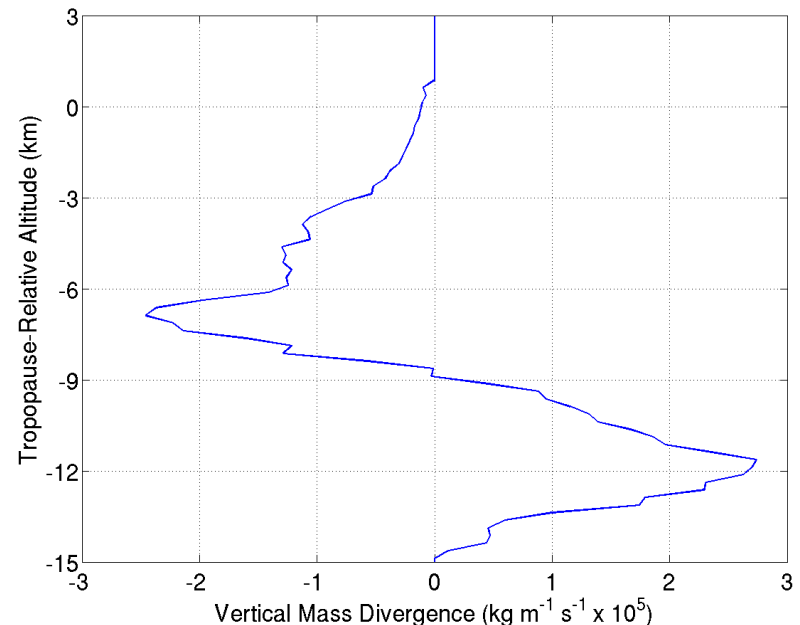
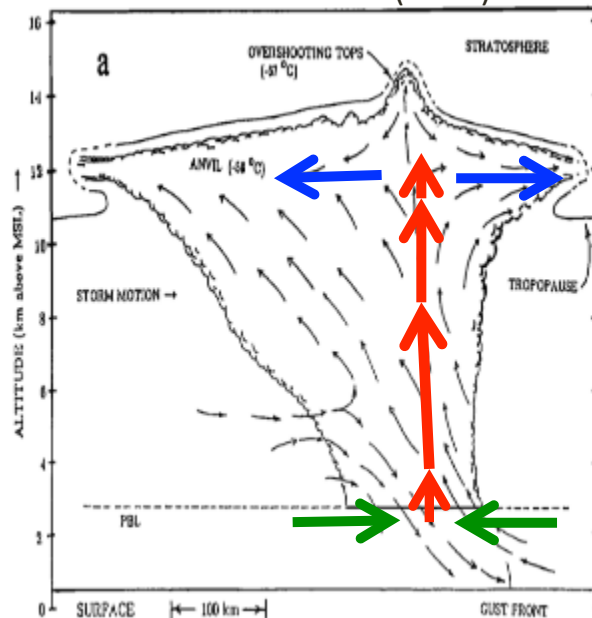
**Red** = Determined from Testing

	Weak Convection (WC)	<b>Quasi-Isolated Strong Convection (QISC)</b>	<b>Mesoscale Convective System (MCS)</b>
Radar Reflectivity Characteristics	< 40 dBZ everywhere in object	<b>≥ 40 dBZ at least 1 point in object</b>	<b>≥ 40 dBZ at least 1 point in object</b>
Areal Characteristics (Reflect. > 0 dBZ)	<b>Can be any size</b>	<b>&lt; 7000 km<sup>2</sup></b>	<b>≥ 7000 km<sup>2</sup></b>

# Deep Convection and Analysis

- Deep convection defined as  $w \geq 2 \text{ m s}^{-1}$  at  $z = 4 \text{ km}$  AND  $w \geq 5 \text{ m s}^{-1}$  at  $z = 8 \text{ km}$
- Vertical mass divergence  $[\partial(\rho w)/\partial z]$  was calculated for each column of deep convection and columns were summed together (Mullendore et al. 2009).
  - Only positive vertical velocities were used

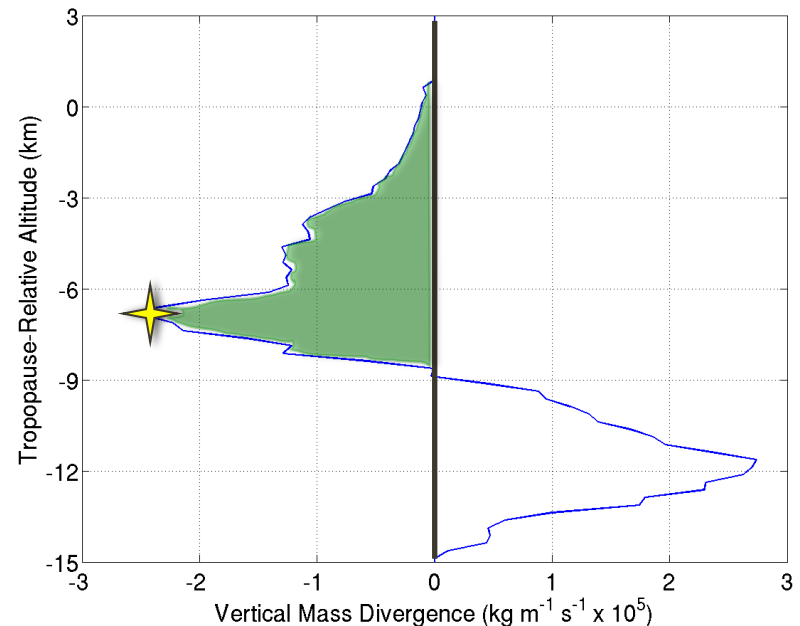
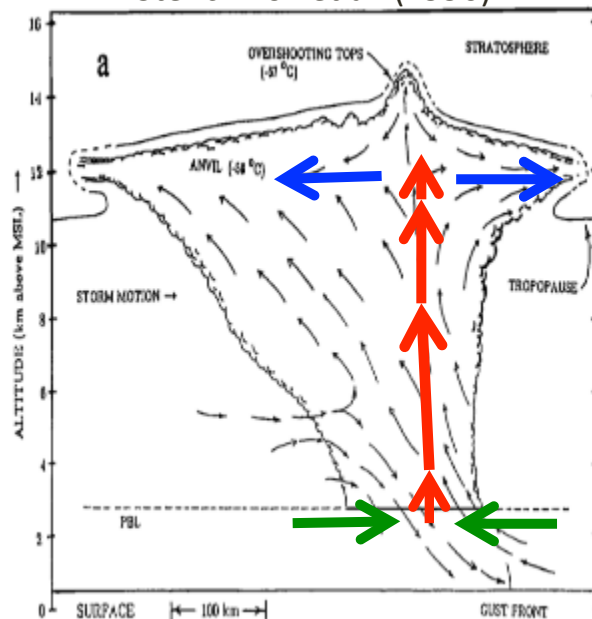
Stenchikov et al. (1996)



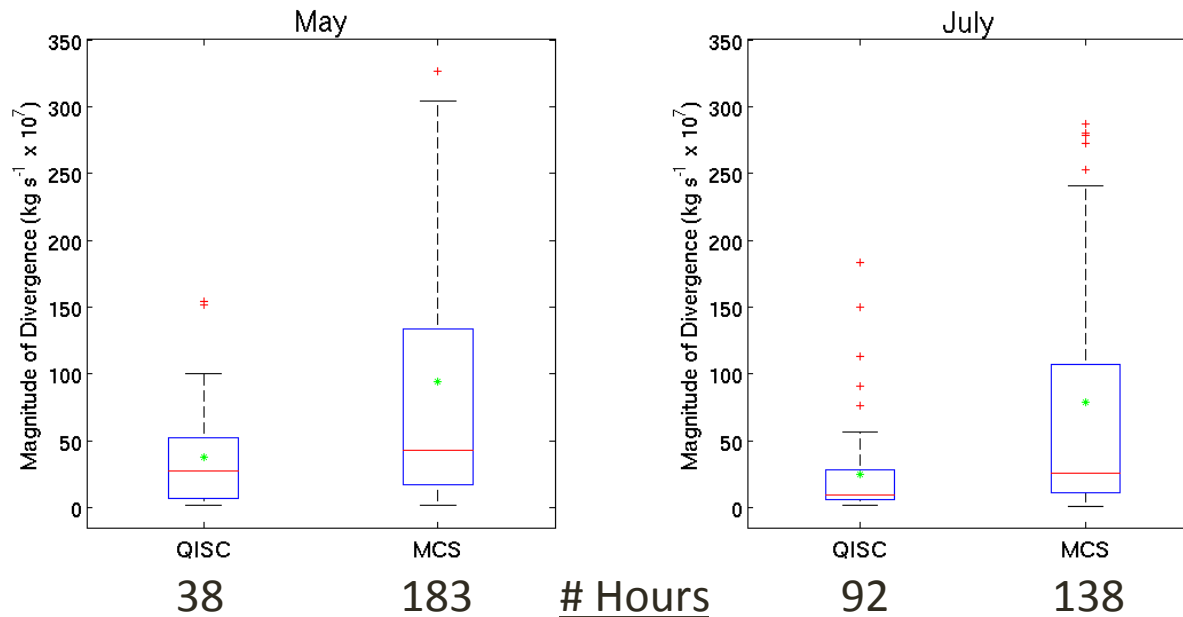
# Deep Convection and Analysis

- Analyzed vertical convergence (horizontal divergence) for analysis
  - Level of Maximum Detrainment (LMD, Mullendore et al. 2009; 2013) determined from maximum vertical convergence
    - Calculated relative to NARR tropopause heights mapped onto analysis domain
  - Magnitude of the vertical convergence calculated as proxy for amount of detrainment

Stenchikov et al. (1996)



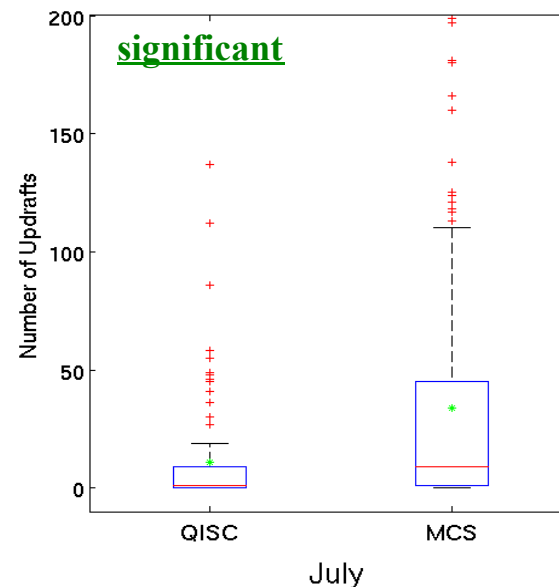
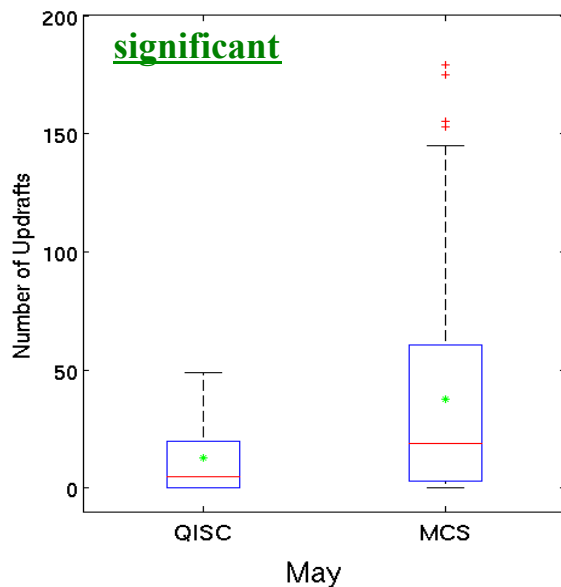
# Transport Magnitude Per Object



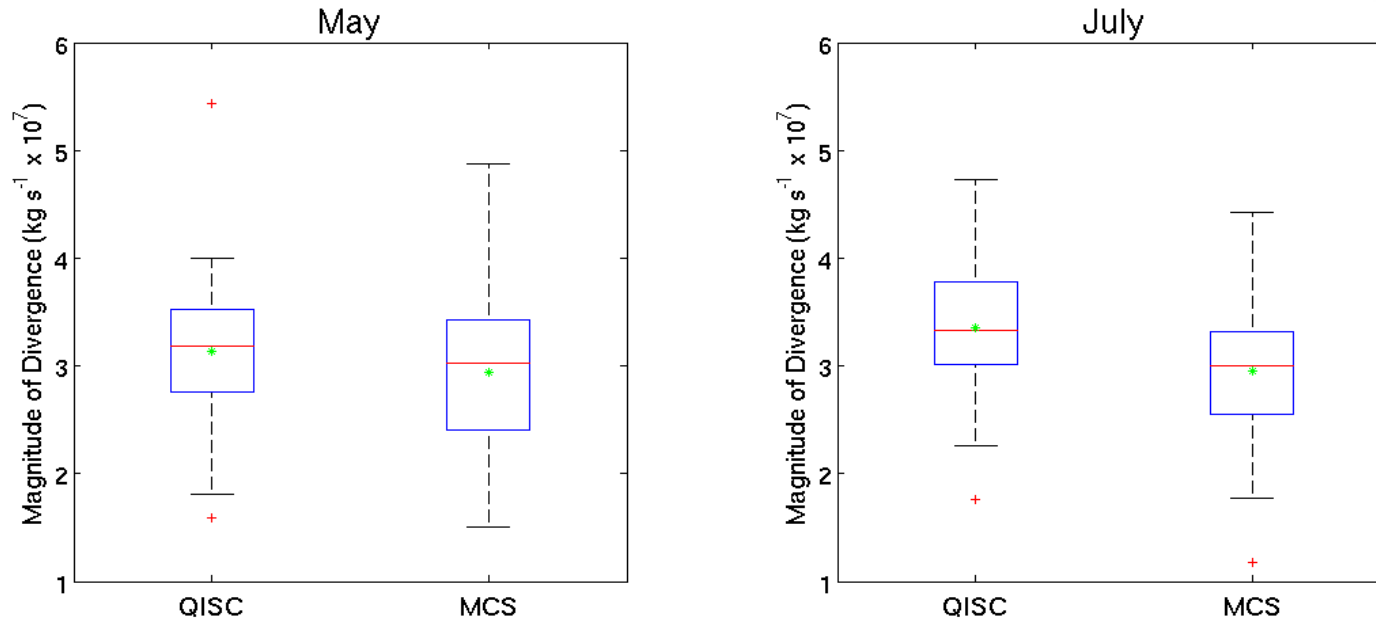
- Box and Whisker Distributions of magnitude
  - Each hour where QISC or MCS dominant represents 1 point in the plot
  - Green Asterisk – Mean value
- For Both May and July:
  - MCS detrains **statistically significantly** more mass out of the updrafts *per deeply convective complex* than QISC

# Number of Deeply Convective Updrafts

- Deep Convection Defined as:
  - $w \geq 2 \text{ m s}^{-1}$  at  $z = 4 \text{ km}$  AND  $w \geq 5 \text{ m s}^{-1}$  at  $z = 8 \text{ km}$
- There is a relationship between the amount of deep convective updrafts and the detrainment out of the updrafts per deeply convective storm complex



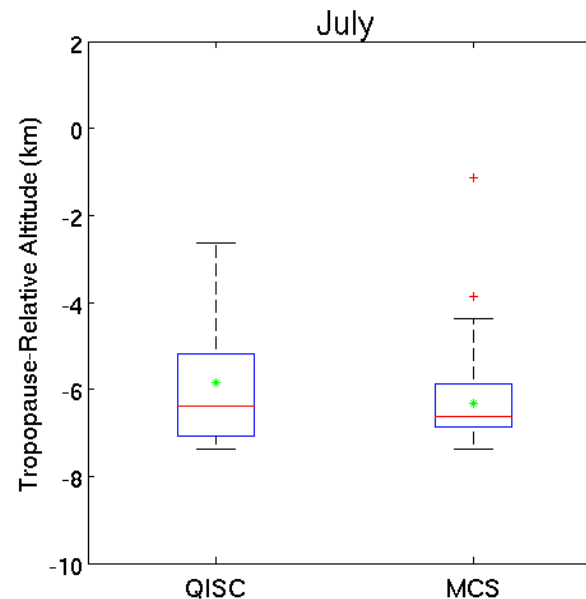
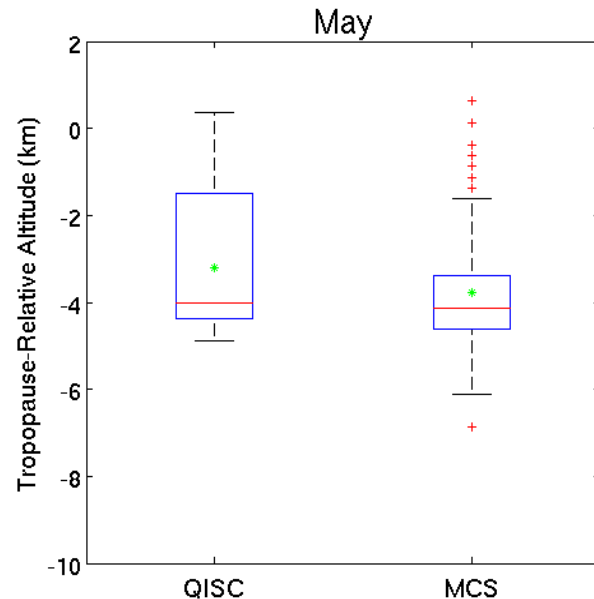
# Transport Magnitude Per Updraft



- Magnitude of Detrainment Per Updraft
- For May:
  - QISC does **not statistically significantly** detrain more mass out of the updrafts *per updraft* than MCS
- For July:
  - QISC detrains **statistically significantly** more mass out of the updrafts *per updraft* than MCS

# Transport Altitude (Tropopause-Relative)

- Due to difference in environment, tropopause height at least 2 km higher in July than in May
  - Could Lead to reason why July maxes at  $\sim -2.5$  km and May  $> 0$  km
- In both May and July, the QISC LMD, relative to the altitude of the tropopause, is **statistically significantly** higher than that of the MCS



# Results

- Summary of Results:
  - Differences are MCS – QISC
  - Ratios are MCS / QISC

	Differences and ratios of mean values of transport characteristics between MCS and QISC regimes			
	Detrainment <i>Per Deeply Convective Storm Complex</i>	# of Deeply Convective Updrafts	Detrainment <i>Per Updraft</i>	Tropopause-Relative LMD Altitude
May	2.4:1	2.9:1	0.9:1	-583 m
July	3.1:1	3.1:1	0.8:1	-487 m

# Conclusions

- The MCS regime capable of detraining  $> 2$  times the amount of mass out of updrafts per storm complex than QISC.
  - Due in part to significantly more updrafts in MCS's
- QISC updrafts, individually, are much stronger and more efficient transporters than MCS updrafts
  - Able to detrain more mass at a higher altitude relative to the tropopause
    - Better chance for QISC to have irreversible transport to the stratosphere than an MCS
- Objective classification of storms for deep convective transport studies important
  - Accounts for variability in overall transport budget due to different storm types
    - Not realized at global transport model scales (where convection parameterized)